GEOCHEMICAL ANALYSIS



Check for updates

Improved Apparatus for Phosphoric Acid Digestion of Carbonates to Determine the Carbon, Oxygen and Clumped Isotope Compositions

Xi Liu^{a,b,c} , Wenfeng Deng^{a,b,c} , Yangrui Guo^{a,b,c}, and Gangjian Wei^{a,b,c}

^aState Key Laboratory of Isotope Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou, China; ^bCAS Center for Excellence in Deep Earth Science, Guangzhou, China; ^cSouthern Marine Science and Engineering Guangdong Laboratory, Guangzhou, China

ABSTRACT

The carbon, oxygen and clumped isotope compositions of carbonates are typically analyzed by carbonate-phosphoric acid digestion. Although this approach has been widely used since the 1950s, this procedure has problems in the operability. To address these issues, an improved apparatus with properly sized accessories, including an I-shaped vessel, a sample boat, a magnet, and a magnetic stir bar, has been developed for sample digestion. Similar to a long neck flask, this apparatus has a straight and axisymmetric structure, providing advantages in terms of cost and operability, and thus can be easily implemented in both open and sealed vessel digestion. To evaluate the reliability of the stable and clumped isotope compositions using this new vessel, international and interlaboratory carbonate standards were analyzed with this new apparatus and two commonly used Y-shaped and T-shaped vessels. The results using the new apparatus were consistent with those obtained by the others and in previous studies, suggesting suitability for the traditioinal phosphoric acid digestion of carbonates.

ARTICLE HISTORY

Received 1 February 2021 Accepted 13 March 2021

KEYWORDS

carbonates; carbon isotope; clumped isotope; oxygen isotope; phosphoric acid digestion

Introduction

Carbonate-phosphoric acid digestion is the cornerstone of studies on carbon, oxygen and clumped isotopes in carbonates (McCrea 1950; Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006; Rosenbaum and Sheppard 1986), which provide important information on paleoclimate, paleoenvironment, paleoelevation, diagenesis and ore deposits (Gilg et al. 2008; Deng et al. 2019; Ghosh, Garzione, et al. 2006; Li et al. 2019). The traditional digestion procedure is manually carried out on an offline system, which consists of containers for powders reacting with anhydrous phosphoric acid and a glass/ stainless vacuum line for removal of CO_2 (McCrea 1950; Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006). Commercial carbonate devices that use sample vials as reaction containers offer automated procedures for carbonates with low sample requirements, high efficiency, and reduced labor (Groot 2009; Révész and Landwehr 2002;

CONTACT Wenfeng Deng 🐼 wfdeng@gig.ac.cn 🝙 511 Kehua Street, Guangzhou, Guangdong 510640, China 🚯 Supplemental data for this article is available online at https://doi.org/10.1080/00032719.2021.1904252 © 2021 Taylor & Francis Group, LLC

Spötl and Vennemann 2003). Hence, the automated procedures are preferentially employed for the measurements on stable isotopes in carbonates that are soluble in acid, such as calcite, aragonite and dolomite, that completely react with phosphoric acid in less than 30 mins (Dean et al. 2015; Currie et al. 2016; Deng et al. 2017; Liu, Deng, and Wei 2019). In recent years, technical improvements have been implemented on automated systems for clumped isotopes in carbonates ($100-120 \mu g$) (Schmid and Bernasconi 2010; Schmid, Radke, and Bernasconi 2012; Hu et al. 2014; Meckler et al. 2014), allowing measurements of small or high-resolution samples (Leutert et al. 2019; Meinicke et al. 2020). However, manual procedures are used when analyzing isotopic compositions of acid-resistant carbonates (such as magnesite and siderite), CO₂ selectively extracted from impure carbonates (such as coexisting calcite-dolomite pairs), and determining phosphoric acid fractionation factors in carbonate minerals (Swart, Burns, and Leder 1991; Rosenbaum and Sheppard 1986; Sharma, Patil, and Gopalan 2002; Fernandez et al. 2016; Swart et al. 2019; Al-Aasm, Taylor, and South 1990).

In the traditional procedure, a McCrea-type (Y-shaped) vessel brings the carbonate powder in contact with phosphoric acid by tilting its body when thermal equilibrium is achieved (Figure S1(A)) (McCrea 1950). The thermal equilibrium is essential to stabilize the acid fractionation factors of oxygen and clumped isotopes (McCrea 1950; Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006; Kim, Coplen, and Horita 2015). Therefore, the McCrea-type vessel has been widely used since 1950 in the traditional procedure, although other containers have been developed (Figure S1) (McCrea 1950; Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006; Groot 2009; Kim, Coplen, and Horita 2015). Acid digestion handled by these reaction containers includes sealed vessel (SV) and open vessel (OV) (Swart, Burns, and Leder 1991; Kim, Coplen, and Horita 2015; Groot 2009; Fernandez et al. 2016). The CO₂ produced by SV digestion is transferred into the cleaning line after completion of the acid reaction, while CO₂ produced by OV digestion is continuously collected in the cleaning line (Swart, Burns, and Leder 1991; Fernandez et al. 2016). The McCrea-type vessel may be used in either mode, but it is generally used in SV digestion because more accessories and manual operations are needed for OV digestion (Fernandez et al. 2016).

