

Tectonic evolution of the Proto-Qiangtang Ocean and its relationship with the Palaeo-Tethys and Rheic oceans



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Abstract: An evaluation of the potential geodynamic connections between the evolution of Paleozoic oceans in NW Gondwana and NE Gondwana is challenging. Until recently, most syntheses emphasized only two Paleozoic oceans (the Proto-Tethys and the Palaeo-Tethys) in the east Tethys realm. However, the discovery of early Paleozoic ophiolites along Palaeo-Tethys sutures located south of Proto-Tethys sutures challenges these traditional views. After a comprehensive review of relevant early Paleozoic tectonomagmatic events, we herein recognize and propose a model for the tectonic evolution of a hitherto unrecognized early Paleozoic ocean, which we call the Proto-Qiangtang Ocean. This ocean was short lived; it opened in the late Cambrian, began to subduct in the Middle Ordovician, and closed diachronously westwards between the Late Ordovician and the middle Silurian. Its closure by middle Silurian time indicates that was a spatially and temporally distinct ocean from the Palaeo-Tethys Ocean. The early tectonic evolution of the Proto-Qiangtang Ocean shares many characteristics with that of the Rheic Ocean. Both opened in the late Cambrian in the back-arc region of the Iapetus–Proto-Tethys Ocean, and the Proto-Qiangtang Ocean is considered to represent the eastern extension of the Rheic Ocean. This correlation has important implications for the Paleozoic tectonic evolution and palaeogeography of northern Gondwana.

Supplementary material: Compiled tectono-magmatic ages relevant to the evolution of the Proto-Qiangtang Ocean are available at <https://doi.org/10.6084/m9.figshare.c.6263311>

Several successive oceans were generated in the late Neoproterozoic and Paleozoic outboard of the northern Gondwanan margin, and their evolution profoundly influences Paleozoic palaeogeographical reconstructions. The vestiges of the Iapetus and Rheic oceans in NW Gondwana, and the Proto-Tethys and Palaeo-Tethys oceans in NE Gondwana are the most widely documented. The Iapetus and Proto-Tethys oceans opened in the late Neoproterozoic and closed in the Silurian (van Staal *et al.* 1998, 2009; Murphy *et al.* 2010; Li *et al.* 2018). The Rheic Ocean opened in the late Cambrian–Early Ordovician and closed in the Devonian and Carboniferous (van Staal *et al.* 2009, 2012; Murphy *et al.* 2010; Nance *et al.* 2010; Kroner and Romer 2013; Wu *et al.* 2021). The Palaeo-Tethys Ocean opened in the Silurian–Devonian and closed in the Triassic (Metcalf 2013; X.Z. Zhang *et al.* 2016;

Dan *et al.* 2018a; W. Xu *et al.* 2020). There have been many attempts to compare the geodynamic evolution of these oceanic tracts and to deduce their palaeogeography in the Paleozoic. There is a broad consensus that the Iapetus and Proto-Tethys oceans represent a continuous oceanic tract (Stampfli and Borel 2002; Torsvik and Cocks 2017; Li *et al.* 2018); however, the identity of the eastern extension of the Rheic Ocean and its potential relationship with the Palaeo-Tethys Ocean is hotly debated (e.g. Stampfli and Borel 2002; Metcalfe 2013; Stampfli *et al.* 2013; Torsvik and Cocks 2017). The main reason for this controversy is that the complex evolution of these oceans has been overprinted by events associated with the evolution of the younger Neo-Tethys Ocean and other marginal seas.

In recent years, a few early Paleozoic ophiolites have been discovered in the east Tethys realm

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along the Palaeo-Tethys sutures that lie to the south of sutures associated with the closure of the Proto-Tethys Ocean (B.D. Wang *et al.* 2013; Zhai *et al.* 2016; Nguyen *et al.* 2019). Two different models have been proposed to account for these early Paleozoic ophiolites. Some researchers consider these ophiolites to reflect the opening of the Palaeo-Tethys Ocean in the late Cambrian (B.D. Wang *et al.* 2013; Zhai *et al.* 2016). Alternatively, others have proposed the existence of an early Paleozoic ocean, located south of the Proto-Tethys Ocean in the east Tethys realm, that closed during the Silurian (H.T. Tran *et al.* 2014; X.Z. Zhang *et al.* 2014; Faure *et al.* 2018; T.V. Tran *et al.* 2020; W. Xu *et al.* 2020). In this paper, we review the Paleozoic tectonic framework in east Tethys and the tectonomagmatic events thought to be related to this early Paleozoic ocean, here called the Proto-Qiangtang Ocean, and briefly compare its geodynamic evolution with those of the Palaeo-Tethys and Rheic oceans.

Geological framework in eastern Asia

Multiple suture zones, including those that reflect the closure of the Proto-, Palaeo- and Neo-Tethys oceans, divide eastern Asia into many small terranes or blocks. These suture zones are uniformly younger from north to south: that is, the Silurian Proto-Tethys sutures in the Qinling–Qilian–Kunlun Orogen, the Triassic Palaeo-Tethys sutures in northern Tibet and adjacent regions; and the Cretaceous–Paleogene Neo-Tethys sutures in southern

Tibet and adjacent regions (Fig. 1) (Metcalf 2013). The terranes south of the Palaeo-Tethys sutures are collectively called Cimmerian microcontinents (Şengör 1979), and include the Karakorum, Southern Qiangtang, Amdo, Lhasa, Baoshan, Tengchong and Sibumasu terranes in eastern Asia, which probably remained attached to northeastern Gondwana until the opening of the Neo-Tethys Ocean in the Early Permian (Dan *et al.* 2021). Other terranes located between the Cimmerian microcontinents and the Tarim–Alxa–North China blocks are related to the evolution of the Proto- and Palaeo-Tethys oceans. However, another recently documented Cambrian ocean (the Proto-Qiangtang Ocean), which was located between the Northern Qiangtang–South China Block and Gondwana, complicates reconstructions of the Paleozoic evolution of eastern Asia.

The Proto-Tethys Ocean is commonly viewed as an Ediacaran–early Paleozoic ocean that was initiated by rifting of the Qilian–Qaidam Block at c. 600–580 Ma from NE Gondwana (X. Xu *et al.* 2015). Its vestiges are ophiolite complexes with crystallization ages ranging from 540 to 460 Ma, composed of serpentinite, gabbro, diabase and basalt with compositions similar to normal mid-ocean ridge basalt (N-MORB) or enriched mid-ocean ridge basalt (E-MORB). These ophiolites occur in parallel suture zones within the Qinling–Qilian–Kunlun Orogen (sutures 1–6 in Fig. 1) between the Tarim–North China and Tibetan Plateau–South China continental blocks (e.g. Dong and Santosh 2016; Dong *et al.*



Fig. 1. Tectonic framework of East Asia and adjacent areas, highlighting the locations of Tethys sutures. The numbers represent sutures given in the legend and the terrane names are labelled in the figure.

