Organic Geochemistry 101 (2016) 38-48

Contents lists available at ScienceDirect

**Organic Geochemistry** 

journal homepage: www.elsevier.com/locate/orggeochem

## Differences in the thermal evolution of hopanes and steranes in free and bound fractions

### Liangliang Wu, Ansong Geng\*

The State Key Laboratory of Organic Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Wushan, Guangzhou 510640, PR China

#### ARTICLE INFO

Article history: Received 3 June 2016 Received in revised form 24 July 2016 Accepted 13 August 2016 Available online 17 August 2016

Keywords: Thermal evolution Bound fraction Hopanes Steranes Catalytic hydropyrolysis

#### ABSTRACT

Hopanes and steranes are the two of the most commonly used biomarker classes in the application of organic geochemistry to petroleum exploration. The same carbon skeletons also occur as a bound fraction, and can be used in a relatively high maturity range compared to their extractable (free) counterparts as a result of protection by the kerogen macromolecular structure. The pools of free and bound biomarkers are expected to be thermally degraded over geological time. There has been little work to address the chemical stabilities of hopanes and steranes in both free and bound forms. This study uses anhydrous pyrolysis to simulate the thermal evolutions of biomarkers from two Type II kerogens. The bound biomarkers within the above thermally altered kerogen residues were also released by catalytic hydropyrolysis and discussed in this study. The anhydrous pyrolysis results show that source-related parameters based on hopanes are more stable than those based on steranes. The hydropyrolysis results show that the bound hopane distributions are quite stable even at 460 °C (Easy%Ro = 2.86), while the bound steranes is lower than that of hopanes in both free and bound fraction which can be explained by the different chemical structure and mode of incorporation of their precursors into kerogen. © 2016 Elsevier Ltd. All rights reserved.

#### 1. Introduction

Hopanes and steranes are among the two most abundant biomarker classes in that portion of most sedimentary rock samples that is soluble in organic solvents because their biogenic precursors, namely the hopanoids and steroids, occur ubiquitously in living organisms (Peters et al., 2005). They are widely used in thermal maturity assessment, organofacies studies and oil-oil/oil-source rock correlations (Peters et al., 2005 and references therein). Generally, the extract of source rocks (bitumen) or the crude oil which contains free biomarkers will be routinely analyzed in molecular geochemistry studies. The use of such solvent-soluble (free) biomarkers is limited to the immature through oil-generation window, because molecular indices based on relative thermal stability and stereochemical changes are completed in that range. Hopane and sterane biomarkers are seldom utilized where high levels of thermal stress have been attained (van Graas, 1990; Farrimond et al., 1998; Liang and Chen, 2005).

A great number of studies have reported that hopanes and steranes also can be released from kerogen and/or solid bitumen by thermal pyrolysis, chemical degradation and catalytic hydropy-

http://dx.doi.org/10.1016/j.orggeochem.2016.08.009 0146-6380/© 2016 Elsevier Ltd. All rights reserved.

rolysis (Seifert, 1978; Tissot and Welte, 1978; Mackenzie et al., 1982a, 1982b; Mycke and Michaelis, 1986; Eglinton and Douglas, 1987; Mycke et al., 1987; de Leeuw et al., 1989; Abbott et al., 1990; Hoffman et al., 1992; Richnow et al., 1992; Adam et al., 1993; Love et al., 1995; Muhammad and Abbott, 2012). Various conditions (temperature and time) have been used for thermal pyrolysis. Generally, the temperature used for thermal pyrolysis is > 250 °C. Chemical degradation is a mild degradation method. Many kinds of chemical reagents were used for chemical degradation, such as Li/EtND<sub>2</sub> and Ni(0)cene/LiAlD<sub>4</sub> (Hoffman et al., 1992; Richnow et al., 1992), and BCl<sub>3</sub>/LiAlH<sub>4</sub> (Michaelis et al., 1988; Jenisch et al., 1990). The nature of bonding inferred from the reagents used in chemical degradation processes (Mycke and Michaelis, 1986; Hoffman et al., 1992; Richnow et al., 1992) and the pyrolysis temperatures at which products are obtained (Seifert, 1978; Eglinton and Douglas, 1987) strongly indicated that such hopanes and steranes were covalently bound. The covalently bound hopanes and steranes are different from solvent soluble hopanes and steranes in composition and distribution characteristics (Tissot and Welte, 1978; Love et al., 1995, 2005; Lockhart et al., 2008; Liao et al., 2012; Wu et al., 2013). Because the steranes and hopanes obtained by the different methods (extraction and pyrolysis) are chemically distinct, previous workers have developed the idea of bound biomarkers and free biomarkers (e.g.,







<sup>\*</sup> Corresponding author. Fax: +86 20 87685791. *E-mail address:* asgeng@gzb.ac.cn (A. Geng).

Abbott et al., 1990). Additionally, previous work also points out that the compounds released by various degradation processes were not covalently bonded to macromolecular structure but rather were occluded within it (Strausz et al., 1999a, 1999b; Liao and Geng, 2002; Russell et al., 2004; Gordadze et al., 2015; Snowdon et al., 2016). Thus, it is not clear that bound moieties released by pyrolysis and HyPy technique can be readily differentiated from occluded species (Snowdon et al., 2016). Therefore, most workers use the term bound biomarker (including both covalently bound and occluded biomarkers) to represent the products released from geological macromolecular structures.

The thermal evolution of bound biomarkers may be retarded to some extent compared to their free counterparts due to protection afforded by the macromolecular structure (Tissot and Welte, 1978; Love et al., 1995, 2005; Abbott et al., 2001; Lockhart et al., 2008; Muhammad and Abbott, 2012). This is especially apparent at low maturity stages (Love et al., 1995; Murray et al., 1998; Meredith et al., 2008; Wu et al., 2013). Maturity-related parameters based on bound biomarkers show constant values with increasing maturity in highly mature source rocks ( $T_{max} > 460 \circ C$ ) (Lockhart et al., 2008; Wu et al., 2013; Liao et al., 2015) and artificially matured solid bitumen (simulation temperature > 350 °C) (Liao et al., 2012). It therefore seemed that assessment of maturity based on the bound biomarkers at high- to over maturity stage was not possible. However, recently, Liao et al. (2012) reported that sourcerelated biomarkers, including the distribution of regular steranes, are quite stable in hydropyrolysis (HyPy) products from laboratory heated bitumens of various maturities (Easy%Ro = 1.08-2.86). Wu et al. (2013) also found that the bound source-related biomarkers from the Permian Dalong Formation in the Sichuan Basin, China at the mature Guangyuan outcrop sections and overmature Wangcang outcrop sections are very similar. Those data indicate that biomarkers released from kerogen can be used in oil-source correlation studies in highly mature to over-mature areas. The bound biomarkers released by HyPy also were also used for oil-source correlation in the Majiang paleo-reservoir (Fang et al., 2014).