To improve the analysis of carbonate clumped isotopes, some reaction containers in OV digestion are designed for a common or an individual acid bath (CAB or IAB, respectively), in which an aliquot of acid reacts with one or several samples (Passey et al. 2010; Kluge and John 2015; Defliese, Hren, and Lohmann 2015). Consequently, memory effects or cross-contamination may occur in CAB (Swart, Burns, and Leder 1991; Passey et al. 2010). For an IAB carried out through a new reaction container with T-shaped glass tube (Figure S1C), its convenience involves the tipping of the McCreatype vessel replaced with a push operation using a sample boat and magnets on the horizontal glass arm of the container (Kluge and John 2015). However, this T-shaped vessel has a limitation that samples and the boat are not heated until they are in phosphoric acid (Figure S1(C)). Hence, an unstable fractionation of oxygen and clumped isotopes theoretically exists because carbonates initially react with phosphoric acid under thermal disequilibrium (Guo, Deng, and Wei 2019; Kluge and John 2015; Groot 2009). To date, the effect of this thermal disequilibrium on the measured results has not been technically identified when a 90 °C temperature is used (Guo, Deng, and Wei

2019; Kluge and John 2015). The thermal disequilibrium on acid fractionation factors is probably changed or enlarged by the boat size, the ratio of sample to acid, and the content of noncarbonate components in the samples.

The McCrea-type (Y-shaped) and Kluge-type (T-shaped) vessels (Figure S1A, S1C)) are the most common design for IABs (McCrea 1950; Swart, Burns, and Leder 1991; Groot 2009; Swart et al. 2019; Y. Guo, Deng, and Wei 2019; Kluge and John 2015), because carbonates are separated from phosphoric acid using the two glass arms in a straightforward manner. However, there are disadvantages of these vessels. First, the assembly consists of many steps requiring special care; otherwise, unexpected contact between the sample and acid will occur. Second, the vessel may to tip over because of the asymmetric design; thus, the vessel should be handled with caution. The asymmetry also complicates the manufacturing process. In particular, unlike the McCrea-type vessel, the horizontal glass arm of the Kluge-type vessel is sealed using a stainless steel tube fitting (Kluge and John 2015), which increases the difficulty of assembly, the vessel asymmetry, and a possible leakage point.

Two straight-tube reaction containers generally include one large and one small glass tube for carbonate powders and phosphoric acid, respectively (Figure S1B, S1D) (Swart, Burns, and Leder 1991; Krishnamurthy, Atekwana, and Guha 1997; Fosu et al. 2019). The simple and symmetric structure of these vessels facilitates the manufacture, sample load and acid injection. Nevertheless, when using these vessels, tilting is needed to trigger the reaction between carbonates and acid. The vessel in Figure S1B is rarely used because of weak seal by the septum and potential contamination from glue to attach the acid boat (Krishnamurthy, Atekwana, and Guha 1997). The sample boat method (Figure S1D) has been used (Swart, Burns, and Leder 1991; Fosu et al. 2019; Murray, Arienzo, and Swart 2016) with valve (Swart, Burns, and Leder 1991; Murray, Arienzo, and Swart 2016) and flame (Fosu et al. 2019) methods to seal the vessel. Flame sealing makes the vessel leak-free and suitable for carbonate digestions that take longer times (Fosu et al. 2019).

In this study, an improved apparatus was designed for acid digestion that includes a simple I-shaped vessel with a magnet, a sample boat and a magnetic stir bar. The assembly, movement and sample introduction of this I-shaped vessel are more convenient than previous designs and hence allows the routine phosphoric acid digestion of carbonates to determine stable and clumped isotope compositions.

Experimental

Samples

Six carbonate samples were analyzed for intervessel and interlaboratory comparisons. Three are reference materials distributed by the International Atomic Energy Agency (IAEA): NBS18, NBS19 and IAEA-C1 (Rozanski et al. 1992; Friedman, O'Neil, and Cebula 1982). NBS18 and NBS19 are used as carbonate stable-isotope standards, and their clumped isotope compositions have also been reported (Li et al. 2019; Fosu et al. 2019; Chang et al. 2020; Wacker, Fiebig, and Schoene 2013). IAEA-C1 is a reference material for ¹⁴C activity and made from a slab of Carrara marble (Rozanski et al. 1992). When analyzing carbonate clumped isotope compositions, IAEA-C1 was used as an in-



Figure 1. Schematic of the assembled I-shaped vessel.

house standard to monitor instrumental drift and make water-equilibrated or heated gases (Guo, Deng, and Wei 2019; Guo, Deng, and Wei 2019; Guo et al. 2020). BACS and ISCS are also in-house standards for carbonate clumped isotopes prepared from tropical *Porites* coral aragonite and Iceland spar calcite, respectively (Y. Guo, Deng, and Wei 2019). The interlaboratory standard, ETH-4, for carbonate clumped isotopes is a synthetic distributed by the Geological Institute of ETH Zürich (Bernasconi et al. 2018).

Vessels for the phosphoric acid digestion of carbonates

To evaluate the reliability of measurements using the I-shaped vessel (Figure 1), acid digestion was also performed on the Y-shaped and T-shaped vessels similar to Figure S1A and Figure S1C. Considering that there are possible fractionation differences in oxygen and clumped isotopes between OV and SV digestions at the popular reaction temperature of 90 °C (Swart, Burns, and Leder 1991; Fernandez et al. 2016; Swart et al. 2019), both modes of the I-shaped vessel were used to assess the differences. In all

digestions, each sample was weighed to 4-6 mg in a quartz glass sample boat (Φ 6 × 10 mm), and ~104% phosphoric acid is used.

