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·《地球科学与环境学报》更名二十周年纪念专辑·

西秦岭造山带早中生代花岗岩 成分多样性及形成机制

豆敬兆^{1,2}, 何俊², 黄曦光², 陈福坤^{2*}

(1. 中国科学院广州地球化学研究所 中国科学院矿物学与成矿学重点实验室, 广东 广州 510640;

2. 中国科学技术大学 地球和空间科学学院, 安徽 合肥 230026)

摘要: 花岗质岩体产状及组成上的差异主要受控于源区性质和岩浆过程。以西秦岭造山带内几个代表性早中生代花岗岩为例, 简要介绍其形成中所记录的岩浆过程, 探讨花岗岩成分多样性的原因。西秦岭造山带内 S 型花岗岩是变泥质岩部分熔融的产物; 根据地球化学特征可分为两类, 即高 Sr 含量、低 Rb/Sr 值和稀土元素总含量的 Group A, 及低 Sr 含量、高 Rb/Sr 值和稀土元素总含量的 Group B, 其分别可由白云母水致熔融和脱水熔融形成。糜署岭岩体的暗色微粒包体中发育 3 类晶形和成分不同的锆石, 其与寄主岩石的锆石具有一致的 U-Pb 年龄; 其中, 暗色微粒包体中的类型 3 锆石与寄主岩石中锆石具有一致的晶形和 $\epsilon_{\text{Hf}}(t)$ 值, 捕获于寄主岩石, 而类型 1 和类型 2 锆石具有较高的 Th/U 值和 $\epsilon_{\text{Hf}}(t)$ 值, 结晶于混合程度不同的基性岩浆, 其基性端元可能来自幔源岩浆; 因此, 糜署岭岩体的暗色微粒包体记录了壳-幔岩浆混合过程。中川岩体由形成时代没有明显差异的 5 个岩相呈同心环状产出; 边部的似斑状花岗闪长岩被含斑中粒黑云母二长花岗岩侵入, 中粒黑云母二长花岗岩被中粒电气石二云母花岗岩侵入, 后者又被细粒黑云母二长花岗岩侵入; Rb-Sr-Ba 微量元素特征显示各岩性之间不存在成分分异趋势; 因此, 中川岩体的环带结构是由多期次岩浆侵位、聚集形成的。

关键词: 花岗岩; 地球化学; 早中生代; 成分多样性; 不一致熔融; 岩浆混合; 多期岩浆侵位; 西秦岭

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Compositional Diversity of Early Mesozoic Granites in the West Qinling Orogen, China and Their Genetic Mechanism

DOU Jing-zhao^{1,2}, HE Jun², HUANG Xi-guang², CHEN Fu-kun^{2*}

(1. CAS Key Laboratory of Mineralogy and Metallogeny, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou 510640, Guangdong, China; 2. School of Earth and Space Sciences, University of Science and Technology of China, Hefei 230026, Anhui, China)

Abstract: The occurrence and compositional diversity of granitic pluton is determined by properties of source rocks and conditions of partial melting, and can be further modified by magmatic processes such as magma mixing, crystal fractionation and assimilation during the

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作者简介: 豆敬兆(1990-), 男, 河北邢台人, 理学博士, 博士后, E-mail: jzdou@mail.ustc.edu.cn.

* 通讯作者: 陈福坤(1964-), 男, 福建龙海人, 教授, 博士研究生导师, 理学博士, E-mail: fkchen@ustc.edu.cn.

transfer. However, the role and relative importance of each process during pluton construction are still elusive. Several representative Early Mesozoic granitic plutons in the West Qinling orogen were chosen to illustrate how the magmatic processes control the compositional variation of granitic plutons. The S-type granites in the West Qinling orogen were derived from partial melting of metapelite and can be classified into two categories based on geochemical features, i. e., high Sr content, and low Rb/Sr ratio and REE content (Group A), and low Sr content, and high Rb/Sr ratio and REE content (Group B). Groups A and B might have been produced through fluid-present and fluid-absent incongruent melting of muscovite, respectively, which can well explain respective contrasting major and trace elemental characteristics, and zircon and monazite saturation temperatures. Three types of zircon were identified in the microgranular mafic enclaves (MMEs) from Mishuling pluton based on the morphology. They have consistent zircon U-Pb age but different trace elemental and Hf isotopic compositions. Type-3 zircon develops {100} prisms and has $\epsilon_{\text{Hf}}(t)$ similar to these of zircon in the host granite. Hence type-3 zircon is interpreted as xenocryst captured from the host granitic magma. Type-1 and type-2 zircons are characterized by {100} + {110} and {110} prisms, respectively. They have relatively higher Th/U ratio and $\epsilon_{\text{Hf}}(t)$ values than type-1 zircon, and may have crystallized from hybridized magmas with varying degrees of the mixing. The mafic endmember of the hybridized magma likely has mantle-derived affinity. Thus, the MMEs in Mishuling pluton record crust-mantle interaction. Zhongchuan pluton is comprised of five rock units from the periphery to the center, which mostly have indistinguishable zircon U-Pb ages. The peripheric porphyritic granodiorite is intruded by phenocryst-bearing (K-feldspar and quartz) medium-grained biotite monzogranite; medium-grained biotite monzogranite is crosscut by medium-grained two-mica tourmaline granite, which is inversely invaded by fine-grained biotite monzogranite in the center. Moreover, the correlations between contents of Rb and Ba, and contents of Rb and Sr for these rock units are not observed. Therefore, Zhongchuan pluton might have been formed by incremental assembly of at least five magma batches.