Although it was widely accepted that source-related biomarkers within kerogen (bound fraction) are more stable than their free counterparts, we still know little about the thermal stability and thermal evolution of different kinds of biomarkers. The Dalong Formation was deposited under very similar conditions, but shows quite different maturity at different locations. Biomarkers obtained by both Soxhlet extraction and HyPy from these rocks have been well investigated (Wu et al., 2013). In the present study, two early mature Type II kerogens from the Dalong Formation were selected for artificial thermal alteration followed by HyPy experiments on the kerogen residues. The aim of this study is to compare the effects of thermal maturation on source-related biomarker parameters based on free biomarkers and bound biomarkers. This is very important for the interpretation of bound biomarkers released from high to over-mature samples.

#### 2. Experimental

#### 2.1. Samples and preparation

The investigated samples are siliceous rock (GY-8) and mudstone (GY-17) selected from the fresh outcrop of the Permian

Table 1
---------

Basic information on the investigated source rocks.

Dalong Formation in Sichuan Basin, China (Lat.  $32^{\circ}19'11''N$ , Long.  $105^{\circ}27'18''E$ ). None of the selected samples were strongly weathered. The Dalong Formation contains interbedded marine dark siliceous rock and mudstones, and mainly occurs in the Guangyuan-Wangcang trough and Western Hubei-Chengkou trough (Wang et al., 2006; Tenger et al., 2008; Li et al., 2009). The basic geochemical characteristics of the samples have already been presented by Wu et al. (2012, 2013) and are summarized in Table 1. The TOC (total organic carbon) content is 8.76% for GY-8 and is 3.67% for GY-17. Their measured vitrinite reflectance are in the range of 0.58–0.68 %Ro. They also have very similar values of  $T_{max}$  (438 °C) and hydrogen Index (343–357 mg/g TOC). The investigated samples are classified as early mature Type II kerogens.

The source rock samples were crushed to  $\leq$  80 mesh powders. Then, kerogen isolation was performed by acid treatment. First the ground samples were treated with 3 M HCl (80 °C for 4 h), and washed with distilled water and separated by centrifugation. Then a mixture of 1 M HCl and HF (1:1, v:v) was used (80 °C for 4 h) and recovered by additional water washing and centrifugation. Before the artificial thermal alteration experiment was conducted, the isolated kerogen from GY-8 was Soxhlet extracted using an azeotropic ternary solvent system (benzene/ acetone/methanol, 5:5:2, v:v:v) for 2 weeks to remove the residual bitumen, while the GY-17 kerogen was not Soxhlet extracted with the ternary solvent (Fig. 1). The aim of different kerogen preparations was to study the influence of free bitumen on the thermal evolution of bound biomarkers within kerogen.

#### 2.2. Artificial thermal alteration experiments

Tubes were loaded with about 0.6 g of the prepared kerogen (GY-8 and GY-17) and sealed under an inert nitrogen atmosphere. Isothermal pyrolysis was then performed for 72 h at 350, 380, 400, 430 and 460 °C, and each sample was heated from room temperature to the final temperature. The calculated vitrinite reflectance (% Ro) by the Easy%Ro method (Sweeney and Burnham, 1990) for the kerogens heated at 350 °C, 380 °C, 400 °C, 430 °C and 460 °C were 1.08, 1.49, 1.74, 2.27 and 2.86, respectively (Table 2). After thermal pyrolysis, the pyrolysate (pyrolysis product) was recovered by repeated dichloromethane (DCM) sonication. Prior to HyPy, kerogen residues were further Soxhlet extracted for two weeks using the ternary solvent azeotrope to remove soluble organic matter (Fig. 1).

#### 2.3. Hydropyrolysis experiments

Catalytic hydropyrolysis (HyPy) of the samples was conducted using the apparatus and procedure described in detail elsewhere (Snape et al., 1989; Love et al., 1995; Wu et al., 2013). The solvent-extracted kerogen was impregnated with an aqueous solution of ammonium dioxydithiomolybdate [(NH<sub>4</sub>)<sub>2</sub>MoO<sub>2</sub>S<sub>2</sub>] to give a nominal loading of 5 wt% molybdenum as catalyst. Then, the catalyst loaded kerogen was put into the HyPy reactor tube. The HyPy experiment involves two steps at the same hydrogen pressure (15.0 MPa) and hydrogen flow (4 L/min). The first step was heating from ambient temperature to 300 °C (5 min) at 250 °C/min. The aim of the first step is to remove the weaker covalent bonds

Sample ID	Lithology	δ <sup>13</sup> C (‰)	%Ro	TOC (%)	S1	S2	$T_{\max}$ (°C)	HI
GY-8	Silicalite	-27.6	0.58	8.75	1.22	30.1	438	343
GY-17	Mudstone	-27.8	0.68	3.67	0.61	13.3	438	357

Data from Wu et al. (2013).



Fig. 1. Sample preparation and experiment process scheme.

Table 2

The TOC-normalized yields of the total EOM and the saturated, aromatic and polar fractions for pyrolysates of kerogen at different temperatures and their corresponding hydropyrolysates.

Sample no.	Temp (°C)	Easy%Ro	Anhydrous pyrolysis (mg/g TOC)			HyPy (mg/g TO	HyPy (mg/g TOC)		
			Total EOM	Saturated	Aromatic	Total EOM	Saturated	Aromatic	
GY-8 series									
Original	-	0.58	56.0	4.1	17.3	497.7	112.8	171	
GY-350	350	1.08	15.5	0.6	6.7	45.1	1.2	18.9	
GY-380	380	1.49	13.5	0.5	5.7	37.5	0.3	16	
GY-400	400	1.74	2.7	1.1	0.8	16	0.6	8	
GY-430	430	2.27	1.5	0.1	0.7	2.4	0.4	0.4	
GY-460	460	2.86	0.9	0.2	0.3	2.6	0.7	0.4	
GY-17 series									
Original	-	0.68	45.4	8.3	20.6	205	19.3	163.3	
GY-350	350	1.08	21.8	8.7	3.6	68.9	21.0	33.0	
GY-380	380	1.49	8.7	2.2	2.2	56.6	14.9	17.9	
GY-400	400	1.74	7.3	5.1	-	56.4	14.8	38.6	
GY-430	430	2.17	-	-	-	99.9	6.1	46.4	
GY-460	460	2.86	-	-	-	50.8	5.2	20.6	

"-" = undetectable.

leaving the stronger bonds for HyPy (Love et al., 1995, 1997, 1998). After this step, the silica gel in the collecting tube was replaced. The second heating run was from ambient temperature to  $250 \,^{\circ}C$  (5 min) at  $300 \,^{\circ}C/min$  and then to  $520 \,^{\circ}C$  (5 min) at  $8 \,^{\circ}C/min$ . The hydropyrolysates (hydropyrolysis products) were collected in a liquid nitrogen cold trap and recovered by DCM/methanol (93:7, v:v) for subsequent analysis (Meredith et al., 2004).

#### 2.4. Product analysis

The hydropyrolysates of the thermally altered kerogen residues were conducted as follows. First, the asphaltenes were precipitated from the products by adding 50:1 (v:v) cold *n*-hexane, and then removed by centrifugation. The maltene fractions were then fractionated by silica/alumina (3:1, v:v) column chromatography into saturated, aromatic and polar fractions by elution with *n*-hexane, DCM/*n*-hexane (3:1, v:v) and DCM/methanol (2:1, v:v), respectively.