2018), and are thought to represent vestiges of several branches of the same ocean (Dong *et al.* 2018; Song *et al.* 2018). Arc magmatism in the Kunlun–Qilian segment of the orogen began at *c.* 530–520 Ma and terminated at *c.* 450–440 Ma (C.L. Zhang *et al.* 2019). These events, together with 440–430 Ma regional metamorphism, suggest that ocean closure had occurred by the early Silurian (e.g. Song *et al.* 2013, 2018). Although previously thought to have been closed by north-directed subduction (present orientation), restoration of Mesozoic–Cenozoic deformation suggests that the ocean closed by southward subduction beneath the NE Gondwanan margin (Li *et al.* 2018). The latter interpretation is supported by the documentation of *c.* 530–450 Ma arc magmatism in the Tianshuihai Terrane of the West Kunlun segment, which is located to the south of the Proto-Tethys Suture Zone (suture 5 in Fig. 1) (C.L. Zhang *et al.* 2018).

Multiple Triassic sutures (sutures 7–19 in Fig. 1) are related to the evolution of the Palaeo-Tethys Ocean and relevant oceanic basins. The Longmu Co–Shuanghu, Changning–Menglian and Bengong–Raub sutures are generally considered to be relicts of the main Palaeo-Tethys Ocean (Li 1987; Metcalfe 2013; X.Z. Zhang *et al.* 2017; Dan *et al.* 2018a). This ocean probably opened between the late Silurian and the Middle Devonian (Metcalfe 2013; X.Z. Zhang *et al.* 2014, 2016), although some researchers argue for an older opening based on a few Cambrian–Ordovician ophiolites discovered in the main sutures (B.D. Wang *et al.* 2013; Zhai *et al.* 2016). The tectonic evolution of the main Palaeo-Tethys Ocean is recorded by Early Carboniferous and Early Permian ophiolites (Jian *et al.* 2009; Zhai *et al.* 2013; X.Z. Zhang *et al.* 2016), Late Devonian–Middle Triassic arc igneous rocks (Dan *et al.* 2018b, 2019), and Late Triassic blueschists and eclogites (K.J. Zhang *et al.* 2006; Dan *et al.* 2018a; H.N. Wang *et al.* 2020b). The penecontemporaneous subsidiary sutures north of, and associated with, the main Palaeo-Tethys sutures are generally thought to be relicts of back-arc basins (Y.J. Wang *et al.* 2018), that closed in the Late Triassic (Y.J. Wang *et al.* 2018; Tang *et al.* 2021), coeval with the closure of the main Palaeo-Tethys Ocean.

Abundant early Paleozoic igneous rocks are exposed south of the main Palaeo-Tethys sutures. They are composed predominantly of S-type granites and minor mafic rocks, and were mostly intruded between 510 and 460 Ma (Gao *et al.* 2019; Dan *et al.* 2020). Varied geodynamic settings, including subduction, lithospheric delamination and back-arc extension (Cawood *et al.* 2007, 2021; Y.M. Liu *et al.* 2019; W. Xu *et al.* 2020), have been proposed to explain these igneous rocks. Their composition, as well as their temporal and spatial characteristics, are

typical of a silicic large igneous province (known as Pinghe SLIP) generated above a mantle plume (Dan *et al.* in press). Emplacement of plumes along the periphery of Gondwana may have been common in the late Neoproterozoic–early Paleozoic (Murphy *et al.* 2021), consistent with geodynamic models that predict plumes preferentially nucleating along the edges of large low-shear-wave-velocity provinces (LLSVPs) (e.g. Tan *et al.* 2002; Steinberger and Torsvik 2012), which may have underlain Gondwana at that time (Doucet *et al.* 2020).

Tectonomagmatic events relevant to the Proto-Qiangtang Ocean

Over the past two decades, application of precise U–Pb (zircon) dating techniques has constrained the crystallization ages of early Paleozoic tectonomagmatic events from central Qiangtang, across Sanjiang, Indochina to Hainan Island. As zircon is rarely crystallized from mafic magmas and is often recrystallized during high-grade metamorphic events, we employed several criteria to evaluate whether or not the zircon dates represent the crystallization age of their host rock. First, cathodoluminescence (CL) images of zircon grains crystallized from felsic magmas typically exhibit oscillatory zoning but only weak internal oscillatory or broad-banded zoning if crystallized from mafic magmas (Corfu *et al.* 2003; Dan *et al.* 2021). Second, if the zircon crystallized from the same magma as its host rock, its Hf isotopic values should be compatible with whole-rock Nd isotopic values. Third, zircon grains that have experienced the same high-grade metamorphism as their host are typically recrystallized, producing oval shapes that are highly luminescent in CL images and have Th/U ratios of >0.1 (Chen *et al.* 2010; Dan *et al.* 2012b). After checking the original data, some dating results have been reinterpreted as the ages of xenocrystic zircon grains in mafic rocks or metamorphic ages in metamorphic rocks (see Supplementary Table S1).

In the central Qiangtang region, a *c.* 500 km-long belt of early Paleozoic ophiolites has been discovered in the Taoxinghu, Guoganjialianshan and Shuanghu areas (Figs 2a & 3a), spatially associated with the Longmu Co–Shuanghu Palaeo-Tethys Suture Zone. This belt is typically exposed in isolated blocks or mélanges and is in faulted contact with late Paleozoic sedimentary rocks. The ophiolites are composed of serpentinites, gabbros, basalts and plagiogranites (Fig. 4a), and metamorphosed mafic rocks, some of which are now amphibolites. No sheeted dykes have been found. Whole-rock geochemical data show that the mafic rocks are akin to MORBs but are depleted in high-field strength elements (HFSEs), including Nb and Ti (Fig. 4b, c)

(Zhai *et al.* 2016). Zircon U–Pb dating of gabbros/amphibolites and plagiogranites constrain the formation of oceanic crust to between *c.* 500 and 440 Ma (Hu *et al.* 2014; Zhai *et al.* 2016). Dating of amphibolites and high-pressure granulites indicate that the oceanic lithosphere underwent *c.* 430–420 Ma peak

metamorphism (X.Z. Zhang *et al.* 2014), an event accompanied by the intrusion of a few *c.* 430 Ma diorites and granites (H. Liu *et al.* 2021). These high-pressure granulites exhibit clockwise pressure–temperature–time (*P–T–t*) paths, indicating that the early Paleozoic ocean between the Southern and