Key words: granite; geochemistry; Early Mesozoic; compositional diversity; incongruent melting; magma mixing; incremental assembly of magma batches; West Qinling

0 引言

花岗岩(广义上)是大陆地壳的重要组成部分,且常与金属矿床相伴生,具有一定的成矿专属性^[1-2],因此,深入了解花岗岩成因机制对认识大陆地壳的形成与演化,理解相关矿床成因具有重要意义^[3]。花岗岩在地壳内常以复式岩体或杂岩体的形式出露,岩体内岩石成分和结构变化多样^[4],其中的造岩矿物(如长石或角闪石)也常呈现矿物成分环带;在不同尺度上,成分和结构的不均一性与花岗质岩浆在地壳内形成、迁移、分异过程及储存条件等因素密切相关。因此,阐明花岗质岩体形成中所经历的岩浆过程有助于揭示花岗岩成分多样性的原因。

花岗质岩体的形成经历了源区岩石部分熔融、熔体抽取、岩浆上升及侵位等过程^[5]。源区岩石成

分的不均一^[6-8]、源岩的不一致熔融^[9-10]及转熔矿物的带入^[11-12]等决定着初始熔体成分。在岩浆上升、侵位过程中,同化混染围岩、分离结晶以及不同期次岩浆之间的混合将进一步改变本已复杂的初始熔体成分^[13-17]。近年来,花岗质岩浆多期次侵位机制^[18],特别是花岗岩可能代表熔体抽取后的堆晶岩^[19-23],这一新颖观点逐渐受到重视;在此模型下,花岗岩的成分变化可能与熔体的抽取程度有关,晶体-熔体分离发生的过程与机制成为当前研究的前沿领域^[24-25]。此外,学者们对花岗质岩浆在地壳内的储存状态的认识由以熔体为主的岩浆房储存逐渐转变为以晶粥形式储存^[26-27];在“晶粥储库”新的视角下,以往用于解释岩浆房过程的分离结晶、岩浆混合等作用发生的物理过程及机理可能需要重新审视^[28]。

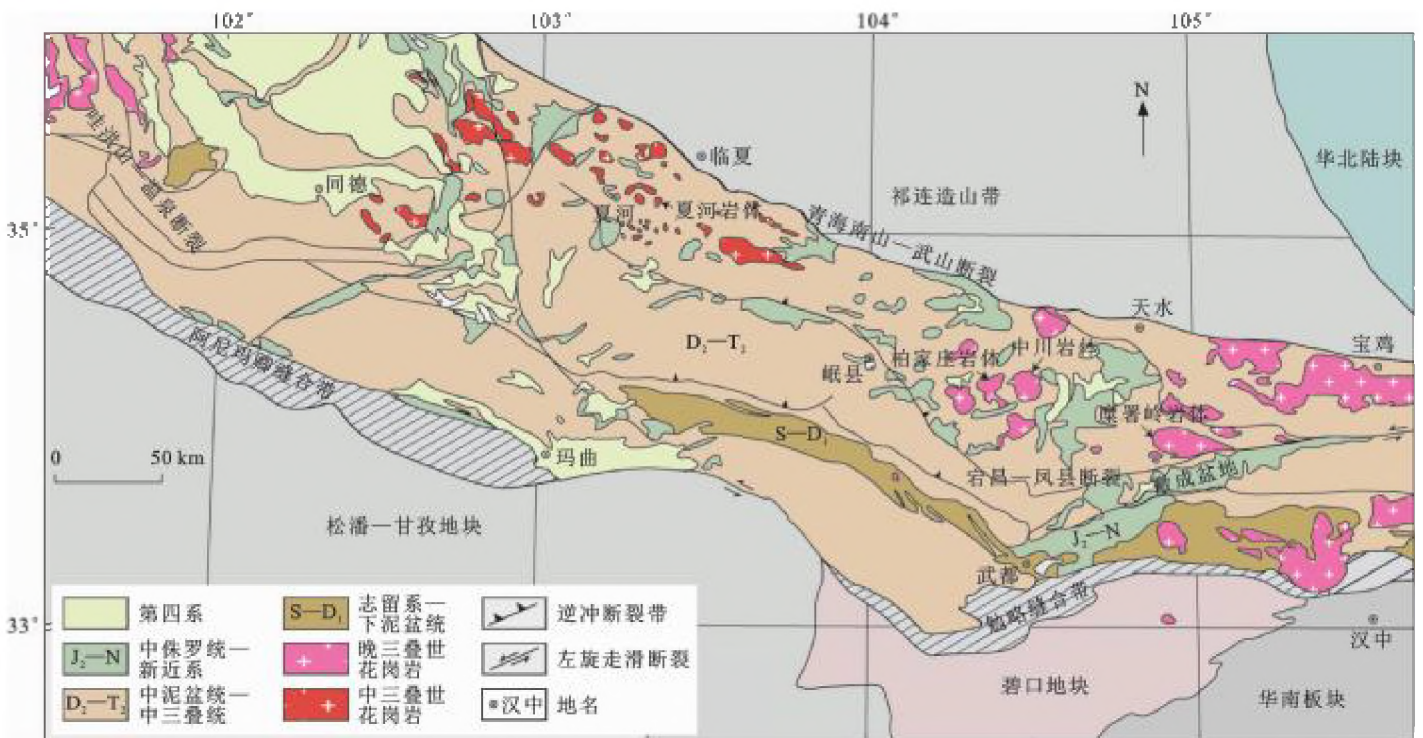
西秦岭造山带内出露大量花岗质岩体,岩石类型及成分多样。前人对该区域的花岗岩开展了大量年代学和地球化学研究,在岩石成因及形成的构造环境等方面做了大量奠基性工作,这为进一步研究花岗岩成分多样性的原因提供了绝佳场所。本文以西秦岭造山带内出露的几个代表性早中生代花岗岩为研究对象,通过解剖和总结典型岩体成因,结合近年来花岗岩成因理论研究进展,对源区部分熔融条件、岩浆混合及岩浆多期侵位等岩浆过程如何影响花岗质岩体成分变化加以探讨,以期抛砖引玉。

