西藏冈底斯南部陆陆碰撞早期成矿作用分析。

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Abstract Many porphyry Cu-Mo and Cu-Pb-Zn polymetal deposits have been found in the southern Gangdese belt, forming a porphyry copper ore belt and a polymetal ore belt, respectively. Previous geochronologic results indicate that most deposits in the southern Gangdese belt were formed after 30Ma and in post-collisional extension evnironment. In this paper, we analyze zircon U-Pb age of igneous rocks closely associate with the Jialong skarn Fe deposit and the Sadang Au-Ag deposit, respectively, and molybdenite Re-Os age of the Duodigou skarn Mo deposit in the southern Gangdese. The Jialong biotite monzonite granite closely associated with skarn Fe deposit has zircon LA-ICP-MS U-Pb age of 61. 1 \pm 0. 4Ma with MSWD = 0. 94. The Sadang andesite closely associated with Au-Ag mineralization has zircon LA-ICP-MS U-Pb age of 62. 6 \pm 0. 5Ma with MSWD = 1. 51. The molybdenite Re-Os model ages of the Duodigou skan Mo deposit varies from 64. 3 \pm 0. 8Ma to 69. 2 \pm 3. 3Ma, with an average weight age of 66. 7 \pm 6. 4Ma, MSWD = 8. 1. Our new results suggest that the three deposits own their origin to the magmatic events triggered by early Indus-Eurasia collision. Our results, together with former work, reveal that the southern Gangdese has undergone large-scale mineralization associated with the strong magmatic activities during the early Indus-Eurasia collision. It is, therefore, suggested that the southern Gangdese could be an important target area for mineralization related to early Indus-Eurasia collision.

Key words Gangdese; Continental collision; Ore forming age; Tibet

摘 要 冈底斯带南部发育有大量的斑岩铜钼矿床和矽卡岩型铜铅锌多金属矿床,形成了斑岩铜矿带和多金属矿带。前人的研究表明,成矿带内的矿床形成年代大都小于30Ma,处于碰撞后期伸展构造环境。本文对冈底斯带中南部的甲龙矽卡岩型铁矿、撒当金银矿床(点)和多底沟矽卡岩型钼矿床(点)开展了年代学研究,结果显示:甲龙铁矿黑云母二长花岗斑岩的锆石 LA-ICP-MS U-Pb 年龄为61.1±0.4Ma,MSWD=0.94;撒当赋矿安山岩锆石 LA-ICP-MS U-Pb 年龄为62.6±0.5Ma,MSWD=1.51;多底沟钼矿床(点)3件辉钼矿 Re-Os 模式年龄为64.3±0.8Ma~69.2±3.3Ma,加权平均模式年龄为66.7±6.4Ma(MSWD=8.1)。三个矿床(点)的同位素年龄表明成岩成矿事件和印度-欧亚板块陆陆碰撞早期构造岩浆事件有关。结合前人工作,我们提出冈底斯中南部发生了大规模与陆陆碰撞早期岩浆事件有关的成矿作用,形成了大面积分布的矿床,具有良

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关键词 冈底斯带;陆陆碰撞;成矿时代;西藏

中图法分类号 P597.3; P611

1 引言

近年来在冈底斯南部发现一系列斑岩铜钼矿床及多金 属矿床,形成了巨型冈底斯金属成矿带,引起了人们的广泛 关注。目前工作表明,冈底斯成矿带矿床主要形成于后碰撞 构造环境,成矿时代集中在 18~13Ma 之间 (Hou et al., 2009; Yang et al., 2009; 侯增谦等, 2003, 2004, 2012; 芮宗 瑶等, 2003; 孟祥金等, 2003; 李光明和芮宗瑶, 2004; 林武 等, 2004; 郑有业等, 2007; 杨志明等, 2008; 梁华英等, 2008; 唐菊兴等, 2011),主要以大型超大型斑岩铜钼矿床为 主,如驱龙、甲马、冲江、厅宫、邦浦等;部分矿床形成于晚碰 撞期,成岩成矿时代在 30~24Ma 之间(Chung et al., 2009; 李光明等, 2006; 莫济海等, 2006; 闫学义等, 2010; 范新 等, 2011; 张松等, 2012), 主要分布于冈底斯带南缘克鲁-冲 木达矿集区,主要矿床有冲木达矽卡岩铜钼矿床(莫济海等, 2008)、努日矽卡岩型钼铜钨矿床(李光明等, 2011; 陈雷等, 2011)、明则斑岩钼矿床(陈玉水等, 2011)等。许多学者对 后碰撞期斑岩矿床的时空分布及成矿岩体地球化学特征开 展了工作,认为成矿斑岩具有埃达克岩亲和性(侯增谦等, 2012);而对成矿斑岩形成动力学过程则有不同的看法,主要 有:俯冲的新特提斯洋壳板片熔融(Qu et al., 2004)、俯冲的 印度下地壳熔融(Xu et al., 2010)、加厚的新生镁铁质下地 壳熔融(Hou et al., 2004; Chung et al., 2009)及俯冲板片物 质改造的岩石圈地幔熔融(Gao et al., 2007, 2010)等观点。