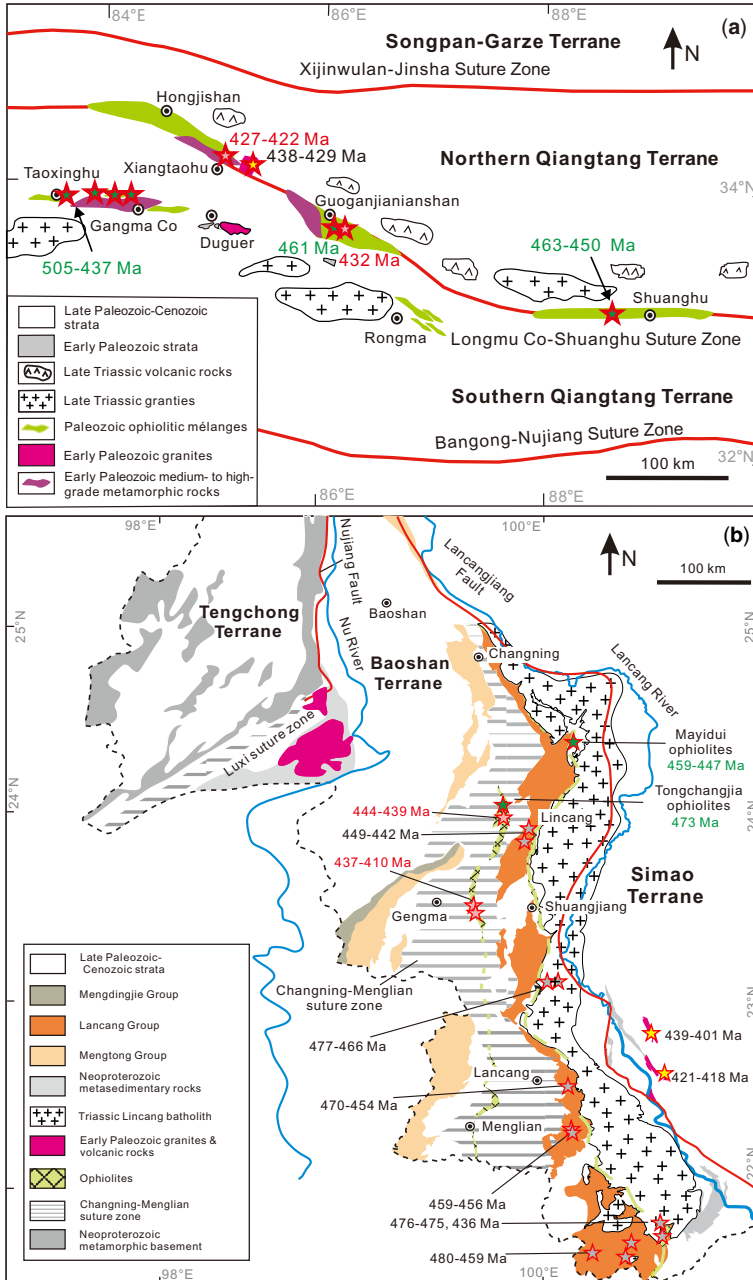


Fig. 2. Simplified regional geological maps for the study areas: (a) Qiangtang, (b) Sanjiang (only shown in China).

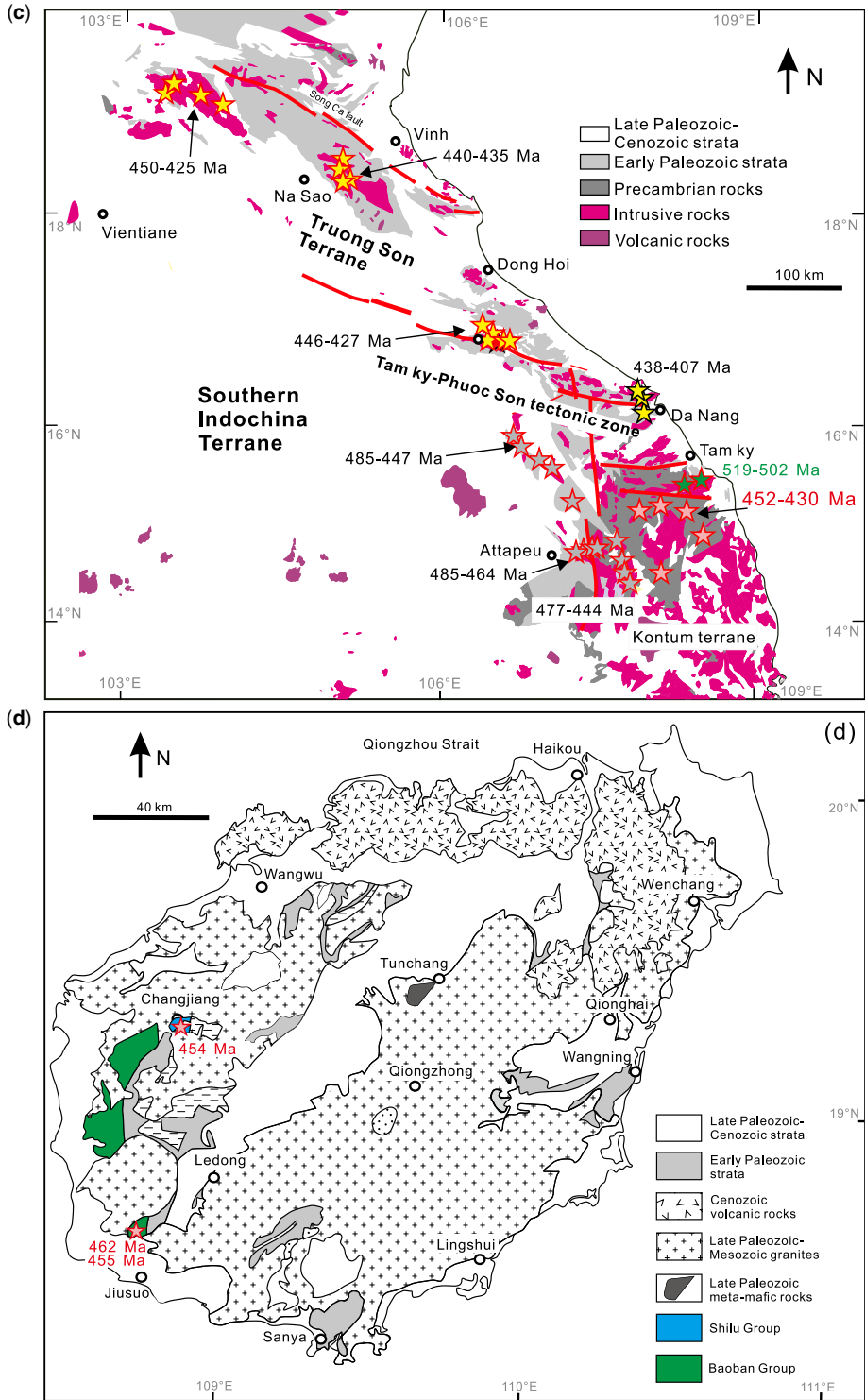


Fig. 2. Continued (c) Indochina and (d) Hainan Island. Age data are based on Supplementary Table S1.