1 区域地质背景

西秦岭造山带位于古亚洲洋构造域、特提斯构造域及滨太平洋构造域的接合地带^[29-30]。其南以勉略—阿尼玛卿缝合带与扬子陆块、松潘—甘孜地块及碧口地块相隔,东以徽成盆地与东秦岭造山带相连,北临祁连山造山带和华北陆块,西接柴达木地块和昆仑造山带(图 1)。西秦岭造山带内广泛发育显生宙地层,主要包括泥盆系、石炭系、二叠系及三叠系,岩性以砂岩、板岩及灰岩为主。区域内早中生代岩浆岩广泛出露,以花岗闪长岩、黑云母二长花岗岩为主。西秦岭中—东段的花岗质岩体多呈面状、椭圆状发育,如温泉、中川、教场坝、间井、碌础坝、糜署岭等岩体;岩体从边缘至中心,岩石结构、矿物组合

及成分常呈现规律变化,即矿物粒度逐渐变小,岩石 SiO₂ 含量逐渐升高;岩体边缘相角闪石发育,而中心相一般不发育角闪石,可能出现白云母、电气石等矿物;暗色微粒包体也主要分布在边缘相中。西秦岭西段的岩体呈 NW 向带状分布,如美武、夏河、同仁等岩体。西秦岭造山带内早中生代火山岩出露较少,主要分布在宕昌、德乌鲁和麦秀等地区,形成时代为 246~229 Ma^[31-34]。此外,西秦岭造山带西段出露少量中三叠世中性、基性和超基性岩,包括闪长岩、辉长岩、纯橄榄岩、异剥橄榄岩、橄榄辉石岩等^[35-36]。

前人对西秦岭造山带内的花岗岩开展了大量年代学、岩石学和地球化学研究,取得的主要认识有:①花岗岩时空分布特征显示,西秦岭中—东段和西段的花岗岩侵位时代分别为 220~210 Ma 和 250~240 Ma^[37-40];②早中生代花岗岩具有高钾钙碱性特征,以 I 型花岗岩为主^[37-39],并有少量岩石表现出埃达克质岩石特征^[41-42];③花岗岩中暗色微粒包体是岩浆混合的产物,记录了壳—幔相互作用^[43-46]。前人对于西秦岭造山带内早中生代花岗岩形成的构造环境也有不同看法。有学者认为早中生代花岗岩形成于同碰撞环境^[47]或与勉略洋向北俯冲有关^[48];另有学者认为 250~240 Ma 花岗岩的形成与阿尼玛卿—勉略洋向北俯冲有关,220~210 Ma 花岗岩则



图件引自文献[30],有所修改

图 1 西秦岭造山带区域地质简图

Fig. 1 Regional Geological Sketch Map of the West Qinling Orogen

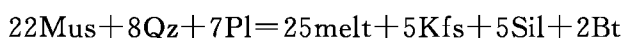
形成于后碰撞环境^[37-39]。

2 早中生代花岗岩记录的岩浆过程

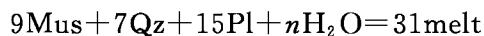
2.1 源岩不一致熔融过程

地壳内花岗质原生岩浆常通过源区岩石中矿物不一致熔融的方式产生^[9,49]。在部分熔融过程中,若熔体形成及抽取的速率大于元素扩散的速率,初始熔体的成分并非由源岩中矿物所决定,而取决于参与部分熔融的矿物相种类及比例,并受转熔矿物相的影响^[11]。在变泥质岩部分熔融过程中,存在两种反应类型^[50],分别为白云母脱水熔融和水致熔融。

白云母脱水熔融反应式为



白云母水致熔融反应式为



式中:Mus 代表白云母;Qz 代表石英;Pl 代表斜长石;Kfs 代表钾长石;Sil 代表矽线石;Bt 代表黑云母;melt 代表熔体。

由上述反应式可以看出,在脱水部分熔融过程中,参与反应的白云母较多,而斜长石较少,并伴有转熔钾长石的形成;由于 Rb、Sr 和 Ba 主要分别赋存在白云母、斜长石和钾长石中,白云母脱水部分熔融所产生的熔体具有较高的 Rb/Sr 值以及低 Sr 含量、Sr/Ba 值;而白云母水致熔融过程中,由于更多斜长石的参与,形成的熔体则具有相对低的 Rb/Sr 值以及高 Sr 含量、Sr/Ba 值。因此,变泥质岩部分熔融形成的花岗质原生岩浆的 Rb-Sr-Ba 关系常被用来推测部分熔融方式。这两种部分熔融反应类型也被认为是喜马拉雅造山带内淡色花岗岩的形成机制^[51-53]。