印度板块与欧亚板块碰撞引发了大规模的岩浆活动,形成了大面积分布的林子宗火山岩及一系列侵入岩(Chung et al., 2005; 莫宣学等, 2003, 2005, 2007; 周肃等, 2004; 潘桂棠等, 2006),强烈岩浆活动为矿床的形成提供了良好的条件。初步工作表明,冈底斯中南部在印度板块与欧亚板块碰撞早期发生了与中酸性岩浆作用有关成矿事件(李应栩等, 2011; 高一鸣等, 2011; 高顺宝等, 2012; 王保弟等, 2012),但目前对该时期成矿的时空分布工作不多。

本文分析了林周县甲龙铁矿和撒当火山岩型金银矿床 (点)矿化岩体锆石 LA-ICP-MS U-Pb 年龄,以及多底沟矽卡 岩型钼矿床(点)辉钼矿 Re-Os 模式年龄,发现这些成矿岩体 或矿床均形成于陆陆碰撞早期。结合前人相关成果,我们初 步分析了印度-亚欧板块碰撞早期与中酸性岩浆作用有关矿 床的时空分布及成矿作用。

2 区域地质概况及矿床地质简介

冈底斯带处于雅鲁藏布江缝合带与怒江-班公湖缝合带

之间,是一条东西长达 2500km,南北宽 150~300km,面积达 45 万平方千米的巨型构造-岩浆带(潘桂棠等,2006)。冈底 斯带岩浆岩主要形成于古新世-始新世,少量形成于中新世和侏罗纪(Chung et al.,2003; Chu et al.,2006)。古新世-始新世火成岩主要有覆盖于上白垩统设兴组的林子宗火山岩及冈底斯花岗岩基。林子宗火山岩主要为钙碱性-钾玄岩系列的中酸性火山、火山碎屑岩,形成时代在 65~41Ma 之间(莫宣学等,2003);冈底斯花岗岩岩基主要分布于冈底斯带中南部,东西向 2000 多千米范围内均见出露,形成时代在 53~47Ma 之间(Mo et al.,2005)。中新世火成岩主要有后碰撞钾质-超钾质火山岩及与斑岩铜钼矿床有关的钾质二长花岗斑岩。

本研究的矿床位于冈底斯带中南部,斑岩铜钼矿带北缘,包括甲龙铁矿床赋矿岩体、撒当金银矿点赋矿火山岩及多底沟钼矿床(点)赋矿岩体(图1)。甲龙铁矿床位于林周县境内,目前探明储量约为100Mt,达小型规模。矿区出露地层主要有:白垩系上统设兴组砂岩粉砂岩和塔克那组砂岩、泥岩、灰岩,及白垩系下统楚木龙组石英砂岩。矿区内岩浆岩主要为侵入白垩系地层中的黑云母二长花岗岩及斑状黑云母二长花岗岩(西藏自治区地质矿产局,1993)。甲龙砂卡岩型铁矿产于斑状黑云母二长花岗岩与楚木龙组石英砂岩的接触带内,矿体为透镜状产于内接触带,金属矿物主要为磁铁矿,矿石具粒状结构、块状构造。斑状黑云母二长花岗岩主要由斜长石、钾长石、石英和黑云母组成。野外观察显示,斑状黑云母二长花岗岩与地层的内接触带发育石榴子石及透辉石等早砂卡岩阶段形成的矿物,远离岩体则阳起石及绿帘石等相对较多。

撒当金银矿床(点)位于林周县境内,拉萨北部约70km处。矿区出露的地层主要有第四系沉积物、始新统帕那组凝灰岩、年波组凝灰岩夹砂砾岩及古新统典中组安山岩。矿化主要呈脉状产于典中组安山岩中,目前已发现3条长100多米、宽数十厘米的矿脉,主要金属矿物有方铅矿、闪锌矿、黄铜矿、黄铁矿、自然金及金银矿等。赋矿火山岩发生了强烈的蚀变,主要蚀变类型有硅化、黄铁矿化、绢云母化、碳酸盐化及沸石化。目前该矿化点为个体企业开采,探矿工程不多,矿床规模不明。

多底沟砂卡岩型钼矿床(点)位于拉萨东部十多千米处。矿区的地层主要有第四系沉积物、白垩系叶巴组中酸性火山岩及侏罗系上统多底沟组灰岩夹粉砂岩;区内第三系岩浆活动较为活跃,出露第三系黑云母二长花岗岩及黑云母花岗闪长岩(西藏自治区地质矿产局,1993)。多底沟矿床钼矿化主要产于角闪黑云母二长花岗岩与侏罗系多底沟组灰岩的接触带中。角闪黑云母二长花岗岩主要由斜长石、钾长石、石

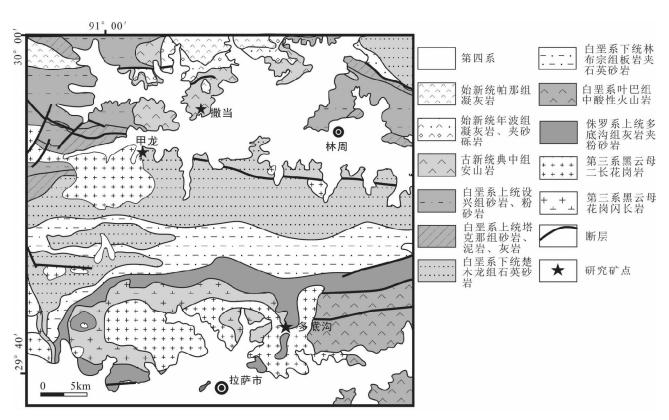


图 1 研究矿点区域地质简图(据西藏自治区地质矿产局,1993修改)