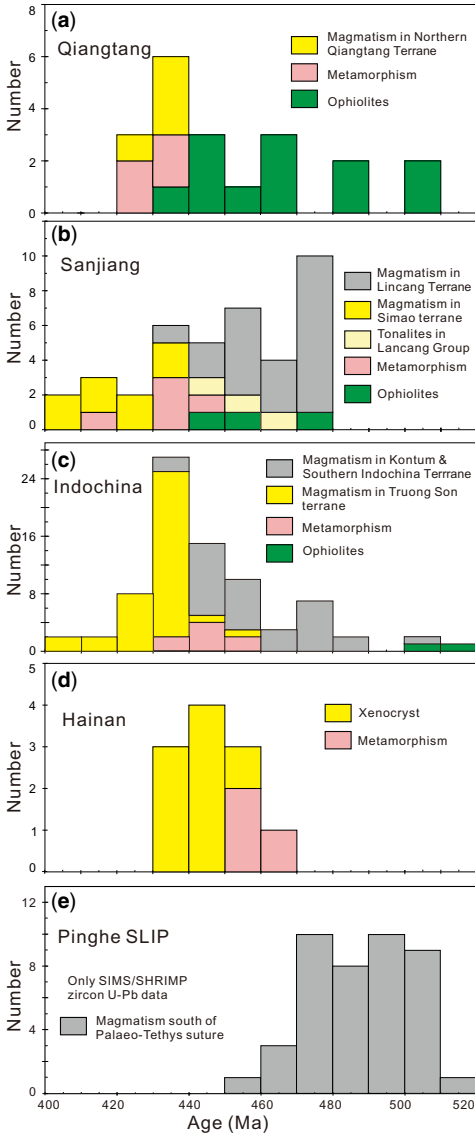


Fig. 3. A compilation of tectonomagmatic ages relevant to the Proto-Qiangtang Ocean. (a)–(d) Zircon U–Pb age or monazite U–Pb or U–Th–Pb age for individual segments. (e) Magmatic zircon U–Pb age for the Pinghe SLIP shown for comparison. Note: (1) only zircon or monazite dates dated by *in situ* methods are plotted in (a)–(d), and the monazite dates interpreted as metamorphic ages were obtained from Indochina and Hainan Island; (2) xenocrystic zircon U–Pb dates are not plotted in (a)–(c) but are shown in (d) as sparse early Paleozoic data obtained in Hainan Island; (3) only precise SIMS/sensitive high-resolution ion microprobe (SHRIMP) zircon U–Pb ages are used for (e); (4) the bin width in each panel is 10 myr; and (5) the data sources for (a)–(d) are listed in [Supplementary Table S1](#), and the data source for (e) is [Dan *et al.* \(in press\)](#).

Northern Qiangtang terranes was closed by the middle Silurian (*c.* 430 Ma) ([X.Z. Zhang *et al.* 2014](#)).

In the Sanjiang region, which consists of the Tengchong, Baoshan and Simao terranes with intervening sutures/faults ([Fig. 2b](#)), U–Pb zircon dating yielded ages of *c.* 470–450 Ma for the Tongchangjie, Mayidui and Wanhe ophiolites along the Changning–Menglian Palaeo-Tethys Suture Zone ([Figs 2b & 3b](#)) ([B.D. Wang *et al.* 2013](#); [G.C. Liu *et al.* 2021b](#)). However, as the early Paleozoic zircon grains from the Wanhe mafic rocks show oscillatory zoning in CL images, and disequilibrium between zircon Hf and whole-rock Nd isotopes, these grains are reinterpreted as xenocrysts. The Tongchangjie and Mayidui ophiolites are in faulted contact with late Paleozoic sedimentary rocks, and are composed of serpentinites, gabbros and metamorphic rocks ([Fig. 4a](#)), some of which are now amphibolites and greenschists. The mafic rocks (gabbros and greenschists) have a MORB affinity but with a depletion in Nb and Ti ([Fig. 4b, c](#)) ([G.C. Liu *et al.* 2021b](#)), traits similar to the early Paleozoic ophiolites in central Qiangtang. Zircon U–Pb ages reveal that the oceanic lithosphere was formed between 470 and 450 Ma ([B.D. Wang *et al.* 2013](#); [G.C. Liu *et al.* 2021b](#)), and metamorphosed at *c.* 440 Ma ([B.D. Wang *et al.* 2013](#)), similar in age to the *c.* 430–410 Ma metamorphism in the Qingping area ([Peng *et al.* 2020, 2021](#)). The *c.* 440–410 Ma age of metamorphism constrained by zircon U–Pb dating is broadly consistent with a *c.* 410 Ma age of metamorphism yielded by the crossite/chlorite Ar–Ar method for blueschists in the Lincang Group ([Cong *et al.* 1993](#)), which is a Cambrian–Ordovician sequence of metasedimentary rocks that was deposited in a continental margin environment ([H.N. Wang *et al.* 2020a](#); [Wei *et al.* 2022](#)). A few *c.* 470–450 Ma tonalites with adakitic geochemical affinities occur as veins intruding the amphibolites and as blocks in the Lincang Group along the suture zone ([D.B. Wang *et al.* 2016](#); [Wu *et al.* 2020](#)). Although a few *c.* 480–440 Ma felsic metavolcanics/granites have been documented in the Lincang Group and Lincang batholith (both comprising the Lincang Terrane) ([Nie *et al.* 2015](#); [Wei *et al.* 2022](#)), only younger 440–400 Ma granites/rhyolites and mafic rocks (amphibolites and basalts) have been reported in the Simao Terrane ([Fig. 2b](#)) ([Jian *et al.* 2009](#); [H.C. Liu *et al.* 2019](#)).

In the Indochina Block ([Fig. 2c](#)), which is further subdivided into several subterrane (Truong Son, Kontum and Southern Indochina terranes) by early Paleozoic sutures or faults, ophiolites occur as blocks within the Tam Ky–Phuoc Son Suture Zone ([Izokh *et al.* 2006](#); [Pham *et al.* 2006](#)). These ophiolites are composed of serpentinites, metagabbros and metabasalts ([Fig. 4a](#)). Zircon U–Pb dating of plagiogranites within the suture zone suggest that