西秦岭造山带内出露少量三叠纪白云母或二云母花岗岩,如柏家庄、蒋家坪、光头山等岩体,它们被认为是变沉积岩部分熔融形成的 S 型花岗岩^[54-56]。这些岩体具有较低的 CaO/Na₂O 值[图 2(a)],暗示其可能是源区变泥质岩部分熔融的产物^[57]。地球化学特征显示,西秦岭造山带内 S 型花岗岩大致可分为两类:一类具有低 Rb/Sr 值、高 Sr 和 Ba 含量特征,称为 Group A;另一类具有高的 Rb/Sr 值、低 Sr 和 Ba 含量特征,称为 Group B。这两类分别符合白云母水致熔融和脱水熔融趋势[图 2(b)、(c)]。与白云母脱水熔融相比,白云母水致熔融导致源区斜长石相较于白云母熔解得更多,因此,所产生的熔体具有高 CaO、Sr 含量以及高 Eu 异常[图 2(a)、

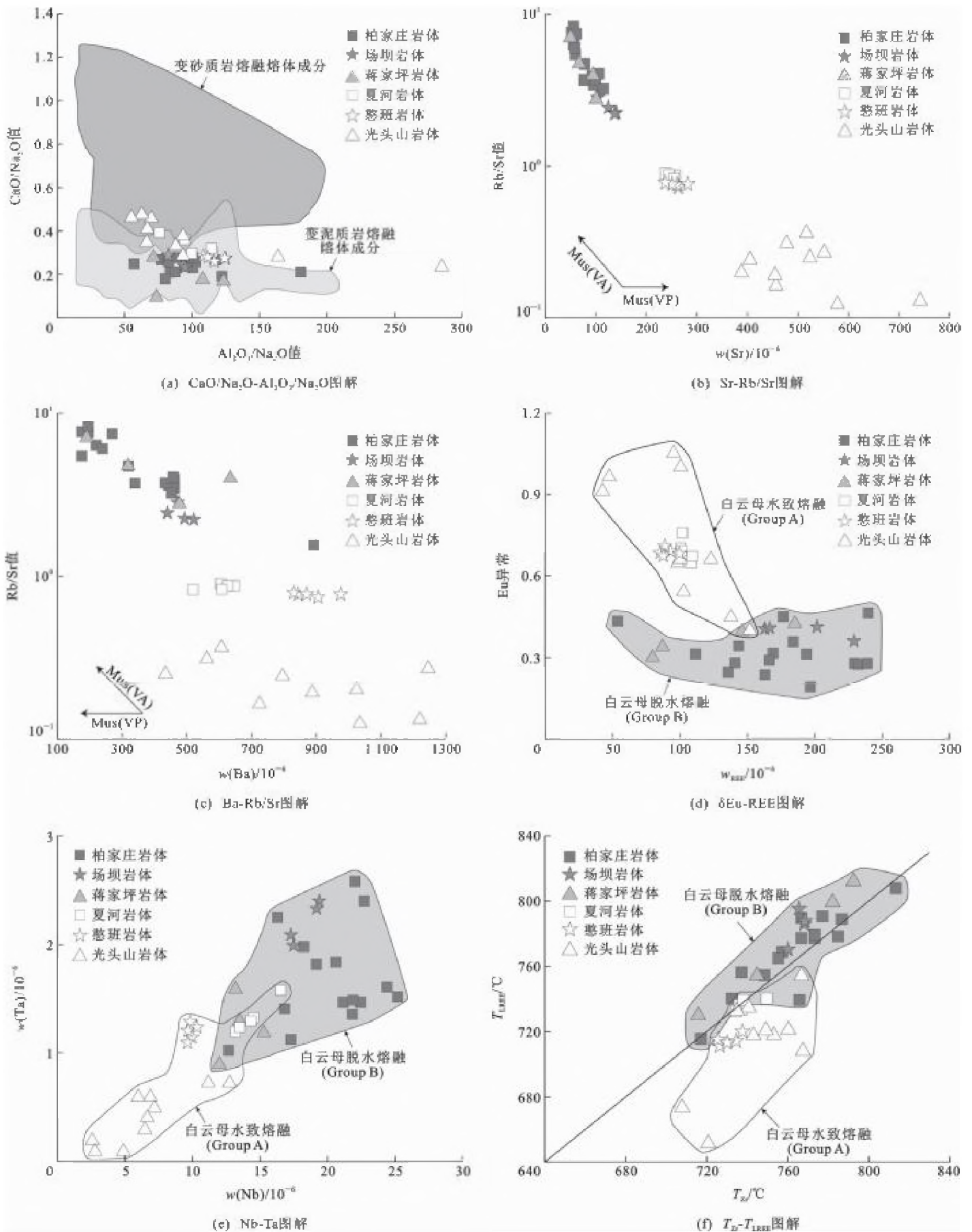
(b)、(d)]。白云母为 Nb、Ta 的载体矿物,在白云母脱水熔融过程中,白云母比斜长石熔解得更多,且产生转熔钾长石,因此所形成的熔体具有较高 Nb、Ta 含量,但较低 Eu 异常[图 2(d)、(e)]。白云母水致熔融过程由于有流体参与,所以具有相对较低的固相线温度;锆饱和温度和独居石饱和温度计算结果显示,Group A 比 Group B 具有更低的温度[图 2(f)],符合上述特征。而由于白云母脱水熔融具有相对高的温度,源区可能有更多磷灰石、独居石溶解,这解释了 Group B 比 Group A 具有更高的稀土元素总含量[图 2(d)]。因此,西秦岭造山带内两类 S 型花岗岩主量、微量元素成分变化应当记录了变泥质岩源区白云母水致熔融和脱水熔融过程。