The saturated biomarkers were analyzed using a Thermo Scientific Trace GC Ultra gas chromatograph coupled to a Thermo Scientific Trace DSQ II mass spectrometer. A DB-1 fused silica capillary column (60 m × 0.32 mm i.d. × 0.25 µm film thickness) was used. The GC oven was held isothermal at 70 °C for 2 min, programmed to 290 °C at 4 °C/min, with a final hold time of 30 min. Helium was used as carrier gas with a constant flow rate of 1.5 mL/min. The ion source temperature was 250 °C, and the temperature of injector was 290 °C. The ion source was operated in the electron impact (EI) mode with electron energy of 70 eV. GC–MS analyses were operated in SIM mode. The selected ions monitored included *m*/*z* 191 (hopanes) and *m*/*z* 217 (steranes).

#### 3. Results and discussion

Table 2 shows the total EOM (extractable organic matter), saturated, and aromatic yields normalized to TOC from anhydrous pyrolysis products (pyrolysates) of kerogen at different temperatures and their corresponding HyPy products (hydropyrolysates). The Rock Eval data of thermal altered GY-8 kerogen residues were also shown in Table 3. The yield of kerogen pyrolysates ( $C_{15+}$ ) decreases from 15.5 mg/g TOC (350 °C) to 0.9 mg/g TOC (460 °C) for GY-8 kerogen and decreases from 21.8 mg/g TOC (350 °C) to 7.3 mg/g TOC (400 °C) for GY-17 kerogen with increasing tempera-

 Table 3

 Rock Eval data for the artificial thermal altered kerogens of GY-8.

Sample ID	Temp (°C)	S1 (mg/g kerogen)	S2 (mg/g kerogen)	$T_{\max}$ (°C)	HI
GY-350	350	2.24	7.29	482	11
GY-380	380	1.15	5.61	562	8
GY-400	400	1.15	3.56	581	5
GY-430	430	1.98	2.05	602	3
GY-460	460	1.05	1.27	605	2

ture due to more and more severe thermal decomposition. The yield of hydropyrolysates ( $C_{15+}$ ) also decreases from 45.1 mg/g TOC (350 °C) to 2.6 mg/g TOC (460 °C) for GY-8 kerogen, while they are nearly constant (ca. 60 mg/g TOC) for GY-17 with increasing temperature. For the same sample, the yields of saturated and aromatic hydrocarbon for both pyrolysates and hydropyrolysates also decrease with increasing temperature. However, the yield of hydropyrolysates for GY-17 kerogen residues is higher than that for GY-8 kerogen residues at the same temperature. This is probably because the free bitumen that is not removed from the GY-17 kerogen was incorporated into the kerogen residue during pyrolysis (cf. secondary reactions as per Vu et al., 2008) and released by the following HyPy experiment (see also following discussion on the distribution of  $C_{27}$ – $C_{29}$  steranes for hydropyrolysates).

# 3.1. Results of laboratory heating experiment on GY-8 and GY-17 kerogens

Fig. 2 shows the hopane and sterane profiles for the pyrolysates of GY-8 and GY-17 kerogens at various temperatures (GY-8 series and GY-17 series). For GY-8 series, most of the biomarkers including hopanes and steranes were not detected or their relative concentration is very low in its pyrolysates at all temperature. This is probably because the pyrolysis-released biomarkers were destroyed by thermal stress. Previous studies also found that hopanes and steranes are absent in both hydrous and anhydrous pyrolysis of kerogen at temperature above 330 °C (Comet et al., 1986; Lewan et al., 1986; Lewan, 1997). Comet et al. (1986) pointed out that the existence of elemental sulfur within kerogen can considerably destroy free polycyclic biomarkers. The work by Abbott et al. (1985) also pointed out that elemental sulfur can catalyze the degradation of steranes and aromatic steroids. Of course, the inefficient recovery of sonication with DCM might be another possible reason for the absence of biomarkers in pyrolysates of GY-8 kerogen at various temperatures. However, hopanes and steranes were detected in pyrolysates of GY-17 kerogen at temperatures from 350 °C to 400 °C. Only trace of hopanes and steranes were detected at 430 °C. Hopanes and steranes were not observed in pyrolysate at 460 °C due to the severe thermal decomposition. Since the free biomarkers in bitumen were not removed from GY-17 kerogen before anhydrous pyrolysis, the biomarkers detected in pyrolysates of GY-17 kerogen should be the mixed thermally altered products of free bitumen from source rock and liquid products released from kerogen. The work by Murray et al. (1998) demonstrated that free bitumen is a more important pool of hopanes and steranes than kerogen at the beginning of the oil-generation window (%Ro = 0.5–0.6). Additionally, hopanes and steranes were not detected or their relative concentration is very low in the pyrolysates released from GY-8 kerogen in which the free bitumen had been removed. Thus, the hopanes and steranes detected in pyrolysates of GY-17 kerogen are probably mainly from free bitumen.

The distribution of  $C_{27}$ - $C_{29}$  regular steranes is very stable in the oil-generation window and can be used for oil-oil and oil-source rock correlation and to assess source and depositional environ-

ment (Seifert et al., 1984; Peters et al., 1989, 2005). With the temperature increasing from 350 °C to 400 °C, the relative abundance of  $C_{27}$  5 $\alpha$ ,14 $\alpha$ ,17 $\alpha$ , 20R-cholestane for GY-17 kerogen increases from 0.49 to 0.64 (Table 4), while  $C_{28}$  5 $\alpha$ ,14 $\alpha$ ,17 $\alpha$ , 20R-24-methylcholestane decreases from 0.30 to 0.19. Previous studies on the pyrolysis of source rocks or kerogens have also reported that the relative abundance of  $C_{27}$  regular steranes gradually increases with simulation temperature (Comet et al., 1986; Lewan et al., 1986; Lu et al., 1989; Lewan, 1997).

 $C_{21}$  and  $C_{22}$  steranes (pregnane and homopregnane and their geological isomers) may originate from the hormones, pregnanol and pregnanone (de Leeuw and Baas, 1986) or from thermal cracking of  $C_{27}$ – $C_{29}$  regular steroids (Huang et al., 1994; Abbott et al., 1995). The ratio of  $C_{21}/C_{22}$  steranes (S21/S22) can also be used for oil-source correlation (Huang et al., 1994; Fang et al., 2011; Wu et al., 2012): S21/S22 is in the range 1.22–1.45 for pyrolysates of the GY-17 series.