Fig. 1 Simplified geological map of the study deposits (modified after BGMRX, 1993)

英、黑云母和角闪石组成,岩体与围岩接触带发育矽卡岩化,钼矿化主要呈细脉状及浸染状产于矽卡岩中,局部见厚度超过10cm 辉钼矿石英脉。

3 样品和分析方法

本文对甲龙铁矿斑状黑云母二长花岗岩和撒当金银矿床(点)赋矿安山岩进行了锆石 LA-ICP-MS U-Pb 年龄分析,对多底沟钼矿床(点)的辉钼矿进行了 Re-Os 年龄测定。

甲龙铁矿斑状黑云母二长花岗岩的采样位置为N29°52′57″,E91°02′30″;撒当金银矿床(点)赋矿安山岩样品采于N29°55′44.3″,E91°06′24.4″。每个样品选取约0.5kg的岩石,研磨至80目,经筛选、淘选、磁选等处理后,在显微镜下手选出晶形完整的锆石70~80颗,装人环氧树脂中制成靶、磨光,在光学显微镜下选出无裂纹且包裹体不发育的锆石晶体点进行测定。锆石 LA-ICP-MS 定年在澳大利亚国立大学地球科学研究所完成。分析流程见 Harris et al. (2004)。为了减少继承铅、铅丢失等对年龄的影响,在207 Pb/235 U-206 Pb/238 U图中,和谐度低于95%的年龄数据及观测误差值/预期误差值大于2的分析数据点将被排除。在此基础上,使用累积概率统计图对数据进行处理。主群组锆石年龄在累积概率统计图中多沿直线分布,其年龄代表岩体结晶年龄,位于直线上方的被认为是继承铅,位于直线下方

的被认为是铅丢失 (Harris et al., 2004; Liang et al., 2006)。

多底沟钼矿床(点)的三件辉钼矿样品均取自矿区的钼矿石,采样位置为 N29°42′42.6″, E91°11′59.2″。选取 1kg 左右的样品,用刀片直接从手标本上剥离辉钼矿后,在双目镜下进一步挑选至纯度 98%以上。辉钼矿同位素分析在中国科学院广州地球化学研究所完成。采用 Carius 管封闭溶样分解样品,样品分解以及 Re 和 Os 的分离等化学处理过程参见文献(Sun et al., 2010; Wu et al., 2012),质谱分析仪器是美国热电公司生产的 X7型 ICPMS,采用 Ir 和 Os 天然丰度进行在线监测和校正仪器测试过程中的 Re 和 Os 同位素分馏。¹⁸⁷ Re-¹⁸⁷ Os 衰变常数 λ 引用的值为 1.666 × 10⁻¹¹ a⁻¹ (Shen et al., 1996; Smoliar et al., 1996)。

4 分析结果

与甲龙砂卡岩型铁矿紧密共生的斑状黑云母二长花岗岩和撒当金银矿床(点)矿化安山岩的锆石 LA-ICP-MS U-Pb 年龄的分析结果见表 1 及表 2。

本文对甲龙铁矿蚀变斑状黑云母二长花岗岩进行了 30 颗锆石 LA-ICP-MS U-Pb 年龄分析,所有分析点谐和度均高于 95%。30个分析点锆石 U-Pb 年龄在 64.9~57.5 Ma 之间,用 isoplot 计算加权平均年龄为 60.8 ± 0.7 Ma, MSWD =

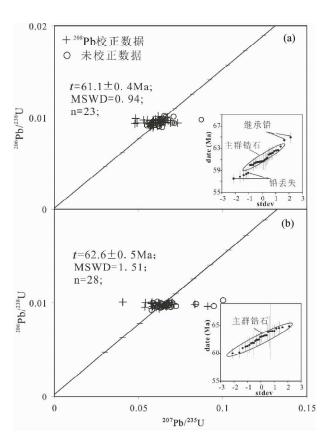


图 2 甲龙斑状黑云母二长花岗岩体(a) 和撒当安山岩(b) 锆石 U-Pb 年龄谐和图, 内插为累计概率统计图 Fig. 2 Condordia plots showing the zircon U-Pb analyses of the Jialong porphyritic biotite monzonite granite (a) and the Sadang andesite (b), the inserts are probability plots

3.01。其 MSWD 值较大,显示其含有继承铅或铅丢失。为了获得更精确的年龄值,我们用累积概率统计图分析锆石年龄分布特征。在累积概率统计图上,23 个数据点沿直线分布,为主群锆石,而位于直线上方的2 个数据和位于直线下方的5 个数据分别看作继承铅和铅丢失(图 2a),均不参与计算。用 isoplot 软件计算 23 个点主群锆石年龄为61.1 ±0.4Ma, MSWD=0.94(图 2a)。