2.2 壳-幔岩浆混合作用

岩浆混合作用在花岗岩形成过程中扮演着重要的角色。由于酸性岩浆之间的成分及黏度等方面差异比基性—酸性岩浆之间更小,故更易发生混合^[58],然而酸性岩浆之间的混合现象较难识别^[16,59]。因此,有关岩浆混合的研究主要关注基性—酸性岩浆之间的混合过程,尤以花岗岩中暗色微粒包体的研究居多,暗色微粒包体常被认为是壳幔岩浆混合的产物^[60-62]。西秦岭造山带内除上述少量 S 型花岗岩外,出露众多具有 I 型花岗岩特征的花岗闪长岩和二长花岗岩,其中包含丰富的暗色微粒包体。已报道的数据显示,西秦岭造山带内暗色微粒包体与寄主岩石具有相似的结晶年龄、全岩 Sr-Nd-Hf 同位素和锆石 Hf-O 同位素特征^[45,63-65],因此,这些暗色微粒包体也被认为是早期堆晶体或与寄主岩石相似岩浆源区的壳源包体^[63,65-67]。

镜下观察显示,西秦岭造山带内的暗色微粒包体具有典型的火成岩结构,黑云母、钾长石及辉石结晶共生,并发育针状磷灰石等淬火结构[图 3(a)],明显不同于堆晶结构。野外露头上观察到暗色微粒包体与寄主岩石相互包裹,并包含有来自寄主岩石的钾长石斑晶,显示寄主岩浆与暗色微粒包体岩浆在晶粥状态下相互作用[图 3(b)]。这些现象说明西秦岭造山带内花岗岩中的暗色微粒包体可能记录了岩浆混合过程。那么这些暗色微粒包体的原始岩浆是壳源还是幔源呢?

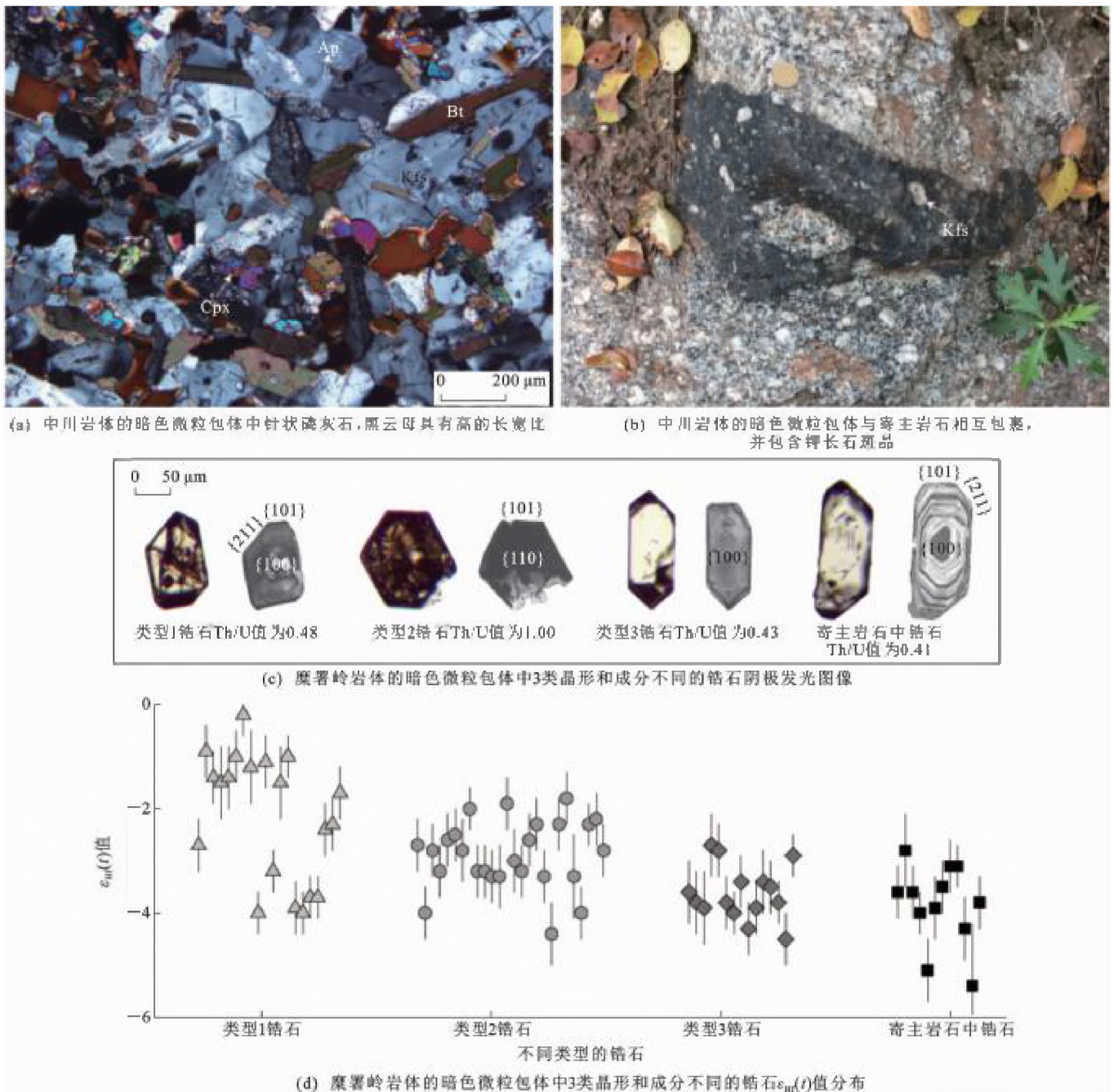
在糜署岭岩体的暗色微粒包体中,识别出 3 类晶形和成分明显不同的锆石[图 3(c)]。类型 1 锆石呈现浅棕色,发育{100}、{110}柱面和{101}、{211}锥面;类型 2 锆石呈现深棕色,并以{110}柱面和{101}锥面为特征;类型 3 锆石呈现浅黄色,发育