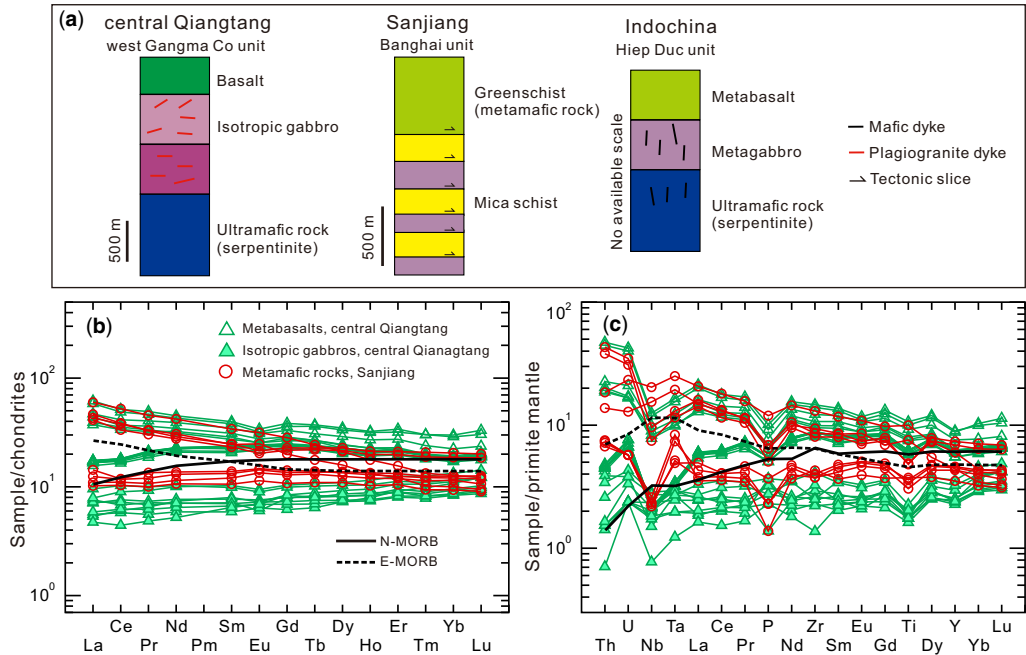


Fig. 4. Schematic columns showing the lithological make-up of (a) the ophiolitic units and (b) and (c) normalized non-cumulate mafic rocks of the ophiolites. Schematic columns of ophiolites are based on previous studies – central Qiangtang: Wu *et al.* (2009) and Zhai *et al.* (2016); Sanjiang: G.C. Liu *et al.* (2021b); Indochina: Izokh *et al.* (2006), Pham *et al.* (2006) and Nguyen *et al.* (2019). Geochemical data are from central Qiangtang (Zhai *et al.* 2016) and Sanjiang: G.C. Liu *et al.* (2021b). The normalized N-MORB and E-MORB values are from Sun and McDonough (1989).

oceanic crust generation occurred at *c.* 510 Ma (Figs 2c & 3c) (Nguyen *et al.* 2019). In contrast to the *c.* 500–440 Ma granitic and minor mafic magmatism documented in both the Kontum and Southern Indochina terranes, only younger (450–400 Ma) granites/rhyolites and mafic rocks were reported in the Truong Son Terrane (Jiang *et al.* 2020; T.V. Tran *et al.* 2020; Y.J. Wang *et al.* 2020, 2021a, b; Nakano *et al.* 2021). Zircon and monazite dating of paragneisses and amphibolites provide a *c.* 450–430 Ma age for amphibolite-facies metamorphism in the Tam Ky-Phuoc Son Suture Zone and the Kontum Terrane, which is interpreted to reflect a collision between the Kontum and Truong Son terranes (Roger *et al.* 2007; Usuki *et al.* 2009; Nakano *et al.* 2013; H.T. Tran *et al.* 2014).

In Hainan Island, no early Paleozoic ophiolites have been reported. A reliable determination of a *c.* 460–450 Ma age for regional amphibolite-facies metamorphism was obtained from the Precambrian Baoban Complex and Shilu Group (Figs 2d & 3d) (D.R. Xu *et al.* 2015; Y.J. Xu *et al.* 2020; H. Zhang *et al.* 2022). The recently identified metabasalts/andesites (Zhou *et al.* 2021) are likely to have formed in the Early Carboniferous (He *et al.*

2018), as samples from both studies collected from same areas yield Early Carboniferous zircon U–Pb ages (He *et al.* 2018), and the zircon grains in the former study show obvious oscillatory zoning in CL images.

In summary, early Paleozoic ophiolites preserved in the Longmuco–Shuanghu Palaeo-Tethys, Changning–Menglian Palaeo-Tethys and Tam Ky-Phuoc Son suture zones were formed and emplaced between *c.* 510 and 440 Ma. Basalts show MORB-like characteristics but with HFSE depletion (Fig. 4), implying their generation in a suprasubduction zone setting. The regional metamorphic ages in the Proto-Qiangtang Ocean realm are *c.* 460–450 Ma for Hainan Island, *c.* 450–430 Ma for Indochina, *c.* 440–410 Ma for Sanjiang and *c.* 430–420 Ma for Qiangtang, suggesting diachronous collision from the SE to the NW.

Discussion

The discovery of early Paleozoic ophiolites in Qiangtang, Sanjiang and Indochina indicates that there was an early Paleozoic ocean located south of

the suture zone inferred to represent the boundary between NE Gondwana and the Proto-Tethys Ocean. This ocean has previously been given local names and has been variously interpreted as a branch of the Palaeo-Tethys or the Proto-Tethys. As the ophiolites are best preserved and were first studied in central Qiangtang, we herein propose that this early Paleozoic ocean be named the Proto-Qiangtang Ocean.

Tectonic evolution of the Proto-Qiangtang Ocean and its relationship with the Palaeo-Tethys Ocean

Based on data from different segments of the Proto-Qiangtang Ocean, most researchers consider the Proto-Qiangtang Ocean to have been generated in the early Paleozoic (B.D. Wang *et al.* 2013; Zhai *et al.* 2016), although some propose a late Neoproterozoic origin for its Indochina segment (T.V. Tran *et al.* 2020). Geochemical comparisons with the Troodos and northern Pamir plagiogranites, including a similar depletion of HFSEs (Nb and Ti), suggest that the *c.* 510 Ma plagiogranites in Indochina (Tam Ky-Phuoc Son Suture Zone) were generated in an intra-oceanic arc setting (Nguyen *et al.* 2019) that has recently been reinterpreted as a back-arc setting (Tang *et al.* 2021). No ophiolites older than these plagiogranites have been documented in the Proto-Qiangtang Ocean realm, implying that this ocean was probably formed no earlier than the middle Cambrian (*c.* 510 Ma). The Proto-Qiangtang Ocean was located south of the subduction zone documented within the Proto-Tethys, implying it was generated in a back-arc tectonic setting (Zhai *et al.* 2016; W. Xu *et al.* 2020), a scenario consistent with the suprasubduction-zone-type geochemistry of the ophiolites reported in the Proto-Qiangtang Suture Zone. A mantle plume, responsible for the generation of the *c.* 510–460 Ma Pinghe SLIP (Dan *et al.* *in press*), could also have played an important role in the generation of this ocean. This SLIP could be a local example of mantle plumes that were preferentially emplaced along the margins of Gondwana in the late Neoproterozoic–early Paleozoic (Murphy *et al.* 2021). Emplacement of plumes preferentially exploits weaknesses in the lithosphere (Obrebski *et al.* 2010; Coble and Mahood 2012), such as those provided by the opening of back-arc basins.