The Y-shaped vessel is used to conduct SV digestion. Different from previous studies (McCrea 1950; Swart, Burns, and Leder 1991; Fernandez et al. 2016; Ghosh, Garzione, et al. 2006; Rosenbaum and Sheppard 1986), carbonate powders were initially loaded into a sample boat and placed at the bottom of the vessel to avoid powders attaching to the vessel. Approximately 2 mL phosphoric acid was introduced into the slant arm using a plastic pipette. After assembly with a vacuum fitting and a ball valve, vessels are connected to the vacuum line and evacuated for approximately 12 h. Subsequent procedures are similar to those employed using McCrea-type vessels (McCrea 1950; Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006). Thermal equilibrium is achieved in a water bath at 90 °C for >0.5 h. The reaction is activated by tilting the vessel and lasts for >0.5 h. When the reaction is completed, the vessel is connected to the CO₂ cleaning line. After the cleaning line under vacuum, the CO₂ is collected by a liquid nitrogen trap. This procedure is used for the IAEA-C1 samples and labeled C1-Y-M.

The T-shaped vessel is similar to the container of Kluge and John (2015) (Figure S1C) and has also been used for OV digestion (Guo, Deng, and Wei 2019; Guo 2019). Approximately 2 mL phosphoric acid and a PTFE magnetic stir bar (Φ 4 × 8 mm) were placed into the vertical arm of the T-shaped vessel. Caution should be exercised when the sample boat loaded with 4-6 mg sample and two magnets are assembled with the horizontal arm to prevent the powder from escaping or the boat from dropping into the vertical arm. The two arms of the T-shaped vessel were carefully sealed using a combination of vacuum fitting and glass tube so that moisture cannot be absorbed by phosphoric acid. After the connection of the assembled vessel to the vacuum line, the vertical arm is heated in a 90 °C water bath. The vessel is evacuated for >0.5 h to achieve a vacuum on the order of 10^{-3} to 10^{-2} mbar. The degassing of phosphoric acid is promoted by stirring. The reaction is activated when the sample boat is pushed into the vertical arm by the magnets. The CO₂ is continuously collected in a liquid nitrogen trap. This process was used for the BACS samples labeled as BACS-T.

The I-shaped vessel designed in this study is a long neck flask (Figure 1) with four assembly steps. First, 2 mL phosphoric acid are transferred to the vessel using a plastic pipette. Phosphoric acid is viscous, and the slender part of the vessel should be free from acid to ensure the smooth movement of the sample boat. Second, a PTFE magnetic stir bar ($\Phi 4 \times 8$ mm) and a samarium-cobalt magnet ($\Phi 6 \times 2$ mm) are jointly slid into the vessel ~1 cm from the vessel mouth. Next, the sample boat ($\Phi 6 \times 10$ mm) is placed onto the stir bar and the outside magnet drives the stir bar and sample boat to the shoulder of the vessel (Figure 1). Third, a vacuum fitting and a sealed glass tube are used to isolate the vessel from moisture. Fourth, the magnet outside the vessel is fixed with a few turns of thread seal tape to improve the robustness of the apparatus as vessels are placed in tube rack, moved, and thermalized together in a water bath. The tape is used due to its low cost and stability at 90 °C.

OV and SV digestions were conducted using the I-shaped vessel. For the former, the operations are similar to those of the T-shaped vessel. The I-shaped vessel is connected to the vacuum line and the tape is removed from the vessel. A 90 °C water bath loaded to ensure a liquid level higher than the top of the sample boat and samarium-cobalt

magnet. Thermal equilibrium and evacuation are carried out for >0.5 h. The samariumcobalt magnet is removed by tweezers, the sample boat falls into the acid, and the reaction is activated. The CO₂ is continuously collected by trap. Samples analyzed by this procedure are denoted with an I as C1-I.

For the SV digestion, most I-shaped vessel operations are similar to those of the McCrea-type (Y-shaped) vessel, except for the activation of the carbonate-phosphoric acid reaction. Several IAEA-C1 samples were analyzed via this process, which are labeled C1-I-M.

Isotopic analyses

After the first liquid nitrogen trap collects the CO_2 , cleaning is conducted according to the procedure of Guo (2019). The CO_2 is released from the first trap by replacing the liquid nitrogen with -100 °C ethanol slush, passes through a -20 °C Porapak-Q trap), and is collected by a second trap for 25 min. Laslty, the CO_2 in the second trap is transferred into a glass tube for analysis.

The carbon, oxygen and clumped isotope compositions of CO_2 were determined by dual-inlet isotope ratio mass spectrometry (DI-IRMS) at the State Key Laboratory of Isotope Geochemistry (SKLaBIG) in the Guangzhou Institute of Geochemistry, Chinese Academy of Sciences (GIG-CAS). The Thermo Scientific 253 Plus DI-IRMS is configured with Faraday cup arrays of 3×10^8 , 3×10^{10} , 1×10^{11} and $1 \times 10^{13} \Omega$ to measure mass 44, mass 45, mass 46 and masses 47-49 (including mass 47.5), respectively. Before each acquisition, the beam intensity of mass 44 is automatically adjusted to 10 V for the sample and reference gas bellows. This signal intensity is sufficient to measure masses 47-49 on the DI-IRMS (Guo, Deng, and Wei 2019; Guo 2019). An analysis includes six acquisitions. Each acquisition consists of 10 cycles of 26 s.

The carbon, oxygen, and clumped isotope compositions are respectively expressed by:

$$\delta^{13}C = \left(\frac{R(^{13}C/^{12}C)_{sample}}{R(^{13}C/^{12}C)_{VPDB}} - 1\right) \times 1000 \%_{oo}$$
$$\delta^{18}O = \left(\frac{R(^{18}O/^{16}O)_{sample}}{R(^{18}O/^{16}O)_{VPDB}} - 1\right) \times 1000 \%_{oo}$$
$$\Delta_{47} = \left[\left(\frac{R^{47}}{R^{47*}} - 1\right) - \left(\frac{R^{46}}{R^{46*}} - 1\right) - \left(\frac{R^{45}}{R^{45*}} - 1\right)\right] \times 1000 \%_{oo}$$

where $R(^{13}C/^{12}C)$ and $R(^{18}O/^{16}O)$ are the abundance ratios of the heavy to light isotopes; R^{47} , R^{46} and R^{45} represent the abundance ratios of masses 47, 46 and 45 to mass 44, respectively; and R^{47*} , R^{46*} and R^{45*} are the 47/44, 46/44, and 45/44 ratios when all isotopes among isotopologues are in the stochastic distribution.