Mus(VA)代表白云母脱水熔融; Mus(VP)代表白云母水致熔融; $w(\cdot)$ 为元素或化合物含量; w_{REE} 为稀土元素总含量; T_{Zr} 和 T_{LREE} 分别代表锆石和独居石饱和温度; 图件引自文献[54]

图 2 西秦岭造山带两类 S 型花岗岩地球化学特征图解

Fig. 2 Diagrams of Geochemical Characteristics of Two Types of S-type Granite in the West Qinling Orogen



类型1、类型2、类型3锆石来自暗色微粒包体;图(c)中, $\{211\}$ 、 $\{101\}$ 、 $\{100\}$ 、 $\{110\}$ 为锆石的不同晶面;Cpx为单斜辉石;Ap为磷灰石;Kfs为钾长石;Bt为黑云母;图(c)、(d)引自文献[46]

图3 西秦岭造山带花岗质岩体内暗色微粒包体的显微结构、野外地质特征及其锆石类型

Fig. 3 Microphotographs, Field Geological Characteristics of Microgranular Mafic Enclaves of the Granitic Plutons and Their Zircon Types in the West Qinling Orogen

$\{100\}$ 柱面和 $\{101\}$ 、 $\{211\}$ 锥面,表现出与寄主岩石的锆石一致的形貌学特征。基性包体和寄主岩石中的锆石颗粒均发育明显的岩浆振荡环带,并具有一致的结晶年龄(215 Ma),与区域上广泛发育的晚三叠世岩浆活动同期。类型1和类型2锆石的Th/U值为0.47~1.63, Hf含量小于 10^{-6} ,与基性岩浆中结晶的锆石具有相似的成分特征^[68],暗示它们结晶于基性岩浆;而类型3锆石具有相对低的Th/U值(0.35~0.53),可能结晶于花岗质岩浆。按照类型1、类型2、类型3锆石的顺序,其 $\epsilon_{Hf}(t)$ 值

依次降低,且类型3锆石与寄主岩石中的锆石具有相似的晶形和一致的Hf同位素组成[图3(c)、(d)]。以上特征显示,类型1和类型2锆石结晶于混合程度不同的基性母岩浆,其基性端元可能为幔源岩浆,而类型3锆石可能结晶于花岗质岩浆,之后又被基性岩浆捕获,进入暗色微粒包体。上述3类锆石所表现的不同晶形是其结晶时所对应的过冷却度环境不同所致。研究表明,在高过冷却度条件下,锆石的 $\{110\}$ 柱面比 $\{100\}$ 柱面更为发育^[69],类型2、类型1和类型3锆石依次分别发育 $\{110\}$ 、 $\{110\}$ +

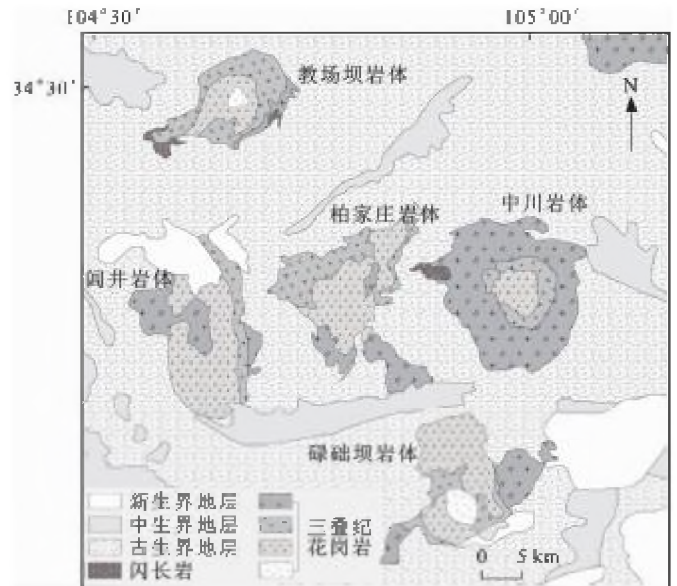
{100}、{100}柱面,表明它们结晶时的过冷却度依次降低。因此,糜署岭岩体暗色微粒包体中的锆石记录了壳-幔岩浆混合作用,强烈的混合过程可能导致了暗色微粒包体与寄主岩石具有相似的全岩 Sr-Nd-Hf 同位素组成。区域上其他花岗质岩体的暗色微粒包体可能也记录了类似过程,但仍有待研究。

2.3 多期次岩浆累积侵位

花岗岩在地壳中多以岩体或岩基形式出露,因此,理解岩体的形成过程有助于认识大陆地壳的生长方式^[70]。有些花岗质岩体表现出环带结构,即从岩体边缘相至中心相,岩石的结构和化学成分呈现有规律的变化;阐明岩体的环带结构成因可以帮助认识岩体的形成机制。从野外地质观察来看,暗色基性包体在环带花岗质岩体中普遍发育,因此,岩浆混合被认为是塑造环带花岗质岩体成分变化的重要原因^[71-72]。然而锆石 ID-TIMS 高精度定年结果显示岩体形成时间跨度可达 10 Ma,远远长于岩浆冷却的时间(约 1 Ma)^[18,73-74],这表明岩浆房可能是由多期次岩浆侵位聚集形成的;各期次岩浆的成分特征可能在深部源区已经确定^[18],且在上升侵位过程中,彼此之间并未出现强烈的相互作用^[75]。此外,相平衡模拟表明,在侵位水平上,晶粥的压实作用引起的熔体抽取过程也可以形成岩体的环带结构^[14]。综上所述,花岗质岩体形成中可能涉及到多种岩浆过程,然而各个岩浆过程在岩体形成中扮演的角色却不清楚^[76-77]。揭示花岗质岩体环带结构形成中的岩浆过程,能够帮助深入理解花岗岩成分多样性的原因。在西秦岭造山带内,晚三叠世花岗质岩体呈现典型的环带结构^[78](图 4),本文以中川岩体为例,在前人研究成果的基础上,结合野外地质特征探讨岩体的形成过程。