The ratios based on free hopanes and tricyclic terpanes, such as gammacerane/C<sub>30</sub>  $17\alpha(H)$ ,  $21\beta(H)$ -hopane (Gam/H30), C<sub>29</sub>  $17\alpha(H)$ , 21β(H)-hopane to C<sub>30</sub> 17α(H),21β(H)-hopane (H29/H30), C<sub>23</sub> tricyclic terpane/ $C_{30}$  17 $\alpha$ (H),21 $\beta$ (H)-hopane (TT23/H30),  $C_{23}$  tricyclic terpane to C<sub>23</sub> and C<sub>24</sub> tricyclic terpanes (TT23/(TT23 + TT24)) are also commonly used to assess the source (Peters et al., 2005). Except for TT23/H30, other ratios are all quite stable with increasing temperature from 350 °C to 400 °C for pyrolysates of the GY-17 series (Table 5). For example, the value of H29/H30 is in the range of 0.65-0.75, and the value of Gam/H30 is in the range of 0.09-0.12. The value of TT23/H30 is 0.78 for thermally altered kerogen of GY-17 at 350 °C, while it is nearly 0.20 at 380 °C and 400 °C (Table 5). Thus, the source related biomarkers parameters based on hopanes seems to be more stable than those based on steranes in the pyrolysates of GY-17 kerogen at various temperatures.

#### 3.2. Results of HyPy experiment on artificially matured kerogens

The HyPy technique was used to release the bound steranes and hopanes from laboratory heated products from the two investigated Dalong Formation kerogens (GY-8 and GY-17) at various temperatures. Their bound hopane  $(m/z \ 191)$  and sterane profiles (m/z 217) are shown in Fig. 3. For the same kerogen series, the distributions of bound hopanes released from the thermally altered kerogens at all temperatures are quite similar, whereas the relative abundance of bound  $C_{29}$  5 $\alpha$ ,14 $\alpha$ ,17 $\alpha$ , 20R-ethylcholestane are gradually decreased with increasing temperature (Fig. 3). However, the distributions of bound C27-C29 regular steranes are quite similar for pyrolysis temperatures < 430 °C (Easy%Ro = 2.27). This is consistent with our previous results for a natural maturity sequence in the Dalong Formation in which the distributions of  $C_{27}$ - $C_{29}$  regular steranes in hydropyrolysates from overmature samples (%Ro = 1.8; Wu et al., 2015) are still similar to those in the mature samples (Wu et al., 2013). These results confirm the suggestion that the bound biomarkers are protected to some extent from thermal degradation by the macromolecular structure of kerogen (Love et al., 1998; Lockhart et al., 2008; Liao et al., 2012; Wu et al., 2013).



**Fig. 2.** Selected ion chromatograms (SIM) for m/z 191 and m/z 217 of the saturated fractions in pyrolysates of thermally altered kerogens: (a) is the siliceous rocks series (original kerogen is GY-8) and (b) is the mudstone series (original kerogen is GY-17). Peak identifications are shown in Table 2.

Fig. 4 is a ternary plot of the  $C_{27}$ - $C_{29}$  regular sterane distributions for hydropyrolysates from two original samples and their thermally altered residues. For the GY-8 GY-17 series, the relative abundance of bound  $C_{29}$  5 $\alpha$ ,14 $\alpha$ ,17 $\alpha$ , 20R-ethylcholestane decreases, but  $C_{27}$  5  $\alpha$  , 14  $\alpha$  , 17  $\alpha$  , 20 R-cholestane increases with increasing temperature. Nevertheless, the distributions of C27-C29 regular steranes in hydropyrolysates are still similar to the original sample up to a special temperature (400 °C for both GY-8 and GY-17 series) because of the protection by the macromolecular network (Fig. 3). The only difference was their decreasing rate of bound  $C_{29}$  5 $\alpha$ , 14 $\alpha$ , 17 $\alpha$ , 20R-ethylcholestane. It seems the decreasing rate of bound  $C_{29}$  5 $\alpha$ ,14 $\alpha$ ,17 $\alpha$ , 20R-ethylcholestane for GY-8 series are faster than that for GY-17 series. Since the lithology of GY-8 sample (siliceous rocks) is different from that of GY-17 (mudstone), their maceral compositions should be also quite different. Thus, this may suggest that the degree of protection of the macrostructure on the bound steranes can differ with kerogen, particularly between kerogens with different maceral composition.

The distributions of bound  $C_{27}$ - $C_{29}$  regular steranes in the thermally altered kerogens of GY-17 are quite different from those of

the original kerogen (Fig. 4b). In contrast, the distributions of  $C_{27}$ - $C_{29}$  regular steranes in the thermally altered kerogens of GY-8 are similar to those in the original kerogen, at least below the temperature of 430 °C. As already mentioned above, the original kerogen of GY-8 was extracted (benzene/acetone/methanol = 5:5:2, v:v:v) for two weeks, while the kerogen of GY-17 was not extracted. Thus, the free bitumen, especially the polar fraction containing hopanoids and steroids, was not removed from the kerogen concentrate of GY-17 and may have been incorporated into the kerogen during pyrolysis (cf. secondary reactions as per Vu et al., 2008). Some researchers have suggested that variable proportions of the steroid and hopanoid pools in geological samples are first incorporated into kerogen and later released as hopanes and steranes during catagenesis (Mycke et al., 1987; Sinninghe Damsté and de Leeuw, 1990). Therefore, the distributions of bound biomarkers within the kerogens of GY-17 at all temperatures may have been influenced by incorporation of free bitumen.

The ratios based on hopanes and tricyclic terpanes remain unified for all hydropyrolysates in each kerogen series (shown in Fig. 5). For example, Gam/H30 is in the range of 0.23–0.28 for the

Table 4Biomarker assignments.

Peak	Compound
TT23	C <sub>23</sub> tricyclic terpane
TT24	C <sub>24</sub> tricyclic terpane
TT25	C <sub>25</sub> tricyclic terpane
TT26	C <sub>26</sub> tricyclic terpane
TT28	C <sub>28</sub> tricyclic terpane
TT29	C <sub>29</sub> tricyclic terpane
tT24	C <sub>24</sub> tetracyclic terpane
H29	$C_{29}$ 17 $\alpha$ (H),21 $\beta$ (H)-hopane
H30	$C_{30}$ 17 $\alpha$ (H),21 $\beta$ (H)-hopane
H31 22S	C <sub>31</sub> 17α(H),21β(H),22S-hopane
H31 22R	C <sub>31</sub> 17α(H),21β(H),22R-hopane
H32	$C_{32}$ 17 $\alpha$ (H),21 $\beta$ (H)-hopane
H33	$C_{33}$ 17 $\alpha$ (H),21 $\beta$ (H)-hopane
Gam	Gammacerane
Mor	$C_{30}$ 17 $\beta$ (H),21 $\alpha$ (H)-hopane
Ts	C <sub>27</sub> 17α-22,29,30-trisnorhopane
Tm	C <sub>27</sub> 18α-22,29,30-trisnorhopane
C27	C <sub>27</sub> 5α,14α,17α, 20R-cholestane
C28	C <sub>28</sub> 5α,14α,17α, 20R-methylcholestane
C29	C <sub>29</sub> -5α,14α,17α, 20R-ethylcholestane
C27 βαS	C <sub>27</sub> 13β,17α,20S-diacholestane
C27 βαR	C <sub>27</sub> 13β,17α,20R-diacholestane
S21	C <sub>21</sub> sterane
S22	C <sub>22</sub> sterane

GY-8 hydropyrolysates and in the range of 0.17–0.20 for the GY-17 hydropyrolysates (Fig. 5b, Table 6). The ratio of S21/S22 is also stable for the all hydropyrolysates in each kerogen series. S21/S22 is in the range 1.38–1.51 and 1.62–1.83 for hydropyrolysates of the GY-8 and GY-17 series, respectively (Fig. 5a, Table 6). However, most of the above biomarker ratios in the hydropyrolysates from the GY-8 series are similar to those in the original kerogen hydropyrolysate, while some biomarker ratios (such as the ratio of S21/S22 and H29/H30) in the hydropyrolysates from GY-17 series are quite different from those in the original kerogen hydropyrolysate (Table 6). This further confirms our previous speculation that the distributions of bound biomarkers within the kerogens of GY-17 at all temperatures may have been influenced by the incorporation of free bitumen.