撒当金银矿床(点)赋矿安山岩31个锆石 LA-ICP-MS U-Pb 年龄分析点中,有3个谐和度小于95%,在计算年龄时被排除,其余28个分析点锆石 U-Pb 年龄变化范围为64.8~60.0Ma,在累积概率统计图上沿直线分布。用 isoplot 进行加权平均年龄计算结果为62.6±0.5Ma, MSWD=1.51(图2b)。

多底沟钼矿床(点)3件辉钼矿同位素分析结果详见表3。结果显示,辉钼矿¹⁸⁷ Re 的含量为 $0.878873 \times 10^{-6} \sim 1.50942 \times 10^{-6}; ^{187}$ Os 的含量为 $1.00405 \times 10^{-9} \sim 1.61907 \times 10^{-9}$,模式年龄为 68.5 ± 0.7 Ma、 69.2 ± 3.3 Ma 和 64.3 ± 0.8 Ma,加权平均模式年龄为 66.7 ± 6.4 Ma (MSWD = 8.1)。

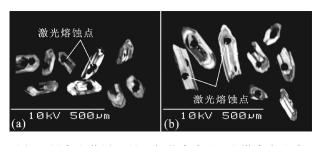


图 3 甲龙斑状黑云母二长花岗岩(a)和撒当安山岩(b)锆石样品阴极发光图

Fig. 3 CL photos of analyzed zircon grains from the Jialong porphyritic biotite monzonite granite (a) and from the Sadang andesite (b)

5 讨论

5.1 矿床的形成时代

甲龙铁矿蚀变斑状黑云母二长花岗岩及撒当金银矿床(点)赋矿安山岩锆石样品均为长柱状,晶形较好,Th/U 比值较大,在 0.41~1.18之间(表 1、表 2)。锆石 CL 图韵律环带发育(图 3),具有岩浆锆石的一般特征,说明其主要是在岩浆结晶过程中形成的。锆石的 U-Pb 同位素体系封闭温度较高,抗后期干扰能力强,不易受后期热事件的扰动,因此获得的主群锆石 LA-ICP-MS U-Pb 年龄可以代表岩体的结晶年龄。甲龙铁矿蚀变斑状黑云母二长花岗岩形成时代为 61.1±0.4Ma;撒当金银矿床(点)赋矿安山岩形成时代为 62.6±0.5Ma。

甲龙铁矿的测试样品采自与矿化关系密切的蚀变斑状黑云母二长花岗岩。野外观察显示矿体主要赋存于岩体与地层的接触带中,岩体内接触带高温砂卡岩矿物如石榴子石及透辉石等相对发育,向外则阳起石、绿帘石及绿泥石等矿物增多。这表明砂卡岩化与斑状黑云母二长花岗岩有关,斑状黑云母二长花岗岩锆石 LA-ICP-MS U-Pb 年龄反映砂卡岩铁矿形成时代,即甲龙砂卡岩型铁矿是在 61.1 ± 0.4 Ma 形成的。

撒当金银矿床(点)矿体主要呈脉状及细脉状产于安山岩中。安山岩发生了强烈的黄铁矿化、硅化、绢云母化及沸石化等蚀变。目前未有直接证据表明矿化与安山岩有内在成因联系,考虑到该类矿化主要产于火山岩中,而且火山岩中还发育较大面积的 Au-Ag 异常,冈底斯南缘沉积地层中未见同类元素组合矿化;加上产于林子宗火山岩中纳如松多银矿床热液蚀变绢云母 Ar-Ar 年龄(57.8Ma)和火山岩锆石年龄相近(杨勇等,2010),显示火山活动发生了成矿作用。因此,我们初步认为火山岩中的金银矿化与火山岩有关,金银矿化时代应和赋矿火山岩时代相近,略晚于62.6±0.5Ma。

多底沟砂卡岩型钼矿床(点)三件辉钼矿样品 Re-Os 模式年龄分别为: 68.5 ± 0.7 Ma、69.2 ± 3.3 Ma 及 64.3 ± 0.8 Ma。三件辉钼矿化样品模式年龄变化不大,加权平均模

表 1 甲龙矽卡岩型铁矿斑状黑云母二长花岗岩锆石 LA-ICP-MS 分析结果

Table 1 Zircon LA-ICP-MS data of the Jialong porphyritic biotite monzonite granite associated with Fe mineralization