The subduction polarity of the Proto-Qiangtang Ocean is controversial, and models of northward (B.D. Wang *et al.* 2013; Faure *et al.* 2018; Jiang *et al.* 2020), southward (Y.J. Wang *et al.* 2020, 2021*b*) or even bivergent subduction have been proposed (G.C. Liu *et al.* 2021*b*; H. Liu *et al.* 2021). The model for a southward subduction polarity is mostly

based on the *c.* 510–460 Ma magmatism located south of the Proto-Qiangtang Ocean. However, if these igneous rocks that constitute the Pinghe SLIP were generated by a mantle plume (Dan *et al.* *in press*; Murphy *et al.* 2021), there is no independent evidence that would support this model. Direct documentation of arc magmatism that would record subduction polarity remains elusive. However, as shown by studies focused on the subduction record of the Rheic Ocean (Pereira *et al.* 2017), the former existence of magmatic arcs could be preserved in detrital zircon age patterns in coeval or younger clastic deposits eroded from that arc (Cawood *et al.* 2012; Dan *et al.* 2012*a*; X.Z. Zhang *et al.* 2017). Our compilation of detrital zircon data reveals contrasting detrital zircon age spectra for different terranes on opposite sides of the Proto-Qiangtang Suture (Fig. 5). Silurian and younger sedimentary rocks north of the Proto-Qiangtang Suture, such as in the Northern Qiangtang, Simao, Truong Son and Hainan Island terranes (Fig. 5*b, d, h, i*), have detrital zircon age spectra characterized by a major 440 Ma peak. The coeval relationship between the 440 Ma detrital zircon peak and the Silurian strata that contain these zircons is typical of arc settings (Cawood *et al.* 2012; Dan *et al.* 2012*a*). In contrast, in Silurian or younger sedimentary rocks south of the Proto-Qiangtang Suture, such as those in the Southern Qiangtang, Baoshan–Tengchong and Kontum terranes (Fig. 5*a, c, f*), the dominant detrital zircon age populations are older than *c.* 500 Ma. Collectively, these data suggest that the Proto-Qiangtang Ocean underwent northward subduction, and the Ordovician–Silurian ages of detrital zircons deposited in strata to the north of this ocean were derived from coeval arc magmatism related to such subduction.

The detrital zircon age populations can be used further to clarify the tectonic affinity of some disputed terranes along the sutures. In the Sanjiang region (Fig. 2*b*), the Lincang Terrane, which is located between the Changning–Menglian Palaeo-Tethys Suture and the Simao Terrane, and consists of the Lincang Group and the Lincang batholiths, is considered to be tectonically linked with Gondwana (Baoshan Terrane) but not with the adjacent Simao Terrane (Feng *et al.* 1996). This interpretation is consistent with the detrital zircon age spectrum, which has no significant age peak younger than *c.* 500 Ma (Fig. 5*e*), indicating that these rocks were located south of the Proto-Qiangtang Ocean in a passive-margin setting. Thus, the *c.* 500–460 Ma igneous rocks in this terrane are not directly related to arc magmatism within the Proto-Qiangtang Ocean but are more likely to be part of the *c.* 510–460 Ma Pinghe SLIP (Dan *et al.* *in press*). A minor Silurian (*c.* 430 Ma) detrital zircon age peak occurs in the uppermost strata of the Lincang Group

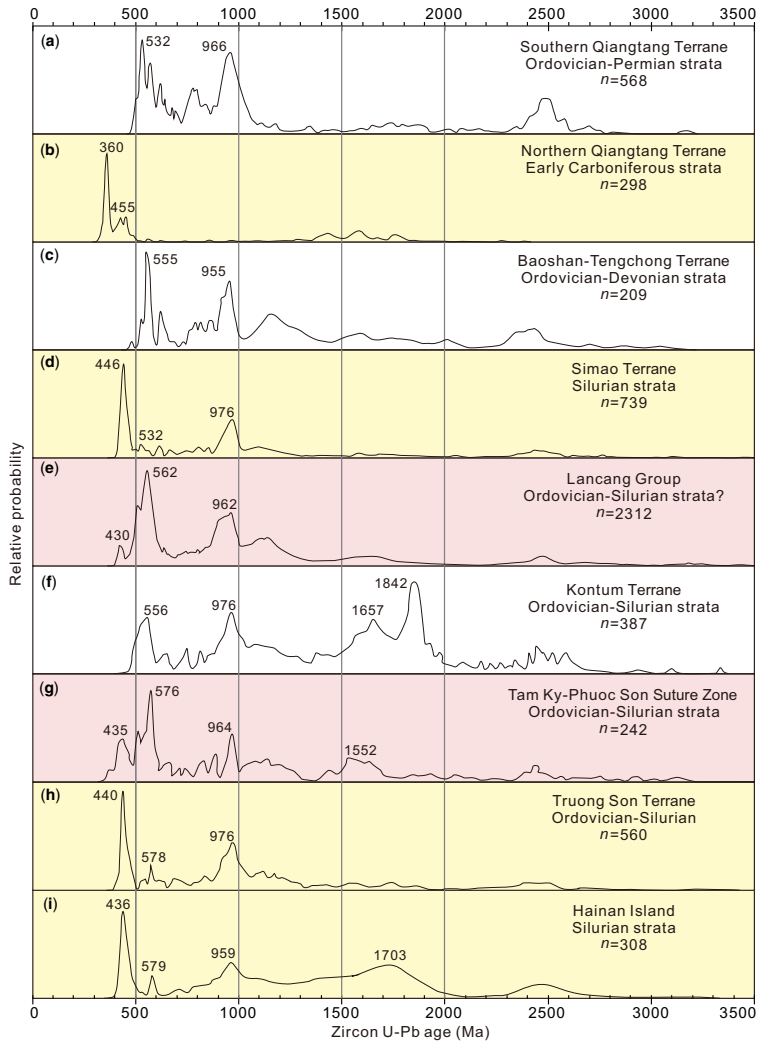


Fig. 5. Relative probability plots comparing detrital zircon age spectra. Note: (1) n in each panel is the number of detrital zircon ages; and (2) there are no detrital zircon age data available for Silurian–Devonian strata in Northern Qiangtang, thus we show data from Early Carboniferous strata in (b). Data sources: the Southern Qiangtang and Northern Qiangtang terranes, X.Z. Zhang *et al.* (2017) and references therein; the Baoshan–Tengchong and Simao terranes and the Lancang Group, H.N. Wang *et al.* (2020a) and references therein and Wei *et al.* (2022); the Kontum Terrane, Tam Ky-Phuoc Son Suture Zone and the Truong Son Terrane, Y.J. Wang *et al.* (2021a) and references therein; and Hainan Island, Zhou *et al.* (2015).