The raw Isodat data files are processed using Easotope (John and Bowen 2016) to export the carbon, oxygen and clumped isotope compositions (δ^{13} C, δ^{18} O and Δ_{47}). Detailed data processing parameters used in Easotope are described in Guo, Deng, and Wei (2019). The internal precisions (1 standard error) for δ^{13} C, δ^{18} O and Δ_{47} are ~0.01 $\%_{00}$. The raw Δ_{47} values are corrected for the linearity effect by the equilibrated gas lines

Sample ID	δ^{13} C (‰)	δ ¹⁸ Ο (‰)	$\Delta_{ extsf{47-AC}}$ (‰)	Number of samples
BACS-I	-1.88 ± 0.02	-4.51 ± 0.02	0.711 ± 0.018	17
BACS-T	-1.88 ± 0.03	-4.47 ± 0.03	0.702 ± 0.015	17
C1-I	2.44 ± 0.01	-2.37 ± 0.02	0.362 ± 0.016	20
C1-T	2.44 ± 0.01	-2.35 ± 0.02	0.369 ± 0.014	17
C1-I-M	2.41 ± 0.01	-2.39 ± 0.02	0.386 ± 0.014	7
C1-Y-M	2.39 ± 0.04	-2.40 ± 0.03	0.374 ± 0.011	11
ETH4-I	-10.20 ± 0.01	-18.90 ± 0.01	0.523 ± 0.007	6
ISCS-I	2.19 ± 0.01	-14.51 ± 0.02	0.546 ± 0.011	21
NBS18-I	-4.99 ± 0.01	-23.12 ± 0.02	0.440 ± 0.015	9
NBS19-I	1.95 ± 0.01	-2.25 ± 0.04	0.368 ± 0.007	8

Table 1. Summary of the stable and clumped isotope compositions using the Y-, T- and I-shaped vessels.

(EGL, see data in Table S1) (Huntington et al. 2009; Dennis et al. 2011) and calibrated to the absolute reference frame (ARF) by the empirical transfer function (ETF) (Dennis et al. 2011). Subsequently, the $\Delta_{47\text{-}ARF}$ values are calibrated to the $\Delta_{47\text{-}AC}$ values at the reference frame with the digestion temperature 25 °C using an acid fractionation factor of 0.068‰ (Guo, Deng, and Wei 2019; Dennis et al. 2011). A factor of 0.068‰ was experimentally determined with NBS19, IAEA-C1 and BACS digestion at 25 °C and 90 °C (Guo, Deng, and Wei 2019; Guo 2019), and is close to the theoretical value of 0.069‰ from Guo et al. (2009) and the measured value of 0.070‰ from Wacker, Fiebig, and Schoene (2013).

Results and discussion

Influence of acid digestion vessel

The carbon, oxygen, and clumped isotope compositions obtained using different vessels are displayed in Table S2. These data include (Table 1): (1) δ^{13} C, δ^{18} O and Δ_{47-AC} values from the I-shaped vessels (BACS-I and C-I) are consistent with those obtained from the T-shaped vessels (BACS-T and C1-T) within the error range; (2) few differences are observed for the I- and Y-shaped vessels, i.e., C1-I-M and C1-Y-M; and the (3) IAEA-C1 samples (C1-I and C1-I-M in Table 1) by OV and SV digestion in the I-shaped vessel show no significant differences in their stable and clumped isotope compositions.

The consistency between C1-I-M and C1-Y-M indicates that the I-shaped vessel is suitable replacement for the Y-shaped device for SV digestion. The I-shaped vessel may also be used in OV digestion because no significant differences were observed in its open and sealed modes (C1-I and C1-I-M in Table 1). The consistency between I- and T-shaped vessels (BACS-I, C1-I and BACS-T, C1-T) signifies that fractionation due to thermal disequilibrium between the two arms of the T-shaped vessel may be ignored when a 90 °C reaction temperature is used, as shown by Guo (2019) and Kluge and John (2015).

Overall, there are no significant differences among the δ^{13} C, δ^{18} O and $\Delta_{47\text{-AC}}$ values using these vessels. This result implies that the conditions in this study are far from those factors that can induce obviously different fractionations of the δ^{18} O and $\Delta_{47\text{-AC}}$ values in different vessels through the isotope exchange reaction between CO₂ and the H₃PO₄-H₂O system (Swart, Burns, and Leder 1991; Swart et al. 2019; Fernandez, Tang, and Rosenheim 2014). The short reaction time at the high temperature (90 °C) may play an important role in restraining isotope exchange. Similarly, different acid fractionation factors may exist in OV and SV digestion (Swart, Burns, and Leder 1991; Fernandez, Tang, and Rosenheim 2014), although this phenomenon was not observed in this study. This phenomenon is important for the carbonate-phosphoric acid method by showing that the acid fractionation factors of carbonates determined by SV digestion are suitable for isotopic calculation in OV digestion (Fernandez et al. 2016; Y. Guo, Deng, and Wei 2019; Kim, Coplen, and Horita 2015). However, due to difference in the acid fractionation factor observed in OV and SV digestions conducted using diverse containers (e.g., McCrea-type vessel and vial) (Swart, Burns, and Leder 1991; Fernandez et al. 2016; Swart et al. 2019) and that the factors affecting the fractionation processes in carbonate-phosphoric acid reactions (e.g., shape, volume and airtightness of the reaction containers) cannot be uniform in all laboratories, the acid fractionation factor differences between OV and SV digestions need to be further evaluated. The designed Ishaped vessel is suitable due to easy switching between OV and SV digestion.