Based on the above data, it is clear that ratios based on bound hopanes and tricyclic terpanes, such as H29/H30 and Gam/H30, are quite stable from 350 °C to 460 °C (Fig. 5), while the distribution of  $C_{27}$ - $C_{29}$  regular steranes are thermally altered with increasing simulation temperature (Fig. 4). Thus the regular steranes should be more easily destroyed by thermal stress than hopanes and tricyclic terpanes. Additionally, the ratio S21/S22 is also based on steranes, but it is quite stable at all temperatures. The following discussion explains why the thermal evolution trend for S21/S22 differs from that of the  $C_{27}$ - $C_{29}$  regular steranes.

#### 3.3. Why biomarker parameters for bound hopanes and steranes differ

The results from anhydrous pyrolysis reveal that the thermal evolution of steranes and hopanes differ in the free fraction. Different thermal evolution trends between steranes and hopanes released by closed-system pyrolysis have been observed previously. Hydrous and anhydrous pyrolysis experiments by Lewan (1997) and Liang et al. (2015) reveal that  $C_{27}$ - $C_{29}$  regular sterane are more easily altered by heating than hopane ratios such as Gam/H30 and H29/H30. Generally, the thermal stability of free biomarkers is controlled by their chemical structure. Most source-related biomarker parameters are the ratios of sterane or hopane homologs, because adjacent homologs show similar thermal stability (Peters et al., 2005). However, extensive thermal cracking can still influence those parameters, due to slight differences in degradation rates of biomarker homologs. Laboratory thermal alteration of triaromatic steroids by Beach et al. (1989) indicated that compounds with longer side chains show faster degradation than the shorter chain homologs. Thus, biomarkers with longer side chains should be more easily thermally degraded than those with shorter side chains. The  $C_{27}$ - $C_{29}$  regular steranes have longer side chains than most commonly used biomarkers. By this line of reasoning,  $C_{27}$ - $C_{29}$  regular sterane distributions are more easily thermally altered than parameters based on biomarkers with shorter branched chains (such as H29/H30 and Gam/H30) in the free fraction.

The influence of thermal stress on bound biomarkers is limited due to protection by the macromolecular structure (Tissot and Welte, 1978; Love et al., 1995, 2005). Sugden and Abbott (2002) reported that epimerization of bound hopanoids still occurs at C-17, C-21 and C-22 during closed system pyrolysis. Lockhart et al. (2008) also suggested that bound constituents undergo the same maturation reaction pathways as their free counterparts. However, the chemical structures of bound biomarkers are quite different from their corresponding free forms, since linkages via covalent bonds link the bound biomarker precursors to the kerogen. Thus, the incorporation of individual biomarkers in kerogen may also influence their thermal stability.

Previous labeling experiments with deuterium indicated that individual biomarkers are incorporated into the macromolecular network via functional groups in the biological precursors (Mycke and Michaelis, 1986; Richnow et al., 1992; Stalker et al., 1998). Bacteriohopanetetrol and similar highly functionalized compounds are precursors of hopanes (Rohmer et al., 1984;

Table	5
-------	---

The biomarker	parameters in	pyrolysates	of the	thermal	altered	kerogen	of	GY-1	17.
---------------	---------------	-------------	--------	---------	---------	---------	----	------	-----

Sample ID	GY-17 series				
Temp (°C)	350	380	400	430	460
$C_{27}/(C_{27} + C_{28} + C_{29})$	0.49	0.58	0.64	-	-
$C_{28}/(C_{27} + C_{28} + C_{29})$	0.30	0.21	0.19	-	-
$C_{29}/(C_{27} + C_{28} + C_{29})$	0.21	0.20	0.17	-	-
S21/S22	1.45	1.22	1.27	-	-
$C_{27} \beta \alpha R / C_{27}$	0.22	0.15	0.17	-	-
$C_{29} \alpha\beta\beta/(\alpha\beta\beta + \alpha\alpha\alpha)$	0.48	0.45	0.45	-	-
$C_{29} 20S/(20S + 20R)$	0.49	0.46	0.51	-	-
TT23/(TT23 + TT24)	0.84	0.78	0.73	-	-
Ts/(Ts + Tm)	0.42	0.40	0.45	-	-
TT23/H30	0.78	0.22	0.20	-	-
H29/H30	0.75	0.65	0.71	-	-
H31 22S/(22S + 22R)	0.60	0.62	0.59	-	-
Gam/H30	0.09	0.12	0.10	-	-

"-" = undetectable.



**Fig. 3.** Selected ion chromatograms (SIM) for *m*/*z* 191 and *m*/*z* 217 of the saturated fractions in hydropyrolysates of thermal altered kerogens: (a) is the siliceous rocks series (original kerogen is GY-8) and (b) is the mudstone series (original kerogen is GY-17). Peak identifications are shown in Table 2.



**Fig. 4.** Ternary plot of C<sub>27</sub>-C<sub>29</sub> ααα20R sterane distributions in the hydropyrolysate of the original sample and its thermal maturation residues: (a) is the series of GY-8 and (b) is the series of GY-17.



Fig. 5. Variations of source-related biomarker parameters in hydropyrolysates with increasing temperature.