(Wei *et al.* 2022), which were likely to have been sourced from the magmatic arc after closure of the Proto-Qiangtang Ocean. Detrital zircon age spectra in the Tam Ky-Phuoc Son Suture Zone are similar to those in the Lancang Group (Fig. 5e, g) and indicate the suture zone was located south of the Proto-Qiangtang Ocean.

The timing of closure of the Proto-Qiangtang Ocean is controversial, with interpretations ranging from the Late Triassic to the Silurian. The Paleozoic

ophiolites were initially interpreted to be relicts of the Palaeo-Tethys Ocean, in which the early Paleozoic ocean had evolved into the Palaeo-Tethys Ocean by the late Paleozoic (B.D. Wang *et al.* 2013; Zhai *et al.* 2016) and then closed in the Late Triassic (K.J. Zhang *et al.* 2006; Dan *et al.* 2018a; W. Xu *et al.* 2020; H.N. Wang *et al.* 2020b). However, more recent documentation of Middle Ordovician–Silurian metamorphic and igneous rocks along the whole belt suggests that the Proto-Qiangtang

Ocean closed during the early Paleozoic (H.T. Tran *et al.* 2014; X.Z. Zhang *et al.* 2014; Faure *et al.* 2018; Jiang *et al.* 2020; T.V. Tran *et al.* 2020; W. Xu *et al.* 2020; Y.J. Wang *et al.* 2020; H. Zhang *et al.* 2022). The *c.* 470 Ma adakitic rocks documented in the suture zones (D.B. Wang *et al.* 2016; Gardner *et al.* 2017) are indicative of the onset of arc magmatism. The *c.* 430 Ma high-pressure granulites accompanied by amphibolite-facies metamorphism in the Qiangtang region indicate this segment of the Proto-Qiangtang Ocean was closed by that time (X.Z. Zhang *et al.* 2014). No high-pressure metamorphic rocks are reported in other segments but early Paleozoic amphibolite-facies metamorphism, combined with (a) coeval magmatism with post-collisional geochemical affinities and (b) post-Silurian termination of arc magmatism based on the detrital zircon age populations (Fig. 5), suggest that the Proto-Qiangtang Ocean was closed by *c.* 430 Ma in these regions. However, younger metamorphic ages from Hainan Island to central Qiangtang indicate diachronous closure of the Proto-Qiangtang Ocean between *c.* 460 and 430 Ma. As the age of collision is similar to the detrital zircon peak in each segment (Fig. 5), arc magmatism was likely to have been of short duration.

Thus, the tectonic evolution of the Proto-Qiangtang Ocean can be summarized in the following stages (Fig. 6). The Proto-Tethys Ocean subduction zone geodynamics generated a back-arc region with inherent structural weaknesses that

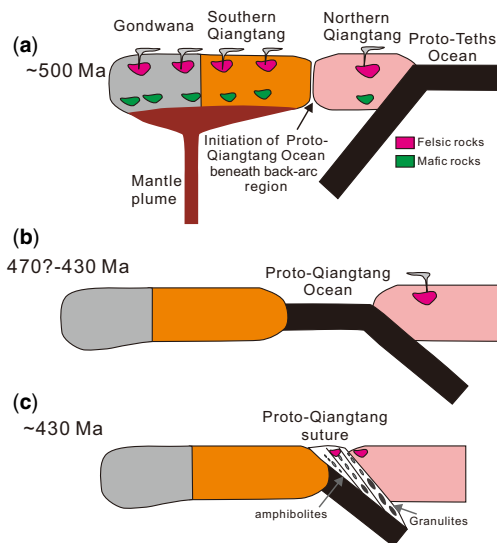


Fig. 6. Schematic model for the tectonic evolution of the Proto-Qiangtang Ocean in the Qiangtang segment.

were exploited by one of several mantle plumes that preferentially ascended along the periphery of Gondwana. The feedback between back-arc extension and mantle-plume emplacement generated a narrow late Cambrian ocean between the Northern Qiangtang–South China–Northern Indochina terranes and the Southern Qiangtang–Baoshan–Southern Indochina terranes. This ocean underwent northward subduction beneath the Northern Qiangtang–South China–Northern Indochina terranes, probably beginning at *c.* 470 Ma, and was closed diachronously between *c.* 460 and 430 Ma from Hainan Island to central Qiangtang. Magmatism after *c.* 450–440 Ma in central Indochina, *c.* 440 Ma in Sanjiang and *c.* 430 Ma in central Qiangtang formed in a syncollisional setting, as supported by the documentation of S-type granites that formed at *c.* 450–440 Ma in the Truong Son Terrane, Indochina (Jiang *et al.* 2020) and at *c.* 436 Ma in Sanjiang (G.C. Liu *et al.* 2021a). Younger episodes (*c.* 420–400 Ma) of magmatism in Indochina and Sanjiang were generated in a post-collisional (Jiang *et al.* 2020) or a rift setting (Jian *et al.* 2009; H.C. Liu *et al.* 2019), possibly related to the opening the Palaeo-Tethys Ocean. The Palaeo-Tethys Suture Zone spatially overlaps with the Proto-Qiangtang Suture, especially in the Qiangtang and Sanjiang segments (Figs 1 & 2), indicating that the rifting preferentially occurred near or along a pre-existing suture (Murphy *et al.* 2006).

Correlation of the Rheic Ocean with the Proto-Qiangtang Ocean

The potential geodynamic relationships between the Paleozoic oceans developed along the northern margin of Gondwana are widely debated. The Proto-Tethys Ocean is generally considered to be the eastern extension of the Iapetus Ocean (e.g. Stampfli and Borel 2002; Murphy *et al.* 2010; Stampfli *et al.* 2013). However, the identity of the eastern extension of the Rheic Ocean is controversial, and proposed counterparts include the Proto-Tethys and Palaeo-Tethys oceans (Stampfli and Borel 2002; Metcalfe 2013; Stampfli *et al.* 2013; Torsvik and Cocks 2017; Li *et al.* 2018). Here, we propose that its eastern extension is probably the Proto-Qiangtang Ocean (Fig. 7).