Intercomparisons of δ^{13} C, δ^{18} O and Δ_{47-AC} among different studies

The published δ^{13} C, δ^{18} O and $\Delta_{47\text{-AC}}$ values of carbonate standards (Li et al. 2019; Fosu et al. 2019; Chang et al. 2020; Wacker, Fiebig, and Schoene 2013; Guo 2019; Fan et al. 2018; Fiebig et al. 2019; Henkes et al. 2013; Müller et al. 2017; Upadhyay et al. 2020) were selected a comparison with those in this study (Figure 2).

An obvious feature is that ETH-4 from different studies is statistically consistent in $\Delta_{47\text{-AC}}$ values but inconsistent in δ^{13} C and δ^{18} O values (Figure 2). Anomalies in δ^{13} C and δ^{18} O values are from three studies (Chang et al. 2020; Fan et al. 2018; Fiebig et al. 2019), which probably means that ETH-4 has heterogeneous stable isotope compositions or was contaminated during distribution. Another obvious difference occurs in δ^{18} O and $\Delta_{47\text{-AC}}$ values of NBS18 from Fosu et al. (2019) and this study (Figure 2). IAEA recommended δ^{18} O value of NBS18 as $-23.2 \pm 0.1 \%_{00}$, and the δ^{18} O values by Fosu et al. (2019) and this study are $-22.76 \pm 0.15 \%_{00}$ and $-23.12 \pm 0.02 \%_{00}$, respectively. The δ^{18} O value of NBS18 from Fosu et al. (2019) is ~0.4 \%_{00} higher than those from IAEA and this study. The result may be attributed to small internal volume of the leak-free vessel, long reaction time and the very negative δ^{18} O value of NBS18 (Fosu et al. 2019), which may affect the rate of oxygen isotopic exchange between CO₂ and H₃PO₄-H₂O system (Swart et al. 2019; Wacker, Fiebig, and Schoene 2013).

Except for ETH-4 and NBS18, δ^{13} C, δ^{18} O and $\Delta_{47\text{-AC}}$ values of other standards are comparable across laboratories (Figure 2). The interlaboratory consistency of the δ^{18} O values is better than for the $\Delta_{47\text{-AC}}$ values, possibly because the calibration processes and parameters of the δ^{18} O values are simpler than for the $\Delta_{47\text{-AC}}$ values (Swart, Burns, and Leder 1991; Ghosh, Garzione, et al. 2006; Groot 2009; Swart et al. 2019; Dennis et al. 2011; Chang et al. 2020). Particular attention should be paid to the symbols in Figure 2 that distinguish samples treated with OV or SV digestion and do not show the diverse reaction containers used for these samples (more details are provided in Table S3). IAEA-C1 and NBS19 notably indicate no differences in the δ^{18} O and $\Delta_{47\text{-AC}}$ values obtained from different apparatuses by OV and SV digestion (Figure 2 and Table S3). This observation suggests that fractionations of oxygen and clumped isotopes observed



Figure 2. Intercomparisons on δ^{13} C, δ^{18} O and Δ_{47-AC} of carbonate standards from different studies (Li et al. 2019; Fosu et al. 2019; Chang et al. 2020; Wacker, Fiebig, and Schoene 2013; Guo 2019; Fan et al. 2018; Fiebig et al. 2019; Henkes et al. 2013; Müller et al. 2017; Upadhyay et al. 2020). $\Delta\delta^{13}$ C = δ^{13} C - δ^{13} C_{x-average}, where δ^{13} C_{x-average} is the δ^{13} C average of all 'x' samples (such as BACS, ETH-4 and IAEA-C1) from Table 1 and Table S3, so are the definitions of $\Delta\delta^{18}$ O and $\Delta\Delta_{47-AC}$. Detailed data are provided in Table 1 and Table S3. The symbols denote different reaction vessels; open and solid symbols represent open vessel (OV) digestion and sealed vessel (SV) digestion, respectively. IV, TV, YV, CAB, KCD, BSV, and C/N represent I-shaped vessel, T-shaped vessel, Y-shaped vessel, common acid bath, Kiel carbonate device, break seal vessel, and common acid bath or Nu carb, respectively. The error bars represent 1 standard deviation.

in extreme conditions (e.g., reaction time of hundreds of hours, low-density phosphoric acid, or samples with extreme isotope compositions) (Swart et al. 2019) cannot be discerned for natural carbonates measured in normal conditions within analytical errors and demonstrates the reliability of the I-shaped vessel for measuring stable and clumped isotope compositions of carbonates by OV and SV digestion.

10 👄 X. LIU ET AL.

Conclusions

An improved apparatus for the phosphoric acid digestion of carbonates is proposed to determine carbon, oxygen and clumped isotope compositions. This apparatus is made of simple and common parts and accessories, including I-shaped vessels, sample boats, magnets and magnetic stir bars with specifically designed sizes, shapes and architectures. The isotopic data of international and interlaboratory carbonate standards obtained from this apparatus are consistent with those generated by other devices and from other laboratories. In addition, the straight and axisymmetric structure of the I-shaped vessel has advantages in manufacturing cost, operability, and ease of switching between modes. Therefore, this improved apparatus is a substitute for other vessels in the analysis of stable and clumped isotope composition of carbonates.

Acknowledgement

We thank Professor Stefano M. Bernasconi for providing ETH standards. This is contribution No.IS-2997 from GIG-CAS.