中川岩体环带结构的形成机制目前有不同看法。李宏卫等认为中川岩体是地壳两次重熔的结果^[79]。暗色微粒包体在岩体边部大量出现,岩浆混合在中川岩体形成中有着重要作用^[44],被认为是导致岩体同心环带结构的重要原因^[80]。柯昌辉等认为从岩体边缘相至中心相,岩浆混合和分异的程度增强^[81]。

详细的野外地质调查显示,中川岩体包含 5 种岩石类型:似斑状花岗闪长岩、含斑中粒黑云母二长花岗岩、中粒黑云母二长花岗岩、中粒电气石二云母花岗岩及细粒黑云母二长花岗岩。暗色微粒包体主要发育在似斑状花岗闪长岩中,在含斑中粒黑云母二长花岗岩中零星出现。野外观察发现,含斑中粒



图件引自文献[78],有所修改

图 4 西秦岭造山带东段花岗质岩体的环带结构
Fig. 4 Zoned Pattern of the Granitic Plutons in the Eastern West Qinling Orogen

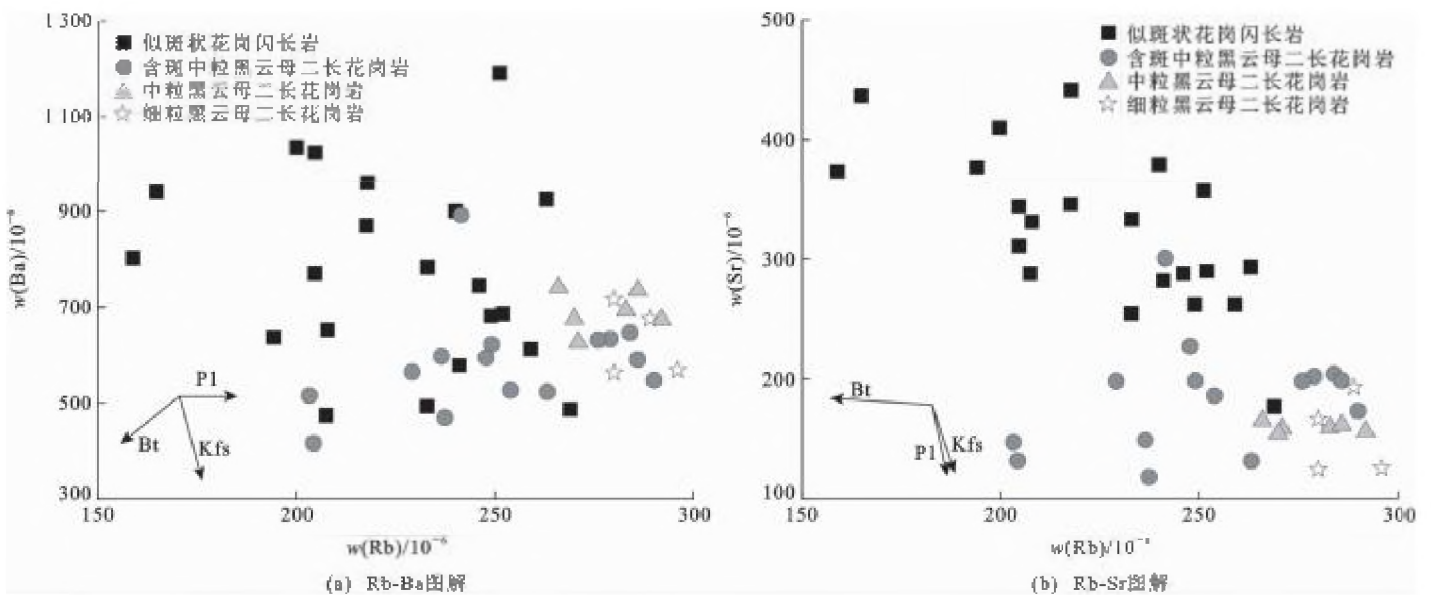
黑云母二长花岗岩包含似斑状花岗闪长岩捕虏体[图 5(a)],中粒电气石二云母花岗岩侵入似斑状花岗闪长岩和中粒黑云母二长花岗岩[图 5(b)、(c)],细粒黑云母二长花岗岩包含中粒电气石二云母花岗岩捕虏体[图 5(d)]。这些侵入关系表明,中川岩体形成过程中涉及到至少 5 期岩浆侵位。全岩微量元素特征显示,似斑状花岗闪长岩、含斑中粒黑云母二长花岗岩、中粒黑云母二长花岗岩和细粒黑云母二长花岗岩之间并没有分离结晶作用的趋势(图 6),反映它们并非是一期次岩浆连续分异的结果,而代表不同期次侵位的岩浆。暗色微粒包体主要分布在岩体边部的似斑状花岗闪长岩中,因此,岩浆混合在中川岩体较早期次岩浆(即似斑状花岗闪长岩)的形成中有着重要作用,但在中川岩体环带结构形成中的角色仍需进一步评估。中川岩体不同岩石单元的岩浆起源仍需同位素证据予以揭示。