测点号	U	Th	²⁰⁶ Pb	± 1s. e.	²⁰⁷ Pb	± 1s. e.	²⁰⁸ Pb	± 1s. e.	$\frac{^{206}\mathrm{Pb}}{^{238}\mathrm{U}}$	±1s. e.	$\frac{^{207}\mathrm{Pb}}{^{235}\mathrm{LI}}$	±1s. e.
1907 77	$(\times 10^{-6})$	U	²³⁸ U	± 15. C.	²³⁵ U	± 13. C.	$\overline{^{232}}$ Th	± 13. C.	age (Ma)	± 13. C.	age (Ma)	113. 0.
42-01	658	0. 54	0. 00897	0. 00013	0. 06439	0. 00440	0. 00301	0. 00009	57. 5	0.8	57.7	3. 6
42-02	363	0.67	0.00998	0.00008	0.06613	0.00190	0.00314	0.00007	64. 4	1.2	66. 2	5.6
42-03	1742	0.46	0.00976	0.00006	0.06324	0.00103	0.00326	0.00004	62. 6	1.0	58.9	4. 7
42-04	411	0.56	0.00941	0.00013	0.05601	0.00386	0.00316	0.00007	60. 2	0.9	48. 9	2. 5
42-05	2563	1.06	0.00930	0.00006	0.05895	0.00106	0.00298	0.00002	60.0	1.2	56. 1	4. 7
42-06	955	0.63	0.00931	0.00008	0.06184	0.00143	0.00299	0.00003	59. 9	0.9	59.6	3. 2
42-07	321	0.69	0.01002	0.00026	0.07100	0.00525	0.00373	0.00012	63. 3	1.0	47.7	2. 7
42-08	438	0.54	0.00944	0.00012	0.06390	0.00334	0.00294	0.00008	60. 9	1.8	64. 7	8. 7
42-09	303	0.56	0.00942	0.00011	0.06484	0.00354	0.00302	0.00006	60.6	1.1	62. 2	4. 3
42-10	270	1. 13	0.00939	0.00012	0.05926	0.00268	0.00305	0.00005	60. 5	1.1	52.7	4. 1
42-11	324	0. 57	0.00939	0.00011	0.05670	0.00509	0.00323	0.00011	60.0	1.2	47.5	8. 2
42-12	333	0.56	0.00935	0.00014	0.05787	0.00488	0.00283	0.00008	60. 4	1.1	61.3	6. 9
42-13	262	0.67	0.00949	0.00013	0.06499	0.00396	0.00312	0.00009	61.0	1.2	59.4	5. 7
42-14	253	0.49	0.00963	0.00011	0.06311	0.00343	0.00307	0.00011	62.0	1.2	61.4	5. 5
42-15	296	0.62	0.00900	0.00011	0.05736	0.00379	0.00281	0.00012	58. 2	1.1	58.3	4. 6
42-16	225	0.56	0.00962	0.00020	0.06805	0.00668	0.00304	0.00011	62. 0	1.1	66. 5	5.8
42-17	376	0.63	0.00940	0.00006	0.05802	0.00200	0.00298	0.00005	60.6	1.5	57.4	8. 3
42-18	371	0. 57	0.00935	0.00013	0.07137	0.00720	0.00292	0.00009	60. 3	0.9	71.6	3.6
42-19	263	0. 57	0.00909	0.00009	0.06202	0.00330	0.00290	0.00009	58. 5	1.2	60. 2	7. 4
42-20	230	0.52	0.00908	0.00012	0.05530	0.00306	0.00283	0.00009	58. 3	1.0	56. 4	4. 8
42-21	377	0.63	0.00902	0.00012	0.06235	0.00345	0.00299	0.00007	57.8	1.1	56.0	4. 6
42-22	195	0.44	0.00940	0.00016	0.07018	0.00550	0.00300	0.00012	60. 4	1.1	67.7	4. 6
42-23	209	0.48	0.00972	0.00013	0.06262	0.00363	0.00334	0.00010	62. 2	1.3	54. 4	6. 3
42-24	475	0.71	0.00977	0.00012	0.06414	0.00329	0.00334	0.00008	62. 4	1.2	52.6	4. 4
42-25	260	0.65	0.00964	0.00008	0.06610	0.00278	0.00322	0.00008	61.8	1.1	58. 2	5. 0
42-26	250	0.52	0.00942	0.00010	0.06020	0.00294	0.00305	0.00007	60. 5	1.0	57.0	4. 4
42-27	740	0.72	0.00939	0.00008	0.06319	0.00170	0.00299	0.00005	60.6	1.0	61.2	3. 9
42-28	262	0.53	0.01010	0.00011	0.06540	0.00284	0.00331	0.00008	64. 9	1.0	60.7	3.8
42-29	233	0.48	0.00977	0.00011	0.06475	0.00307	0.00329	0.00008	62. 6	1.1	57.8	3. 7
42-30	223	0.56	0. 00969	0.00010	0. 08747	0.00493	0.00384	0.00014	61. 2	1.1	62. 2	3.8

式年龄为66.7±6.4Ma,MSWD=8.1。辉钼矿模式年龄基本可代表多底沟矽卡岩型钼矿床(点)的形成时代,多底沟矽卡岩型钼矿床(点)是在66.7±6.4Ma左右形成的。

5.2 冈底斯带南缘陆陆碰撞早期成矿作用讨论

冈底斯带经历了复杂地质构造演化历史,在中新生代先后受雅鲁藏布江洋壳向南俯冲、洋壳消亡和期后的弧-陆碰撞及陆-陆碰撞作用,在冈底斯带南缘形成了与洋壳俯冲及陆陆碰撞有关的火山岩和侵入岩(潘桂棠等,2002; Ding et al.,2003,2005; Hou et al.,2004,2006,2007; Mo et al.,2005,2007; Zhu et al.,2008; Ji et al.,2009)。虽经多年工作,但目前对新特提斯洋最终消亡乃致印度板块与欧亚板块发生碰撞的时代仍有不同看法(Rowley,1998; Yin and Harrison,2000; Aitchison et al.,2007)。Lee and Lawver(1995)提出,印度板块与欧亚板块在约65Ma时发生"软碰撞",而在约45Ma时发生"硬碰撞"。莫宣学等(2003,