Funding

This work was supported by the National Natural Sciences Foundation of China (41722301 and 42003009), and the Key Special Project for Introduced Talents Team of Southern Marine Science and Engineering Guangdong Laboratory (Guangzhou) (GML2019ZD0308).

ORCID

Xi Liu (b) http://orcid.org/0000-0003-3244-5701 Wenfeng Deng (b) http://orcid.org/0000-0002-1722-8031

References

- Al-Aasm, I. S., B. E. Taylor, and B. South. 1990. Stable isotope analysis of multiple carbonate samples using selective acid extraction. *Chemical Geology: Isotope Geoscience Section* 80 (2): 119–25. doi:10.1016/0168-9622(90)90020-D.
- Bernasconi, S. M., I. A. Müller, K. D. Bergmann, S. F. M. Breitenbach, A. Fernandez, D. A. Hodell, M. Jaggi, A. N. Meckler, I. Millan, and M. Ziegler. 2018. Reducing uncertainties in carbonate clumped isotope analysis through consistent carbonate-based standardization. *Geochemistry, Geophysics, Geosystems: G(3)* 19 (9):2895–914. doi:10.1029/2017GC007385.
- Chang, B., W. F. Defliese, C. Li, J. Huang, A. Tripati, and T. J. Algeo. 2020. Effects of different constants and standards on the reproducibility of carbonate clumped isotope (Δ 47) measurements: Insights from a long-term dataset. *Rapid Communications in Mass Spectrometry: RCM* 34 (8):e8678. doi:10.1002/rcm.8678.
- Currie, B. S., P. J. Polissar, D. B. Rowley, M. Ingalls, S. Li, G. Olack, and K. H. Freeman. 2016. Multiproxy paleoaltimetry of the late oligocene-pliocene Oiyug Basin, Southern Tibet. *American Journal of Science* 316 (5):401–36. doi:10.2475/05.2016.01.
- Dean, J. R., M. D. Jones, M. J. Leng, S. R. Noble, S. E. Metcalfe, H. J. Sloane, D. Sahy, W. J. Eastwood, and C. N. Roberts. 2015. Eastern Mediterranean hydroclimate over the late glacial and holocene, reconstructed from the sediments of Nar Lake, Central Turkey, using stable

isotopes and carbonate mineralogy. *Quaternary Science Reviews* 124:162–74. doi:10.1016/j.quas-cirev.2015.07.023.

- Defliese, W. F., M. T. Hren, and K. C. Lohmann. 2015. Compositional and temperature effects of phosphoric acid fractionation on Δ_{47} analysis and implications for discrepant calibrations. *Chemical Geology* 396:51–60. doi:10.1016/j.chemgeo.2014.12.018.
- Deng, W., G. Wei, J. Zhao, and T. Zeng. 2019. Anthropogenic effects on tropical oceanic climate change and variability: An insight from the South China Sea over the past 2000. *Quaternary Science Reviews* 206:56–64. doi:10.1016/j.quascirev.2018.12.027.
- Deng, W., X. Liu, X. Chen, G. Wei, T. Zeng, L. Xie, and J. Zhao. 2017. A comparison of the climates of the medieval climate anomaly, little ice age, and current warm period reconstructed using coral records from the Northern South China Sea: Last millennium climates in the SCS. *Journal of Geophysical Research: Oceans* 122 (1):264–75. doi:10.1002/2016JC012458.
- Dennis, K. J., H. P. Affek, B. H. Passey, D. P. Schrag, and J. M. Eiler. 2011. Defining an absolute reference frame for "Clumped" isotope studies of CO₂. *Geochimica et Cosmochimica Acta* 75 (22):7117–31. doi:10.1016/j.gca.2011.09.025.
- Fan, M., S. A. Ayyash, A. Tripati, B. H. Passey, and E. M. Griffith. 2018. Terrestrial cooling and changes in hydroclimate in the continental interior of the United States across the eoceneoligocene boundary. GSA Bulletin 130 (7–8):1073–84. doi:10.1130/B31732.1.
- Fernandez, A., J. Tang, and B. E. Rosenheim. 2014. Siderite "Clumped" isotope thermometry: A new paleoclimate proxy for humid continental environments. *Geochimica et Cosmochimica Acta* 126:411–21. doi:10.1016/j.gca.2013.11.006.
- Fernandez, A., J. van Dijk, I. A. Müller, and S. M. Bernasconi. 2016. Siderite acid fractionation factors for sealed and open vessel digestions at 70 °C and 100 °C. *Chemical Geology* 444:180–6. doi:10.1016/j.chemgeo.2016.10.015.
- Fiebig, J., D. Bajnai, N. Löffler, K. Methner, E. Krsnik, A. Mulch, and S. Hofmann. 2019. Combined high-precision Δ_{48} and Δ_{47} analysis of carbonates. *Chemical Geology* 522:186–91. doi:10.1016/j.chemgeo.2019.05.019.
- Fosu, B. R., P. Ghosh, D. Mishra, Y. Banerjee, P. K, and A. Sarkar. 2019. Acid digestion of carbonates using break seal method for clumped isotope analysis. *Rapid Commun Mass Spectrom* 33 (2):203–14. doi:10.1002/rcm.8304.
- Friedman, I., J. O'Neil, and G. Cebula. 1982. Two new carbonate stable-isotope standards. *Geostandards and Geoanalytical Research* 6 (1):11–2. doi:10.1111/j.1751-908X.1982.tb00340.x.
- Ghosh, P., C. N. Garzione, and J. M. Eiler. 2006. Rapid uplift of the altiplano revealed through ¹³C-¹⁸O bonds in paleosol carbonates. *Science (New York, N.Y.)* 311 (5760):511–5. doi:10.1126/ science.1119365.
- Ghosh, P., J. Adkins, H. Affek, B. Balta, W. Guo, E. A. Schauble, D. Schrag, and J. M. Eiler. 2006. ¹³C-¹⁸O bonds in carbonate minerals: A new kind of paleothermometer. *Geochimica et Cosmochimica Acta* 70 (6):1439–56. doi:10.1016/j.gca.2005.11.014.
- Gilg, H. A., M. Boni, R. Hochleitner, and U. Struck. 2008. Stable isotope geochemistry of carbonate minerals in supergene oxidation zones of Zn-Pb deposits. *Ore Geology Reviews* 33 (2): 117-33. doi:10.1016/j.oregeorev.2007.02.005.
- Groot, P. A. de., ed. 2009. Chapter 4: Carbon. In *Handbook of stable isotope analytical techniques*, 229–329. Amsterdam: Elsevier. doi:10.1016/B978-0-444-51115-7.00004-8.
- Guo, W., J. L. Mosenfelder, W. A. Goddard, and J. M. Eiler. 2009. Isotopic fractionations associated with phosphoric acid digestion of carbonate minerals: Insights from first-principles theoretical modeling and clumped isotope measurements. *Geochimica et Cosmochimica Acta* 73 (24):7203–25. doi:10.1016/j.gca.2009.05.071.
- Guo, Y. 2019. Clumped isotope fractionation in biogenic carbonates (in Chinese). Beijing: University of Chinese Academy of Sciences. http://ir.gig.ac.cn/handle/344008/36583.
- Guo, Y., W. Deng, and G. Wei. 2019. Kinetic effects during the experimental transition of aragonite to calcite in aqueous solution: Insights from clumped and oxygen isotope signatures. *Geochimica et Cosmochimica Acta* 248:210–30. doi:10.1016/j.gca.2019.01.012.