The Rheic Ocean formed by the separation of Carolina–Avalonia–Ganderia from Gondwana during the late Cambrian–Early Ordovician (Nance *et al.* 2010; van Staal *et al.* 2012, 2021), as recorded in Iberia by ophiolites (Sánchez Martínez *et al.* 2012, 2021) and rift–drift successions (Sánchez-García *et al.* 2003, 2008). The northward drift of Avalonia is commonly thought to be broadly related to the southward subduction and rollback of Iapetus oceanic lithosphere, although, in detail, the mechanisms

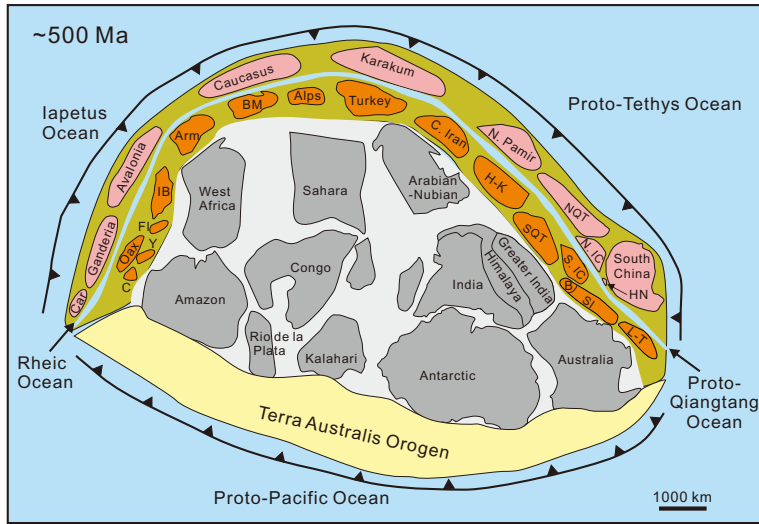


Fig. 7. Palaeogeographical reconstruction of Gondwana in the late Cambrian (*c.* 500 Ma). Note: (1) the shape of the peri-Gondwanan terranes is conjectural, and the scale is for Gondwana; and (2) abbreviations: Arm, Armorica; B, Baoshan; BM, Bohemian Massif; C, Chortis; Car, Carolina; C. Iran, central Iran; Fl, Florida; H-K, Helmand–Karakoram; HN, Hainan Island; IB, Iberia; L-T, Lhasa-Tengchong; N. IC, Northern Indochina; NQT, Northern Qiangtang; Oax, Oaxaquia; S. IC, Southern Indochina; SI, Sibumasu; SQT, Southern Qiangtang; Y, Yucatan.

proposed are highly variable (Nance *et al.* 2010). Coeval with the Rheic Ocean opening, the *c.* 495–470 Ma Ollo de Sapo SLIP was emplaced along the Gondwanan margin of this ocean (García-Arias *et al.* 2018), which is potentially another example of the importance of plume emplacement to the evolution of the Gondwanan margin (Murphy *et al.* 2021). In the Silurian–Devonian, the Rheic Ocean underwent northward subduction beneath Avalonia, with evidence for the associated magmatic arc recorded as detrital zircon populations in Late Devonian–Early Carboniferous siliciclastic rocks (Murphy *et al.* 2010; Nance *et al.* 2010; Pereira *et al.* 2017). Devonian (*c.* 400–380 Ma) ophiolites (Murphy *et al.* 2009; Arenas *et al.* 2021) and two stages of high-pressure metamorphism (*c.* 400 and *c.* 370–330 Ma) (Arenas *et al.* 2014) reported in the Variscan Orogen imply a complex or protracted collision during amalgamation of Laurussia and Gondwana (Kroner and Romer 2013; Franke *et al.* 2017; Wu *et al.* 2021).

The tectonic evolution of the Rheic Ocean described above is similar to that of the Proto-Qiangtang Ocean. First, both oceans opened in the late Cambrian–Early Ordovician along the northern margin of Gondwana by dispersal of peri-Gondwanan terranes. Second, creation of these oceans was related to the southward subduction and rollback of oceanic lithosphere (Iapetus and Proto-Tethys oceans) beneath Gondwana, and both were accompanied by emplacement of a SLIP (the *c.* 495–

470 Ma Ollo de Sapo SLIP in NW Gondwana and the *c.* 510–460 Ma Pinghe SLIP in NE Gondwana), which were likely to have been generated by mantle plumes that were preferentially emplaced in zones of lithospheric weakness along the Gondwanan margin (Murphy *et al.* 2021). Third, the Rheic and Proto-Qiangtang oceans were both closed by north-directed subduction beneath the previously drifted peri-Gondwanan terranes. Fourth, the diachronous Devonian–Carboniferous closure of the Rheic Ocean (Wu *et al.* 2021) is consistent with the westward diachronous closure of the Proto-Qiangtang Ocean, and implies highly diachronous closure over 80 myr between East Asia and West Africa. Thus, we propose that the Proto-Qiangtang Ocean probably represents the eastern extension of the Rheic Ocean. Although more studies are needed to test our hypothesis and to refine the tectonic evolutions of both the Proto-Qiangtang and Rheic oceans, our model provides new constraints on the Paleozoic tectonic evolution and palaeogeography of the northern Gondwanan margin.

Conclusions

This study provides new constraints on the early Paleozoic tectonic evolution and palaeogeography along the northern Gondwanan margin. The Proto-Qiangtang Ocean, located south of the Proto-Tethys Ocean, is a short-lived (*c.* 80 myr), early Paleozoic

ocean. This ocean opened in the late Cambrian (*c.* 510–500 Ma) by the northward drift of peri-Gondwanan terranes from central Qiangtang, across Sanjiang and central Indochina, to Hainan Island, and was possibly generated as a back-arc basin whose opening may have been assisted by the emplacement of a mantle plume. The Proto-Qiangtang Ocean underwent northward subduction beneath the dispersed peri-Gondwanan terranes (Northern Qiangtang–South China–NE Indochina) beginning in the Middle Ordovician (*c.* 470 Ma), and closed diachronously westwards, from Hainan Island to central Qiangtang, between *c.* 460 and 430 Ma. A Silurian closure of this ocean distinguishes it from the Silurian–Devonian opening of the Palaeo-Tethys Ocean. The Proto-Qiangtang Ocean shares a similar tectonic evolution to that of the Rheic Ocean in the early Paleozoic, and together they constitute an early Paleozoic ocean outboard of the northern Gondwanan margin.

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Author contributions **WD**: conceptualization (lead), data curation (lead), formal analysis (lead), methodology (lead), writing – original draft (lead), writing – review & editing (lead); **JBM**: conceptualization (equal), data curation (supporting), formal analysis (supporting), methodology (supporting), writing – original draft (equal), writing – review & editing (equal); **QW**: conceptualization (supporting), data curation (supporting), formal analysis (supporting), methodology (supporting), writing – original draft (supporting), writing – review & editing (supporting); **X-ZZ**: conceptualization (supporting), data curation (supporting), formal analysis (supporting), methodology (supporting), writing – original draft (supporting), writing – review & editing (supporting); **G-JT**: conceptualization (supporting), data curation (supporting), formal analysis (supporting), methodology (supporting), writing – original draft (supporting), writing – review & editing (supporting).

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Data availability All data generated or analysed during this study are included in this published article and its supplementary information file.

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