Table 6			
The biomarker para	meters in hydropyro	olysates of the the	rmal altered kerogens

Sample ID	GY-8						GY-17					
Temp (°C)	Original	350	380	400	430	460	Original	350	380	400	430	460
$C_{27}/(C_{27} + C_{28} + C_{29})$	0.34	0.32	0.35	0.35	0.37	0.41	0.52	0.36	0.37	0.35	0.39	0.43
$C_{28}/(C_{27} + C_{28} + C_{29})$	0.21	0.23	0.27	0.26	0.30	0.29	0.16	0.26	0.28	0.30	0.38	0.29
$C_{29}/(C_{27} + C_{28} + C_{29})$	0.45	0.45	0.38	0.39	0.33	0.30	0.32	0.35	0.35	0.35	0.34	0.28
S21/S22	1.42	1.38	1.43	1.45	1.38	1.51	1.63	1.72	1.71	1.66	1.62	1.83
$C_{27} \beta \alpha R/C27$	-	0.24	0.18	0.23	0.23	0.21	0.13	0.18	0.17	0.16	0.20	0.23
$C_{29} \alpha\beta\beta/(\alpha\beta\beta + \alpha\alpha\alpha)$	0.42	0.50	0.51	0.50	0.51	0.53	0.48	0.37	0.40	0.39	0.37	0.38
C <sub>29</sub> aaa20S/(20S + 20R)	0.50	0.39	0.42	0.39	0.43	0.4	0.47	0.42	0.44	0.42	0.44	0.41
TT23/(TT23 + TT24)	0.64	0.61	0.69	0.66	0.71	0.68	0.60	0.62	0.60	0.60	0.61	0.63
Ts/(Ts + Tm)	0.05	0.49	0.46	0.42	0.45	0.44	0.06	0.56	0.47	0.58	0.54	0.54
TT23/H30	0.12	0.28	0.25	0.36	0.24	0.32	0.27	0.11	0.09	0.08	0.12	0.12
H29/H30	0.73	0.70	0.67	0.66	0.69	0.62	0.93	0.53	0.55	0.55	0.54	0.51
H31 22S/(22S + 22R)	0.60	0.59	0.60	0.60	0.59	0.62	0.58	0.60	0.60	0.62	0.60	0.58
Gam/H30	-	0.23	0.25	0.27	0.28	0.27	-	0.17	0.18	0.18	0.17	0.20

Note: "-" mean undetectable, data for original samples cited from Wu et al. (2013).



Fig. 6. Modes of incorporation of hopane and sterane precursors into macromolecular network (modified from Mycke and Michaelis, 1986; Richnow et al., 1992).

Rohmer, 1987; Richnow et al., 1992; Peters et al., 2005), while steranes originate from sterols (Rohmer, 1987; Richnow et al., 1992; Peters et al., 2005). The probable precursor for tricyclic terpanes is tricyclohexaprenol, which exists in microbial membrane lipids (Ourisson et al., 1982). Thus, hydroxyl groups exist in the long branched side chains of hopanes and tricyclic terpanes precursors, while they are directly linked to six-member rings of sterol precursors (Fig. 6). For hopanes and tricyclic terpanes, their long side chains are fixed into the kerogen through sulfur or oxygen bonds and are protected by the macromolecular structure. On the other hand, the sterols are incorporated into kerogen through short bonds to their six-member A-ring, while their long side chain is exposed to the thermal stress (Fig. 6). When long branched chain cracking occurs, thermally altered hopanes or tricyclic terpanes will be released from the kerogen because they are incorporated into kerogen through long branched chain, whereas part of the thermally altered steranes may remain in kerogen because the cracking of long branched chain cannot break their linkage into kerogen. For example, as cracking products of regular steranes,  $C_{21}$  and  $C_{22}$  steranes can be released from kerogen or solid bitumen (Wu et al., 2012, 2013; Fang et al., 2014). Thus, the thermal stress has little influence on the original distributions of hopanes or tricyclic terpanes, but will change the distribution of steranes within kerogen. This may also explain why the ratio of TT23/(TT23 + TT24) is stable in bound fractions at all simulation temperatures (Fig. 5d), although TT23 and TT24 also have long side chains. Moreover, the C<sub>21</sub> and C<sub>22</sub> steranes do not have long branched chains, so the ratio S21/S22 is as stable as hopane ratios in the hydropyrolysate at all temperatures. Therefore, biomarker parameters based on bound hopanes and tricyclic terpanes are more stable than those based on bound steranes.

#### 4. Conclusions

Two kerogens (GY-8 and GY-17) isolated from different outcrop lithology of the Dalong Formation were subjected to by anhydrous pyrolysis at various temperatures. The bound hopanes and steranes within the original samples were released from the thermally altered kerogens by HyPy. The results from the anhydrous pyrolysis and HyPy experiments indicate that hopanes and tricyclic terpanes show greater thermal stability than steranes in both the free and bound fraction. Because of differences in the length of branched side chains and the mode of incorporation into kerogen, the thermal evolution of steranes is quite different from that of hopanes and tricyclic terpanes in both free and bound fractions. The source-related biomarker parameters based on hopanes and steranes have quite different thermal evolution trends in both free and bound fractions. For the investigated Permian Dalong Formation kerogen, bound regular sterane distributions are thermally altered above a simulation temperature of 430 °C (Easy% Ro = 2.27), but many hopane ratios and S21/S22 remain stable even at 460 °C. Therefore, ratios based on bound hopane and tricvclic terpane homologs seem to be very reliable source-related biomarker parameters even at 460 °C (Easy%Ro = 2.86). However, more studies are needed to confirm the maturity range in which bound hopanes and steranes might be invalid, because the results from artificial simulation experiment are not directly analogous to natural samples.

#### Acknowledgements

This work was supported by National Natural Science Foundation of China (Grant No. 41303033) and National Science and Technology Major Project (Grant No. 2011ZX05008-002). This is contribution No. IS-2284 from GIGCAS. We thank Prof. Brian Horsfield for providing insightful comments and improving this manuscript. We are grateful to Dr. Yuhong Liao, Prof. Jialan Lu, Dr. Yankuan Tian for GC–MS analysis and laboratory assistance. Dr. Ken Peters, and three anonymous reviewers are gratefully acknowledged for their constructive comments and suggestions. We thank Dr. John Volkman and Dr. Maowen Li for their great help and patience in handling the manuscript.