12 🕢 X. LIU ET AL.

- Guo, Y., W. Deng, G. Wei, L. Lo, and N. Wang. 2019. Clumped isotopic signatures in land-snail shells revisited: possible palaeoenvironmental implications. *Chemical Geology* 519:83–94. doi: 10.1016/j.chemgeo.2019.04.030.
- Guo, Y., Deng, W., Wei, G., Chen, X., Liu, X., Wang, X., et al. (2020). Exploring the temperature dependence of clumped isotopes in modern Porites corals. Journal of Geophysical Research: Biogeosciences, 125, e2019JG005402. https://doi.org/10.1029/2019JG005402.
- Henkes, G. A., B. H. Passey, A. D. Wanamaker, E. L. Grossman, W. G. Ambrose, and M. L. Carroll. 2013. Carbonate clumped isotope compositions of modern marine mollusk and brachiopod shells. *Geochimica et Cosmochimica Acta* 106:307–25. doi:10.1016/j.gca.2012.12.020.
- Hu, B., J. Radke, H.-J. Schlüter, F. T. Heine, L. Zhou, and S. M. Bernasconi. 2014. A modified procedure for gas-source isotope ratio mass spectrometry: The Long-Integration Dual-Inlet (LIDI) methodology and implications for clumped isotope measurements. *Rapid Communications in Mass Spectrometry: RCM* 28 (13):1413–25. doi:10.1002/rcm.6909.
- Huntington, K. W., J. M. Eiler, H. P. Affek, W. Guo, M. Bonifacie, L. Y. Yeung, N. Thiagarajan, B. Passey, A. Tripati, M. Daëron, et al. 2009. Methods and Limitations of " 'clumped' CO2 isotope (Delta47) analysis by gas-source isotope ratio mass spectrometry. *Journal of Mass Spectrometry: JMS* 44 (9):1318–29. doi:10.1002/jms.1614.
- John, C. M., and D. Bowen. 2016. Community software for challenging isotope analysis: First applications of " 'easotope' to clumped isotopes". *Rapid Communications in Mass Spectrometry: RCM* 30 (21):2285–300. doi:10.1002/rcm.7720.
- Kim, S. T., T. B. Coplen, and J. Horita. 2015. Normalization of stable isotope data for carbonate minerals: Implementation of IUPAC guidelines. *Geochimica et Cosmochimica Acta* 158:276–89. doi:10.1016/j.gca.2015.02.011.
- Kluge, T., and C. M. John. 2015. Effects of brine chemistry and polymorphism on clumped isotopes revealed by laboratory precipitation of mono- and multiphase calcium carbonates. *Geochimica et Cosmochimica Acta* 160:155–68. doi:10.1016/j.gca.2015.03.031.
- Krishnamurthy, R. V., E. A. Atekwana, and H. Guha. 1997. A simple, inexpensive carbonate - phosphoric acid reaction method for the analysis of carbon and oxygen isotopes of carbonates. *Analytical Chemistry* 69 (20):4256–8. doi:10.1021/ac9702047.
- Leutert, T. J., P. F. Sexton, A. Tripati, A. Piasecki, S. L. Ho, and A. N. Meckler. 2019. Sensitivity of clumped isotope temperatures in fossil benthic and planktic foraminifera to diagenetic alteration. *Geochimica et Cosmochimica Acta* 257:354–72. doi:10.1016/j.gca.2019.05.005.
- Li, L., M. Fan, N. Davila, G. Jesmok, B. Mitsunaga, A. Tripati, and D. Orme. 2019. Carbonate stable and clumped isotopic evidence for late Eocene moderate to high elevation of the East-Central Tibetan plateau and its geodynamic implications. *GSA Bulletin* 131 (5–6):831–44. doi: 10.1130/B32060.1.
- Liu, X., W. Deng, and G. Wei. 2019. Carbon and oxygen isotopic analyses of calcite in calcitedolomite mixtures: Optimization of selective acid extraction. *Rapid Communications in Mass Spectrometry* 33 (5):411–8. doi:10.1002/rcm.8365.
- McCrea, J. M. 1950. On the isotopic chemistry of carbonates and a paleotemperature scale. *The Journal of Chemical Physics* 18 (6):849–57. doi:10.1063/1.1747785.
- Meckler, A. N., M. Ziegler, M. I. Millán, S. F. M. Breitenbach, and S. M. Bernasconi. 2014. Longterm performance of the Kiel carbonate device with a new correction scheme for clumped isotope measurements. *Rapid Communications in Mass Spectrometry: RCM* 28 (15):1705–15. doi: 10.1002/rcm.6949.
- Meinicke, N., S. L. Ho, B. Hannisdal, D. Nürnberg, A. Tripati, R. Schiebel, and A. N. Meckler. 2020. A robust calibration of the clumped isotopes to temperature relationship for foraminifers. *Geochimica et Cosmochimica Acta* 270:160–83. doi:10.1016/j.gca.2019.11.022.
- Müller, I. A., M. E. S. Violay, J.-C. Storck, A. Fernandez, J. van Dijk, C. Madonna, and S. M. Bernasconi. 2017. Clumped isotope fractionation during phosphoric acid digestion of carbonates at 70 °C. *Chemical Geology* 449:1–14. doi:10.1016/j.chemgeo.2016.11.030.
- Murray, S. T., M. M. Arienzo, and P. K. Swart. 2016. Determining the Δ_{47} acid fractionation in dolomites. *Geochimica et Cosmochimica Acta* 174:42–53. doi:10.1016/j.gca.2015.10.029.