2009)、Mo et al. (2007)提出两大陆在 70~65Ma 左右开始碰撞,经过约二十多个百万年,在 40Ma 左右完成两个大陆碰撞,新特提斯洋完全消失;侯增谦等(2006a, b, c, d)把西藏碰撞分为主碰撞(65~41Ma)、晚碰撞(40~26Ma)和后碰撞(25~0Ma)三期。总的来说,目前较多地质、地球物理及地球化学证据支持印度板块与欧亚板块在 70~65Ma 左右时发生碰撞。

甲龙铁矿床赋矿岩体锆石 LA-ICP-MS U-Pb 年龄 61.1±0.4Ma、撒当赋金银矿(点)火山岩锆石 LA-ICP-MS U-Pb 年龄 62.6±0.5Ma、多底沟钼矿(点)辉钼矿 Re-Os 年龄在 66.7±6.4Ma 左右。三个矿床成岩成矿时代和印度板块与欧亚板块开始碰撞时代(70~65Ma)相近,因此,这些矿床应是印度板块与欧亚板块碰撞早期形成的,是碰撞早期构造岩浆事件的产物。已有的工作表明查个勒铜铅锌矿床(高顺宝等,2012; 王保弟等,2012)、亚贵拉铅锌钼多金属矿床(高一鸣等,2011; 黄克贤等,2012)及恰功砂卡岩型铁矿床(李应栩

表 2 撒当金银矿床安山锆石 LA-ICP-MS 分析结果

Table 2 Zircon LA-ICP-MS data of the Sadang andesite associated with Au-Ag mineralization

测点号	U (×10 ⁻⁶)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\mathrm{Pb}}{^{238}\mathrm{U}}$	±1s. e.	$\frac{^{207}{\rm Pb}}{^{235}{\rm U}}$	±1s. e.	$\frac{^{208}\mathrm{Pb}}{^{232}\mathrm{Th}}$	±1s. e.	206 Pb 238 U age (Ma)	±1s. e.	$\frac{^{207} \text{Pb}}{^{235} \text{U}}$ age (Ma)	±1s. e.	备注
44-01	36	1. 17	0. 00888	0. 00038	0. 17327	0. 02230	0. 00320	0. 00023	55. 7	0.8	131. 3	3. 9	剔除
44-02	32	1. 18	0.00958	0.00040	0.09507	0. 01190	0.00303	0.00021	61.8	2.7	88.7	25.7	
44-03	28	0.96	0.01023	0.00036	0. 10061	0.01332	0.00377	0.00036	63. 9	2.9	56. 9	20. 1	
44-04	545	0.58	0.00959	0.00008	0.06471	0.00209	0.00305	0.00005	61.8	2.9	63. 1	30. 1	
44-05	126	0. 20	0. 10803	0.00086	0. 95682	0. 02688	0. 03671	0.00094	660. 2	8.6	660. 1	9. 1	剔除
44-06	334	0.60	0.00996	0.00008	0.05894	0.00195	0.00331	0.00006	63. 9	5. 1	53.7	15. 4	
44-07	192	0.41	0.00953	0.00012	0.06959	0.00426	0.00315	0.00012	61. 2	1.0	65.4	6. 7	
44-08	152	1.06	0.00994	0.00017	0.06705	0.00553	0.00319	0.00009	64. 0	1.2	60. 9	5.7	
44-09	245	0.58	0.00958	0.00008	0.06313	0.00269	0.00317	0.00007	61.5	1.4	57. 6	8. 9	
44-10	785	1. 33	0.00918	0.00007	0.05813	0.00166	0.00288	0.00003	60.0	0.9	61.5	3.6	
44-11	256	0.55	0.01200	0.00024	0. 28108	0.01995	0. 01099	0.00080	64. 8	1.0	40.7	6. 9	
44-12	434	0.73	0.00964	0.00020	0.07251	0.00797	0.00345	0.00007	61. 2	1.9	53. 2	28. 0	
44-13	341	0.52	0.00977	0.00010	0.06058	0.00290	0.00309	0.00008	62. 9	1.5	59. 4	8. 9	
44-14	302	0.47	0.01003	0.00011	0.06972	0.00301	0.00324	0.00008	64. 5	1.1	66. 5	4. 0	
44-15	144	0.85	0.00991	0.00011	0.06672	0.00433	0.00312	0.00008	64. 0	1. 1	65.6	4. 1	
44-16	271	0.62	0.00969	0.00011	0.06060	0.00433	0.00307	0.00010	62. 4	1. 1	59.7	6.0	
44-17	645	0. 95	0.00970	0.00008	0.06241	0.00150	0.00299	0.00004	63. 1	1. 1	68.3	5.8	
44-18	257	0.65	0.00973	0.00008	0.06708	0.00299	0.00299	0.00006	63.0	1.0	70.0	4. 7	
44-19	297	0. 93	0.00975	0.00007	0.06845	0.00262	0.00311	0.00005	63.0	1.0	65.3	4. 5	
44-20	145	0.76	0.00950	0.00014	0.06860	0.00375	0.00314	0.00008	60. 9	0.9	59. 5	4. 6	
44-21	192	0.63	0.00959	0.00012	0.06642	0.00382	0.00303	0.00011	61.8	1.2	65.0	5. 9	
44-22	144	0.38	0.01001	0.00020	0. 10251	0.01141	0.00624	0.00048	60. 9	1.2	33. 3	5.8	剔除
44-23	237	0.69	0.00998	0.00009	0.06919	0.00312	0.00343	0.00007	63.8	1.5	57. 4	14. 2	
44-24	117	0.56	0.00937	0.00035	0.06311	0.00823	0.00305	0.00021	60. 1	1.0	57. 1	4. 2	
44-25	287	0.60	0.00981	0.00008	0.06365	0.00204	0.00316	0.00006	63. 2	2.4	61.4	11.6	
44-26	302	0.61	0.00966	0.00009	0.06824	0.00323	0.00314	0.00007	62. 2	1.0	65.0	4. 1	
44-27	172	0.86	0.01004	0.00012	0.06141	0.00393	0.00323	0.00007	64. 7	1.0	57. 1	5.0	
44-28	398	0. 97	0.00993	0.00011	0.06348	0.00294	0.00304	0.00007	64. 5	1. 2	68. 4	7. 0	
44-29	289	0.59	0.00968	0.00007	0.06509	0.00223	0.00307	0.00006	62. 4	1. 1	63.6	6. 1	
44-30	78	0.44	0.00981	0.00016	0.08441	0.00628	0.00312	0.00016	63. 1	1.0	81.5	3.4	
44-31	242	0.75	0.00971	0.00011	0.06786	0.00347	0.00317	0.00007	62.4	1.3	61.8	7. 0	