#### Associate Editor-Maowen Li

#### References

- Abbott, G.D., Lewis, C.A., Maxwell, J.R., 1985. Laboratory models for aromatization and isomerization of hydrocarbons in sedimentary basins. Nature 318, 651– 653.
- Abbott, G.D., Wang, G.Y., Eglinton, T.I., Home, A.K., Petch, G.S., 1990. The kinetics of sterane biological marker release and degradation processes during the hydrous pyrolysis of vitrinite kerogen. Geochimica et Cosmochimica Acta 54, 2451– 2461.
- Abbott, G.D., Bennett, B., Petch, G.S., 1995. The thermal degradation of  $5\alpha$ (H)cholestane during closed-system pyrolysis. Geochimica et Cosmochimica Acta 59, 2259–2264.
- Abbott, G.D., Bashir, F.Z., Sugden, M.A., 2001. Kerogen-bound and free hopanoic acids in the messel oil shale kerogen. Chirality 13, 510–516.
- Adam, P., Schmid, J.C., Mycke, B., Strazielle, C., Connan, J., Huc, A., Riva, A., Albrecht, P., 1993. Structural investigations of nonpolar sulfur cross-linked macromolecules in petroleum. Geochimica et Cosmochimica Acta 57, 3395– 3419.
- Beach, F., Peakman, T.M., Abbott, G.D., Sleeman, R., Maxwell, J.R., 1989. Laboratory thermal alteration of triaromatic steroid hydrocarbons. Organic Geochemistry 14, 109–111.
- Comet, P.A., McEvoy, J., Giger, W., Douglas, A.G., 1986. Hydrous and anhydrous pyrolysis of DSDP Leg 75 kerogens – a comparative study using a biological marker approach. Organic Geochemistry 9, 171–182.
- de Leeuw, J.W., Baas, M., 1986. Early diagenesis of steroids. In: Johns, R.B. (Ed.), Biological Markers in the Sedimentary Record. Elsevier, Amsterdam, pp. 102– 127.
- de Leeuw, J.W., Cox, H.C., van Graas, G., van de Meer, F.W., Peakman, T.M., Baas, J.M. A., van de Graaf, V., 1989. Limited double bond isomerization and selective hydrogenation of sterenes during early diagenesis. Geochimica et Cosmochimica Acta 53, 903–909.
- Eglinton, T.I., Douglas, A.G., 1987. Quantitative study of biomarker hydrocarbons released from kerogens during hydrous pyrolysis. Energy & Fuels 2, 81–88.
- Fang, Y.X., Liao, Y.H., Wu, L.L., Geng, A.S., 2011. Oil-source correlation for the paleoreservoir in the Majiang area and remnant reservoir in the Kaili area, South China. Journal of Asian Earth Sciences 41, 147–158.
- Fang, Y.X., Liao, Y.H., Wu, L.L., Geng, A.S., 2014. The origin of solid bitumen in the Honghuayuan Formation (O<sub>1</sub>h) of the Majiang paleo-reservoir – evidence from catalytic hydropyrolysates. Organic Geochemistry 68, 107–117.
- Farrimond, P., Taylor, A., Telnæs, N., 1998. Biomarker maturity parameters: the role of generation and thermal degradation. Organic Geochemistry 29, 1181–1197.
- Gordadze, G.N., Giruts, M.V., Koshelev, V.N., Yusupova, T.N., 2015. Distribution features of biomarker hydrocarbons in asphaltenes thermolysis products of different fractional compositions (using as an example oils from carbonate deposits of Tatarstan oilfields). Petroleum Chemistry 55, 22–31.
- Huang, D.F., Zhang, D.J., Li, J.C., 1994. The origin of 4-methyl steranes and pregnanes from Tertiary strata in the Qaidam Basin, China. Organic Geochemistry 22, 343– 348.
- Hoffman, I.C., Hutchison, J., Robson, J.N., Chicarelli, M.I., Maxwell, J.R., 1992. Evidence for sulphide links in a crude oil by lithium in ethylamine. Organic Geochemistry 19, 371–387.
- Jenisch, A., Richnow, H.H., Michaelis, W., 1990. Chemical structural units of macromolecular coal components. In: Durand, B., Behar, F. (Eds.), Advances in Organic Geochemistry 1989. Organic Geochemistry 16, pp. 917–929.
- Lewan, M.D., Bjorøy, M., Dolcater, D.L., 1986. Effects of thermal maturation on steroid hydrocarbons as determined by hydrous pyrolysis of Phosphoria Retort Shale. Geochimica et Cosmochimica Acta 50, 1977–1987.
- Lewan, M.D., 1997. Experiments on the role of water in petroleum formation. Geochimica et Cosmochimica Acta 61, 3691–3723.
- Li, H.J., Xie, X.N., Liu, Z.L., Yan, J.X., Zhou, L., Xiong, X., Su, M., 2009. Organic matter enrichment of Dalong Formation in Guangyuan area of the Sichuan Basin. Geological Science and Technology Information 28, 98–103 (in Chinese).
- Liang, D.G., Chen, J.P., 2005. Oil-source correlations for high and overmatured marine source rocks in South China. Petrology Exploration Development 32, 8–14 (in Chinese).
- Liang, M.L., Wang, Z.D., Zheng, J.J., Li, X.G., Wang, X.F., Gao, Z.D., Luo, H.Y., Li, Z.P., Qian, Y., 2015. Hydrous pyrolysis of different kerogen types of source rock at high temperature-bulk results and biomarkers. Journal of Petroleum Science and Engineering 125, 209–217.