- Passey, B. H., N. E. Levin, T. E. Cerling, F. H. Brown, and J. M. Eiler. 2010. High-temperature environments of human evolution in East Africa based on bond ordering in paleosol carbonates. *Proceedings of the National Academy of Sciences of the United States of America* 107 (25):11245–9. doi:10.1073/pnas.1001824107.
- Révész, K. M., and J. M. Landwehr. 2002. δ^{13} C and δ^{18} O isotopic composition of CaCO₃ measured by continuous flow isotope ratio mass spectrometry: Statistical evaluation and verification by application to devils hole core DH-11 calcite. *Rapid Communications in Mass Spectrometry* 16 (22):2102–14. doi:10.1002/rcm.833.
- Rosenbaum, J., and S. M. F. Sheppard. 1986. An isotopic study of siderites, dolomites and ankerites at high temperatures. *Geochimica et Cosmochimica Acta* 50 (6):1147. https://doi.org/10. 1016/0016-7037. doi:10.1016/0016-7037(86)90396-0.
- Rozanski, K., W. Stichler, R. Gonfiantini, E. M. Scott, R. P. Beukens, B. Kromer, and J. V. D. Plicht. 1992. The IAEA ¹⁴C intercomparison exercise 1990. *Radiocarbon* 34 (3):506–19. doi:10. 1017/S0033822200063761.
- Schmid, T. W., and S. M. Bernasconi. 2010. An automated method for "clumped-isotope' measurements on small carbonate samples. *Rapid Communications in Mass Spectrometry: RCM* 24 (14):1955–63. doi:10.1002/rcm.4598.
- Schmid, T. W., J. Radke, and S. Bernasconi. 2012. Clumped-isotope measurements on small carbonate samples with a Kiel IV carbonate device and a MAT 253 mass spectrometer. *Thermo Fisher Application Note*:30233. https://www.thermofisher.com/search/results?query=30233&persona=DocSupport
- Sharma, S. D., D. J. Patil, and K. Gopalan. 2002. Temperature dependence of oxygen isotope fractionation of CO₂ from magnesite-phosphoric acid reaction. *Geochimica et Cosmochimica Acta* 66 (4):589–93. https://doi.org/10.1016/S0016-7037. (01)00833-X. doi:10.1016/S0016-7037(01)00833-X.
- Spötl, C., and T. W. Vennemann. 2003. Continuous-flow isotope ratio mass spectrometric analysis of carbonate minerals. *Rapid Communications in Mass Spectrometry: RCM* 17 (9):1004–6. doi:10.1002/rcm.1010.
- Swart, P. K., S. J. Burns, and J. J. Leder. 1991. Fractionation of the stable isotopes of oxygen and carbon in carbon dioxide during the reaction of calcite with phosphoric acid as a function of temperature and technique. *Chemical Geology: Isotope Geoscience Section* 86 (2):89–96. https:// doi.org/10.1016/0168-9622. (91)90055-2. doi:10.1016/0168-9622(91)90055-2.
- Swart, P. K., S. T. Murray, P. T. Staudigel, and D. A. Hodell. 2019. Oxygen isotopic exchange between CO₂ and phosphoric acid: Implications for the measurement of clumped isotopes in carbonates. *Geochemistry, Geophysics, Geosystems* 20 (7):3730–50. doi:10.1029/2019GC008209.
- Upadhyay, D., J. Lucarelli, A. Arnold, R. Flores, H. Bricker, R. Ulrich, G. Jesmok, L. Santi, W. Defliese, R. A. Eagle, et al. 2020. Carbonate clumped isotope analysis (Δ_{47}) of 21 carbonate standards determined via gas source isotope ratio mass spectrometry on four instrumental configurations using carbonate-based standardization and multi-year datasets. Preprint. EarthArXiv doi:10.31223/X5VC80.
- Wacker, U., J. Fiebig, and B. R. Schoene. 2013. Clumped isotope analysis of carbonates: Comparison of two different acid digestion techniques. *Rapid Communications in Mass Spectrometry: RCM* 27 (14):1631–42. doi:10.1002/rcm.6609.