表 3 多底沟矽卡岩型钼矿床(点)辉钼矿 Re-Os 年龄分析结果

Table 3 Re-Os data of molybdenite from the Duodigou skarn deposit

Sample	$\frac{\text{Re}}{(\times 10^{-6})}$	187 Re ($\times 10^{-6}$)	187 Os (×10 $^{-9}$)	模式年龄 (Ma)	ΔT (Ma)
LHY-1	1. 40395	0. 878873	1. 00405	68. 5	0. 7
LHY-40	2. 16881	1.35768	1.56723	69. 2	3.3
LHY-80	2. 41121	1. 50942	1. 61907	64. 3	0.8

等,2011)形成时代都在65Ma左右(表4),与本文报道的甲龙铁矿、撒当金银矿床(点)、多底沟钼矿床(点)同位素年龄相近。这些矿床的形成也和印度板块与欧亚板块碰撞早期构造岩浆事件有关。

印度板块与欧亚板块碰撞早期与中酸性岩浆活动有关

的矿床类型及元素组合较多,有矽卡岩型铁矿床、矽卡岩型铅锌钼矿床、斑岩-矽卡岩型铜(钼)铅锌矿床及火山岩型银金矿床;碰撞早期矿床分布面积较大,在空间上在冈底斯带中南部形成了一条东西向伸展近800km、南北宽近80km的成矿带(图4);碰撞早期岩浆活动也形成大型-特大型矿床,如亚贵拉铅锌钼矿床达特大型规模(黄克贤等,2012),恰功铁矿达大型规模(王方国等,2005)。这表明印度板块与欧亚板块碰撞早期在冈底斯中南部发生了与中酸性岩浆活动有关的较大规模成矿作用,形成一系列不同元素组合矿床。

冈底斯南部碰撞早期岩浆活动强烈,发育广泛分布的林子宗组火山岩及冈底斯花岗岩岩基(Zhou et al., 2004;董国臣,2002;莫宣学等,2005;潘桂棠等,2006)。冈底斯带古新世-始新世侵入岩和火山岩多为钙碱性,富集大离子亲石元素(LILE)和相对亏损高场强元素(HFSE),具岛弧岩浆岩的一般特征(莫宣学等,2003)。目前对碰撞早期岩浆形成

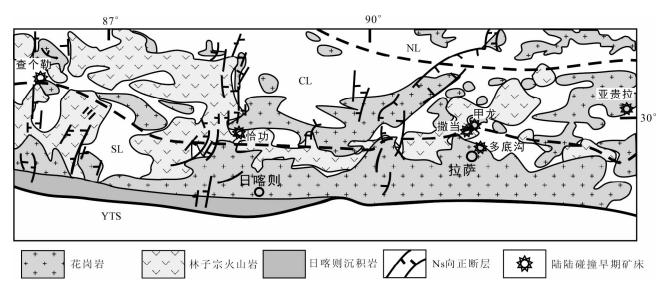


图 4 冈底斯带南缘地质及碰撞早期矿床布简图(据 Zheng et al., 2012 修改)

NL-北拉萨微地体; CL-中拉萨微地体; SL-南拉萨微地体; YTS-雅鲁藏布缝合带

Fig. 4 Simplified geological map of the southern Gangdese ore belt, Tibet (modified after Zheng et al., 2012)