- Liao, Y.H., Fang, Y.X., Wu, L.L., Geng, A.S., Hsu, C.S., 2012. The characteristics of the biomarkers and  $\delta^{13}$ C of *n*-alkanes released from thermally altered solid bitumens at various maturities by catalytic hydropyrolysis. Organic Geochemistry 46, 56–65.
- Liao, Y.H., Fang, Y.X., Wu, L.L., Cao, Q.G., Geng, A.S., 2015. The source of highly overmature solid bitumens in the Permian coral reef paleo-reservoirs of the Nanpanjiang Depression. Marine and Petroleum Geology 59, 527–534.
- Liao, Z.W., Geng, A.S., 2002. Characterization of *n*C<sub>7</sub>-soluble fractions of the products from mild oxidation of asphaltenes. Organic Geochemistry 33, 1477–1486.
- Lockhart, R.S., Meredith, W., Love, G.D., Snape, C.E., 2008. Release of bound aliphatic biomarker via hydropyrolysis from Type II kerogen at high maturity. Organic Geochemistry 39, 1119–1124.
- Love, G.D., Snape, C.E., Carr, A.D., Houghton, R., 1995. Release of bound alkane biomarkers in high yields from kerogen via catalytic hydropyrolysis. Organic Geochemistry 23, 981–986.
- Love, G.D., McAulay, A., Snape, C.E., Bishop, A.N., 1997. Effect of process variable in catalytic hydropyrolysis on the release of bound aliphatic hydrocarbons from sedimentary organic matter. Energy & Fuels 11, 522–531.
- Love, G.D., Snape, C.E., Fallick, A.E., 1998. Differences in the mode of incorporation and biogenicity of the principal aliphatic constituents of a Type I oil shale. Organic Geochemistry 28, 797–811.
- Love, G.D., Bowden, S.A., Jahnke, L.L., Snape, C.E., Campbell, C.N., Day, J.G., Summons, R.E., 2005. A catalytic hydropyrolysis method for the rapid screening of microbial cultures for lipid biomarkers. Organic Geochemistry 36, 63–82.
- Lu, S.T., Ruth, E., Kaplan, I.R., 1989. Pyrolysis of kerogens in the absence and presence of montmorillonite – I. The generation, degradation and isomerization of steranes and triterpanes at 200 and 300 °C. Organic Geochemistry 14, 491– 499.
- Mackenzie, A.S., Brassell, S.C., Eglinton, G., Maxwell, J.R., 1982a. Chemical fossils: the geological fate of steroids. Science 217, 491–504.
- Mackenzie, A.S., Lamb, N.A., Maxwell, J.R., 1982b. Steroid hydrocarbons and the thermal history of sediments. Nature 295, 223–226.
- Meredith, W., Russell, C.A., Cooper, M., Snape, C.E., Love, G.D., Fabbri, D., Vane, C.H., 2004. Trapping hydropyrolysates on silica and their subsequent thermal desorption to facilitate rapid fingerprinting by GC–MS. Organic Geochemistry 35, 73–89.
- Meredith, W., Snape, C.E., Carr, A.D., Nytoft, H.P., Love, G.D., 2008. The occurrence of unusual hopanes in hydropyrolysates generated from severely biodegraded oil seep asphaltenes. Organic Geochemistry 39, 1243–1248.
- Michaelis, W., Jenisch, A., Richnow, H.H., Kruse, U., Mycke, B., 1988. Organofacies des Ölschiefers von Messel. Courier Forschung Institut Senckenberg 107, 89– 103.
- Muhammad, A.B., Abbott, G.D., 2012. The thermal evolution of asphaltene-bound biomarkers from coals of different rank: a potential information resource during coal biodegradation. International Journal of Coal Geology 107, 90–95.
- Murray, I.P., Love, G.D., Snape, C.E., Bailey, N.J.L., 1998. Comparison of bound aliphatic biomarkers released via hydropyrolysis with their solvent-extractable counterparts for a suite of Kimmeridge clays. Organic Geochemistry 29, 1487– 1505.
- Mycke, B., Michaelis, W., 1986. Molecular fossils from chemical degradation of macromolecular organic matter. Organic Geochemistry 10, 847–858.
- Mycke, B., Narjes, F., Michaelis, W., 1987. Bacteriohopanetetrol from chemical degradation of an oil shale kerogen. Nature 326, 179–181.
- Ourisson, G., Albrecht, P., Rohmer, M., 1982. The hopanoids, paleochemistry and biochemistry of a group of natural products. Trends in Biochemical Sciences 7, 236–239.
- Peters, K.E., Moldowan, J.M., Driscole, A.R., Demaison, G.J., 1989. Origin of Beatrice oil by co-sourcing from Devonian and Middle Jurassic source rocks, Inner Moray Firth, U.K. American Association of Petroleum Geologists Bulletin 73, 454–471.
- Peters, K.E., Walters, C.C., Moldowan, J.M., 2005. The Biomarker Guide, Biomarkers and Isotopes in Petroleum Exploration and Earth History. Cambridge University Press, New York, 699 pp.
- Richnow, H.H., Jenisch, A., Michaelis, W., 1992. Structural investigations of sulphurrich macromolecular oil fractions and a kerogen by sequential chemical degradation. Organic Geochemistry 19, 351–370.
- Rohmer, M., Bouvier-Nave, P., Ourisson, G., 1984. Distribution of hopanoid triterpenes in prokaryotes. Journal of General Microbiology 130, 1137–1150.
- Rohmer, M., 1987. The hopanoids, prokaryotic triterpenoids and sterol surrogates. In: Schrinner, E., Richmond, M.H., Seibert, G., Schwarz, U. (Eds.), Surface Structure of Microorganisms and Their Interactions with the Mammalian Host. WCH Publishing, Weinlein, Germany, pp. 227–242.
- Russell, C.A., Snape, C.E., Meredith, W., Love, G.D., Clarke, E., Moffatt, B., 2004. The potential of bound biomarker profiles released via catalytic hydropyrolysis to reconstruct basin charging history for oils. Organic Geochemistry 35, 1441– 1459.
- Sweeney, J.J., Burnham, A.K., 1990. Evaluation of a simple method of vitrinite reflectance based on chemical kinetics. American Association of Petroleum Geologists Bulletin 74, 1559–1570.
- Seifert, W.K., 1978. Steranes and terpanes in kerogen pyrolysis for correlation of oils and source rocks. Geochimica et Cosmochimica Acta 42, 473–484.
- Seifert, W.K., Moldowan, J.M., Demaison, G.J., 1984. Source correlation of biodegraded oils. Organic Geochemistry 6, 633–643.
- Sinninghe Damsté, J.S., de Leeuw, J.W., 1990. Analysis, structure and geochemical significance of organically-bound sulphur in the geosphere: state of the art and future research. Organic Geochemistry 16, 1077–1101.

- Snape, C.E., Bolton, C., Dosch, R.G., Stephens, H.P., 1989. High liquid yields from bituminous coal via hydropyrolysis with dispersed catalysts. Energy & Fuels 3, 421–425.
- Snowdon, L.R., Volkman, J.K., Zhang, Z.R., Tao, G.L., Liu, P., 2016. The organic geochemistry of asphaltenes and occluded biomarkers. Organic Geochemistry 91, 3–15.
- Stalker, L., Larter, S.R., Farrimond, P., 1998. Biomarker binding into kerogens: evidence from hydrous pyrolysis using heavy water (D<sub>2</sub>O). Organic Geochemistry 28, 239–253.
- Strausz, O.P., Mojelsky, T.W., Faraji, F., Lown, E.M., Peng, P., 1999a. Additional structural details on Athabasca asphaltenes and their ramifications. Energy & Fuels 13, 207–227.
- Strausz, O.P., Mojelsky, T.W., Lown, E.M., Kowalewski, I., Behar, F., 1999b. Structural features of Boscan and Duri asphaltenes. Energy & Fuels 13, 228–247.
- Sugden, M.A., Abbott, G.D., 2002. The stereochemistry of bound and extractable pentacyclic triterpenoids during closed system pyrolysis. Organic Geochemistry 33, 1515–1521.
- Tenger Qin, J.Z., Fu, X.D., Li, W., Rao, D., Zhang, M.Z., 2008. Basic conditions of marine hydrocarbon accumulation in northwest Sichuan Basin-High quality source rocks. Petroleum Geology & Experiment 30, 478–483 (in Chinese).
- Tissot, B.P., Welte, D.H., 1978. Petroleum Formation and Occurrence: A New Approach to Oil and Gas Exploration. Springer-Verlag, 538 pp.

- van Graas, G.W., 1990. Biomarker maturity parameters for high maturities for high maturities: calibration of the working range up to the oil/condensate threshold. Organic Geochemistry 16, 1025–1032.
- Vu, T.A.T., Horsfield, B., Sykes, R., 2008. Influence of in-situ bitumen on the generation of gas and oil in New Zealand coals. Organic Geochemistry 39, 1606– 1619.
- Wang, Y.G., Wen, Y.C., Hong, H.T., Xia, M.L., Zhan, J., Song, S.Y., Liu, H.Y., 2006. Petroleum geological characteristics of deep water deposits in Upper Permian-Lower Triassic trough in Sichuan Basin and adjacent areas. Oil and Gas Geology 27, 702–714 (in Chinese).
- Wu, L.L., Liao, Y.H., Fang, Y.X., Geng, A.S., 2012. The study on the source of the oil seeps and bitumens in the Tianjingshan structure of the northern Longmen Mountain structure of Sichuan Basin, China. Marine and Petroleum Geology 37, 147–161.
- Wu, L.L., Liao, Y.H., Fang, Y.X., Geng, A.S., 2013. The deference in biomarkers released by hydropyrolysis and by Soxhlet extract from source rocks of different maturities and its geological implications. Chinese Science Bulletin 58, 373– 383.
- Wu, L.L., Liao, Y.H., Geng, A.S., 2015. Investigation of hydropyrolysis released aromatic hydrocarbons from Permian kerogens at different maturities in the Sichuan Basin, China. Journal of Analytical and Applied Pyrolysis 114, 47–59.