在误差范围内,似斑状花岗闪长岩、含斑中粒黑云母二长花岗岩、中粒黑云母二长花岗岩和细粒黑云母二长花岗岩具有一致的锆石 U-Pb 年龄(约 215 Ma, LA-ICPMS 法),且野外地质观察显示,中粒电气石二云母花岗岩呈塑性变形被包裹在细粒黑云母二长花岗岩中。这些现象表明,上述 5 种岩石类型可能近同时侵位。考虑 2.5% 的分析误差,对应约 5 Ma,中川岩体不同期次岩浆侵位在 5 Ma 内也已完成,这与 Glazner 等通过岩体热模拟所获得的认识^[73]是一致的。图 7 简要地描绘了中川岩体的形成过程。



图5 中川岩体不同岩石单元之间的野外侵入关系

Fig. 5 Intrusive Contacts of Different Rock Units in Zhongchuan Pluton



Kfs为钾长石; Bt为黑云母; Pl为斜长石; 数据引自文献[82]

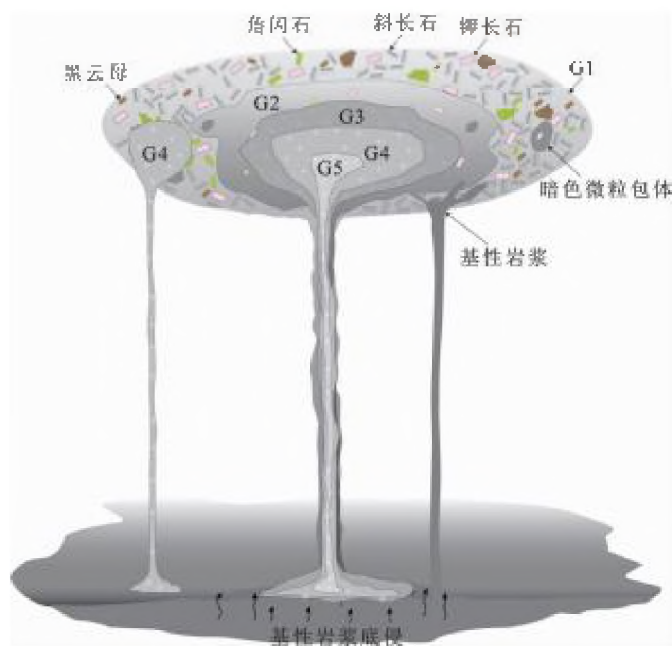
图6 中川岩体不同岩石单元的 Rb-Ba 图解和 Rb-Sr 图解

Fig. 6 Diagrams of Rb-Ba and Rb-Sr for Different Rock Units in Zhongchuan Pluton

3 结 语

(1) 西秦岭造山带内花岗岩记录了源区不一致

熔融、岩浆混合、多期岩浆侵位等多种岩浆过程。两类S型花岗岩明显不同的地球化学特征可能反映了变泥质岩源区白云母脱水熔融和水致熔融过程; 由



G1 为似斑状花岗闪长岩;G2 为含斑中粒黑云母二长花岗岩;
G3 为中粒黑云母二长花岗岩;G4 为中粒电气石二云母花岗岩;
G5 为细粒黑云母二长花岗岩

图 7 中川岩体多期次岩浆侵位模式示意图

Fig. 7 Conceptual Model View of Multi-stage Magmatic Emplacement of Zhongchuan Pluton

于白云母脱水熔融会形成钾长石等转熔矿物,西秦岭造山带内白云母脱水部分熔融形成的 S 型花岗岩(如柏家庄岩体)是否存在转熔成因钾长石,以及其与岩浆成因钾长石如何区别,仍需进一步研究。

(2)西秦岭造山带内花岗岩的暗色微粒包体是壳-幔岩浆相互作用的产物。野外地质证据显示,暗色微粒包体与寄主岩石在晶粥状态下发生过相互作用;这种过程如何影响包体成分特征以及包体在多大程度上保留幔源岩浆信息,仍然没有很好的评估。西秦岭造山带内的环带花岗质岩体是多期次岩浆侵位、聚集形成的;不同期次岩浆之间相互作用的程度,以及每一期次岩浆是否由多期岩浆聚集形成,仍然需要进一步矿物学工作予以揭示。

(3)野外地质观察显示,西秦岭造山带的花岗质岩体显示晶粥储存的特点,如丰富的钾长石斑晶聚集现象,岩石中暗色微粒包体的塑性变形等;在晶粥岩浆系统中,矿物-熔体-流体之间的反应如何塑造花岗岩成分变化仍是未来研究的重点。

(4)由于花岗质岩体是地壳深部岩浆储库的代表,以上问题的解决不仅有助于认识花岗岩体成分变化及形成机制,对理解地壳岩浆系统的演化也具有一定的参考意义。

陈福坤:非常感谢主编彭建兵院士的约稿!欣闻《地球科学与环境学报》迎来更名二十周年,在此

表示衷心的祝贺!贵刊在主编彭建兵院士的带领下,聚焦国际地学研究前沿,注重报道地学领域内交叉学科的最新成果,具有鲜明的办刊特色,吸引了大批学者专家聚此交流。贵刊精心谋划的一系列富有特色的专辑,提高了期刊在国内学术界的知名度和影响力。衷心祝愿《地球科学与环境学报》在新的征程中坚持办刊特色,成为学者们喜爱的刊物,为促进地球科学和环境领域发展作出更大的贡献!

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