NL-northern Lhasa subterrane; CL-central Lhasa subterrane; SL-southern Lhasa subterrane; YTS-Yarlung-Tsangpo suture

表 4 印度-欧亚板块碰撞早期矿床

Table 4 The deposits formed in the early stage of India-Eurasia collision

矿床名称	矿化类型	赋矿岩体	成矿年代	资料来源	
甲龙铁矿	矽卡岩型铁矿化	斑状黑云母二长花岗岩与 地层接触带	61.1 ± 0.4Ma(锆石 LA-ICP-MS)		
撒当金银矿床	火山岩型金银矿化	安山岩	略晚于62.6±0.5Ma(锆石LA-ICP-MS)	本文	
多底沟钼矿床	矽卡岩型钼矿化	角闪黑云母二长花岗岩	66.7 ± 6.4 Ma(辉钼矿 Re-Os 加权平均模式年龄)		
查个勒铜 铅锌矿床	矽卡岩型铜铅锌矿化	花岗斑岩与围岩接触带及 构造破碎带	61.5±0.6Ma(辉钼矿 Re-Os 等时线) 62.1±1.1Ma(锆石 LA-ICP-MS)	高顺宝等,2012	
TI 17 19 1/N	斑岩型铜钼矿化	花岗斑岩	63.3 ± 0.9Ma(锆石 LA-ICP-MS)		
亚贵拉铅 锌钼矿床	矽卡岩型钼矿化,局部伴 生铅锌矿化	石英斑岩与地层接触带的 角岩、大理岩、砂卡岩等	65.0±1.9Ma(辉钼矿 Re-Os 等时线) 62.4±0.6Ma(锆石 LA-ICP-MS)	高一鸣等,2011; 黄克贤等,2012;	
恰功铁矿	矽卡岩型铁矿化	二长花岗斑岩与灰岩的接 触带	66.8±0.7Ma(锆石 LA-ICP-MS) 67.4±0.8Ma(锆石 LA-ICP-MS)	李应栩等,2011;	

过程主要有下列两种观点:(1)向北俯冲雅江洋板块发生回转或断离,软流圈上涌,诱发拉萨地体下地壳发生部分熔融(Chung et al., 2005; Ji et al., 2009);(2)雅江洋板块向北俯冲过程中地幔楔发生部分熔融形成的基性岩浆底侵至加厚的下地壳并诱发其部分熔融(Mo et al., 2005, 2007;董国臣等, 2006)。

主碰撞中晚期沙让斑岩钼矿床(辉钼矿 Re-Os 年龄: 51.0±1.0Ma, Zhao et al., 2012)成矿斑岩微量元素组成特征表明成矿岩浆形成时冈底斯带中部地体没有明显加厚(侯增谦等, 2012)。主碰撞早期与中酸岩浆作用有关矿床主要位于冈底斯带中部北纬30°线附近(图 4),和沙让斑岩钼矿床处于同一纬度区域,因此,主碰撞早期成矿岩浆形成应不是加厚下地壳熔融形成的。印度板块与欧亚板块在约70~

65Ma 时发生初始碰撞,因此,冈底斯中带碰撞早期中酸性岩体及有关矿床可能是向北俯冲的雅江洋板块在 70~65Ma 发生岛弧碰撞时受阻,俯冲洋壳由于推动力减弱而发生调整,俯冲洋壳前缘(可能位于北纬 30°线附近)发生下沉(回转)、后退或断离,软流圈上涌导致冈底斯带中部下地壳物质熔融形成的岩浆,经结晶分异形成的。

冈底斯带中部碰撞早期中酸性岩浆活动强烈,现有工作表明,碰撞早期中酸性岩浆活动伴随一系列成矿作用,形成一系列不同元素组合的矿床。不同元素组合矿床成岩成矿过程有一定的差异,因此,冈底斯带在碰撞早期发生了复杂多样的岩浆成矿事件,有形成碰撞早期与中酸性岩浆活动有关矿床的良好成矿条件,今后应注意在冈底斯带中部地区碰撞早期中酸性岩浆作用有关矿床的找矿工作。

6 结论

- (1)甲龙砂卡岩型铁矿斑状黑云母二长花岗岩的锆石 LA-ICP-MS U-Pb 年龄为 61.1 ± 0.4Ma, MSWD = 0.94; 撒当 赋金银床(点)矿化安山岩锆石 LA-ICP-MS U-Pb 年龄为 62.6 ± 0.5Ma, MSWD = 1.51; 多底沟钼矿床(点)辉钼矿 Re-Os 模式年龄为 64.3 ± 0.8Ma ~ 69.2 ± 3.3Ma, 加权平均模式年龄为 66.7 ± 6.4Ma(MSWD = 8.1)。三个矿床(点)与印度-亚洲板块碰撞早期构造岩浆活动有关;
- (2)冈底斯中部陆陆碰撞早期岩浆作用伴随大规模成矿作用,形成一系列与陆陆碰撞早期中酸性岩浆作用有关的矿床。这些矿床的形成可能与俯冲洋壳在70~65Ma发生岛弧碰撞时受阻,俯冲洋壳前缘由于推动力减弱下沉(回转)或断离诱发岩浆活动有关。

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